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- 24 Table S1: Location and record lengths for the 11 charcoal data records from western
- 25 North America used in this study.

				Record
				Length
Site ID	Site Name	Lat	Lon.	(yrs)
7	Foy	48.17	-114.36	13180
8	Bolan	42.02	-123.46	14545
13	Trail	44.28	-110.17	8050
128	Lily Lake	41.98	-120.21	12214
134	Hunters Lake	37.61	-106.84	14260
220	Ruppert	67.07	-154.25	14000
225	7-M	62.50	-113.72	6878
268	Cooley	49.49	-117.65	7551
1077	Lower Gaylor Lake	37.91	-119.29	11679
1129	Todd Lake	44.03	-121.68	7770
1146	Sanger Lake CA	41.90	-123.65	14465

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32 Figure S1: Representation of the regions as defined for the GFED4s and in this study.

- 33 BONA: Boreal North America, TENA: Temperate North America, CEAM: Central
- 34 America, NHSA: Northern Hemisphere South America, SHSA: Southern Hemisphere
- 35 South America, EURO: Europe, MIDE: Middle East, NHAF: Northern Hemisphere
- 36 Africa, SHAF: Southern Hemisphere Africa, BOAS: Boreal Asia, CEAS: Central Asia,
- 37 SEAS: Southeast Asia, EQAS: Equatorial Asia, AUST: Australia.



*Figure S2: Power spectra of NINO3 SSTs computed using Morlet wavelet analysis with* wavenumber-6 with the period expressed as octaves of the annual cycle. Spectra are computed for non-overlapping a) 20-year segments, b) 100-year segments and c) 500-year segments for ESM2Mb and d) 20-year segments, e) 100-year segments, and f) 450-year segments for CESM1. The thick black line gives the average spectrum for the full preindustrial control runs for the respective ESM. The orange line shows the average spectrum for the HadISST1 dataset (Rayner et al., 2003) for 1870 to 2015. 





Figure S3: Composite analysis showing differences in the means of standardized
anomalies of precipitation for years with positive NINO3 minus years with negative
NINO3 for annual precipitation (a: ESM2Mb, f: CESM1.1, k: CRU TS 3.10 [Harris et
al., 2014]/HadISST1 [Rayner et al., 2003] 1901-2009), DJF precipitation (b: ESM2Mb,
g: CESM1.1, l: CRU TS 3.10/HadISST1), MAM precipitation (c: ESM2Mb, h: CESM1.1,
m: CRU TS 3.10/HadISST1), JJA precipitation (d: ESM2Mb, i: CESM1.1, n: CRU TS

58	3.10/HadISST1), SON precipitation (e: ESM2Mb, j: CESM1.1, o: CRU TS
59	3.10/HadISST1).
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*Figure S4: The interannual variability of precipitation from a) CESM1.1 and b)* 

- 83 ESM2Mb compared to the CRU TS3.10 (Harris et al., 2014) 1976-2005 precipitation
- 84 (green circles).



91 Figure S5: Fire emission anomalies  $[kgC m^{-2}]$  from the GFED4s for August 1997 to

92 September 1998 relative to the entire GFED4s record, 1997 to 2014.





99 Figure S6: Composite analysis showing differences in the means of standardized

100 anomalies of annual fire emissions for years with positive, 50-year low-pass filtered

101 AMO indices minus years with negative, 50-year low-pass filtered AMO indices from

102 *CESM1.1.* The values of the difference in means that are significant at a 95% confidence

103 *level (two-tailed test) are shown in the colorbars as pink lines.* 

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- 108 Datasets
- 109 1. GFEDv4s
- 110 The Global Fire Emissions Database version 4s (GFED4s) is based on a burned area
- 111 dataset derived primarily from MODIS (Giglio et al., 2013). Carbon emissions from fires

112 are modeled using the burned area data applied within the Carnegie-Ames-Stanford-113 Approach (CASA) terrestrial model (Potter et al., 1993; Field et al., 1995; Randerson et 114 al., 1996), which simulates carbon cycling through different pools, including plant litter, 115 and important fire-related processes such as fire-caused vegetation mortality and 116 combustion completeness (Van der Werf et al., 2010). Fires may be consistently 117 underestimated in this inventory in areas with persistent cloud cover that hides active 118 fires from the satellite sensors, and in croplands where fires are often small and escape 119 detection (Van der Werf et al., 2010). An accounting of the emissions from small fires 120 has been made by statistical means (Randerson et al., 2012) and these additional fire 121 emissions are included in the data shown in our study. Fires associated with 122 deforestation and peatlands are isolated from grassland and natural forest fires with 123 ancillary datasets (Giglio et al., 2013).

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## 125 2. GFASv1

126 The Global Fire Assimilation System (GFASv1) also uses MODIS to determine global 127 fire activity, but is based on a fire radiative power product instead of burned area (Kaiser 128 et al., 2012). Land-cover dependent conversion factors translate the radiative power into 129 dry-matter combustion rates from which carbon emissions can be derived. The 130 conversion factors are based on theory (Wooster et al., 2005) but scaled to improve the 131 match of the GFAS emissions to GFEDv3 (Kaiser et al., 2012). Compared to GFEDv3, 132 GFASv1 exhibits larger areas of small fire emissions, an aspect that may have been 133 compensated somewhat by the inclusion of small fires in the more recent GFED4s.

135 **3. FINNv1.5** 

136 The Fire Inventory from NCAR (FINN) provides higher time (daily) and spatial (~1km) 137 resolution that the previous two inventories (Wiedinmyer et al., 2011), although we re-138 grid the data to more coarse resolution and use monthly and annual averages for this 139 study. FINN combines active fire counts and land cover data from MODIS to compute 140 total fire emissions from each square kilometer where a fire is detected. Fuel loadings and 141 fraction of vegetation that is burned by fires are derived from MODIS land cover and 142 vegetation type products. In comparison to GFEDv3.1, Wiedinmyer et al. (2011) found 143 that FINN had generally higher emissions from southern hemisphere South America, 144 substantially higher emissions from southeastern Asia, with lower emissions in Africa, 145 Australia and boreal North America. The two inventories report similar global average 146 fire emissions for the period 2005-2009 as examined by Wiedenmyer et al. (2011).

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## 148 **4. Charcoal sediments**

149 Charcoal sediments are often used as a proxy for local and sometimes regional fire 150 emissions and can produce datasets that extend tens of thousands of years into the past 151 (Marlon et al., 2008). Here we make use of data from the Global Charcoal Database 152 (www.paleofire.org) (Power et al., 2008), which is a collection of over 1000 sediment 153 cores taken from six continents. The time resolution of each core, determined by the rate 154 of sedimentation characteristic to the core site, and the period of record vary substantially 155 between sites. Cores with a combination of lengthy record and high time resolution are 156 uncommon. For this study we selected sites with a length of record of at least 4,000 years 157 and time resolution of less than 20 years between timesteps. Western North America had

158	the largest amount of cores that fit these criteria by a large margin (sites listed in Table
159	S2). The data are prepared following the methodology of Marlon et al. (2008) and
160	programming of Blarquez et al. (2014) after linear interpolation to a 20-year timestep.
161	All concentration data are converted to charcoal influx data and undergo Box-Cox
162	transformation and standardization so that different records can be averaged together.
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164	5. Data availability
165	GFEDv4s: http://www.globalfiredata.org/data.html
166	GFASv1: http://apps.ecmwf.int/datasets/data/cams-gfas/
167	FINNv1.5: http://bai.acom.ucar.edu/Data/fire/
168	GCD: https://www.paleofire.org
169	CRU TS3.10: http://catalogue.ceda.ac.uk/uuid/3f8944800cc48e1cbc29a5ee12d8542d
170	HadISST1: http://www.metoffice.gov.uk/hadobs/hadisst/data/download.html
171	
172	The CESM Large Ensemble Community Project (LENS) has made their ESM output
173	publicly available and interested users are directed to begin at this website:
174	http://www.cesm.ucar.edu/projects/community-projects/LENS/
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