

Understanding the double peaked El Niño in coupled GCMs

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Abstract Coupled general circulation models (CGCMs) simulate a diverse range of El Niño–Southern Oscillation behaviors. “Double peaked” El Niño events—where two separate centers of positive sea surface temperature (SST) anomalies evolve concurrently in the eastern and western equatorial Pacific—have been evidenced in Coupled Model Intercomparison Project version 5 CGCMs and are without precedent in observations. The characteristic CGCM double peaked El Niño may be mistaken for a central Pacific warming event in El Niño composites, shifted westwards due to the cold tongue bias. In results from the Australian Community Climate and Earth System Simulator coupled model, we find that the western Pacific warm peak of the double peaked El Niño event emerges due to an excessive westward extension of the climatological cold tongue, displacing the region of strong zonal SST gradients towards the west Pacific. A coincident westward shift in the zonal current anomalies reinforces the western peak in SST anomalies, leading to a zonal separation between the warming effect of zonal advection (in the west Pacific) and that of vertical advection (in the east Pacific). Meridional advection and net surface heat fluxes further drive growth

of the western Pacific warm peak. Our results demonstrate that understanding historical CGCM El Niño behaviors is a necessary precursor to interpreting projections of future CGCM El Niño behaviors, such as changes in the frequency of eastern Pacific El Niño events, under global warming scenarios.

Keywords El Niño evolution · Coupled general circulation model · CMIP5 · Cold tongue bias · Climate change

1 Introduction

Coupled General Circulation Models (CGCMs) are among our most effective tools for investigating the dynamics of El Niño–Southern Oscillation (ENSO) and the response of ENSO to global warming (Meehl et al. 2006; Yeh et al. 2006, 2009; Collins et al. 2010; Vecchi and Wittenberg 2010). Improvements are continually being made to these models to better represent the salient features of ENSO, such as its amplitude, frequency, seasonality, and stability (AchutaRao and Sperber 2006; Guilyardi et al. 2009; Deser et al. 2012; Guilyardi et al. 2012, 2013; Bellenger et al. 2014; Kim et al. 2014; Guilyardi et al. 2015). Nevertheless, there is considerable diversity in the simulation of ENSO dynamics, both within and across CGCMs (Capotondi et al. 2006; Lloyd et al. 2009; Belmadani et al. 2010; Ham and Kug 2012; Lloyd et al. 2012; Brown et al. 2013; Capotondi 2013; Capotondi et al. 2015a, b; Choi et al. 2015) and even more diversity in how ENSO will change under global warming (Leloup et al. 2008; Guilyardi et al. 2009; Collins et al. 2010; Boucharel et al. 2011; Kim and Jin 2011; DiNezio et al. 2012; Watanabe et al. 2012; Taschetto et al. 2014; Latif et al. 2015).

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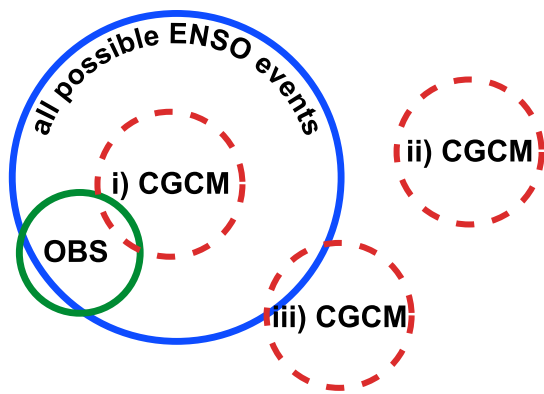


Fig. 1 Venn diagram representing possible relationships between the ENSO behaviors simulated by CGCMs (*red dashed circles*), the full range of possible ENSO behaviors under present-day conditions (*blue circle*) and the observed ENSO behaviors (*green circle*). The *green circle* extends slightly outside the *blue circle* to represent observational errors, such as in measurement or reconstruction. We also note that the *blue circle* is itself evolving on decadal to centennial time-scales due to natural internal variability, as well as due to external radiative forcings

It follows that a current focus of ENSO research is in quantifying the realism of behaviors simulated by CGCMs, which requires comparison of model output with observed features such as sea surface temperature (SST), winds, rainfall, clouds, mixed layer depth, thermocline depth, and ocean currents. However, we have glimpsed only a sample of the possible ENSO behaviors and spatial diversity that could occur (Fig. 1). This is at least partly due to the fact that ENSO modulates climate on multiple timescales, demonstrating strong interannual variability as well as decadal to multidecadal variability (Allan 2000; Allan et al. 2003; Wittenberg 2009; Kug et al. 2010; Choi et al. 2012; Ogata et al. 2013; Meehl et al. 2013; Holbrook et al. 2014; Lee et al. 2014; Wittenberg et al. 2014; Wittenberg 2015), and such long-term variability may not yet be clearly distinguishable from our relatively short observational record. The framework schematized in Fig. 1 presents three of the possible scenarios for the range of ENSO behaviors evidenced in CGCMs: (i) CGCMs simulate realistic behaviors, of which some may mirror the observations; (ii) CGCMs are unable to simulate present-day ENSO behaviors; or (iii) CGCMs capture behaviors that are qualitatively similar to those of the real world as well as some unrealistic ones. [Additional scenarios to these three discussed here are possible, such as the observational or reanalysis data exhibiting biases in their representation of reality, as well as the real-world variability changing due to external radiative forcings.] Scenario (i) is desirable if we are to use CGCMs to understand future externally forced ENSO events, while scenario (ii) implies little faith in the ability of coupled models to perform this task. Based on results from recent

studies (e.g., Wittenberg et al. 2006; Guilyardi et al. 2009; Brown et al. 2013) scenario (iii) is perhaps the most likely, indicating that while CGCMs are useful, their underlying biases should be taken into consideration when interpreting simulated ENSO behaviors.

Observations suggest that there is a continuum of El Niño spatial diversity in warming, with centers of action located from the eastern equatorial Pacific to the central equatorial Pacific (Giese and Ray 2011; Johnson 2013; Capotondi et al. 2015b). A recent trend classifies El Niño events as “eastern Pacific” events or “central Pacific” events depending on the location of maximum sea surface temperature warming at the height of the El Niño event (Ashok et al. 2007; Kao and Yu 2009; Yeh et al. 2009; Lee and McPhaden 2010; Yu and Kim 2013; Yeh et al. 2014). [Although, these classifications are qualitative descriptors of diversity, rather than being indicative of different modes of spatial variability; Capotondi et al. 2015b.] Nevertheless, the patterns of warming simulated by CGCMs do not necessarily closely align with those of observations or flux-forced ocean general circulation models (OGCMs). For instance, while observations and OGCMs show strong, and relatively continuous, variability in SST anomalies (SST') along the equator from the central to the eastern Pacific, the pattern of SST' in CGCMs is split into two separate centers of action, in the western-central and eastern Pacific (Fig. 2). The western-central Pacific peak of warm SST' in Fig. 2, or indeed in composites of El Niño SST' , might be interpreted as the CGCM analog of the central Pacific El Niño event, whose center of action is shifted westwards due to the cold tongue bias (Wittenberg et al. 2006; Kao and Yu 2009; Yeh et al. 2009; Ham and Kug 2012; Taschetto et al. 2014). However, systematic inspection of the evolution of CGCM El Niño events reveals a “double peaked” pattern of warming in CGCMs—with two warm peaks developing

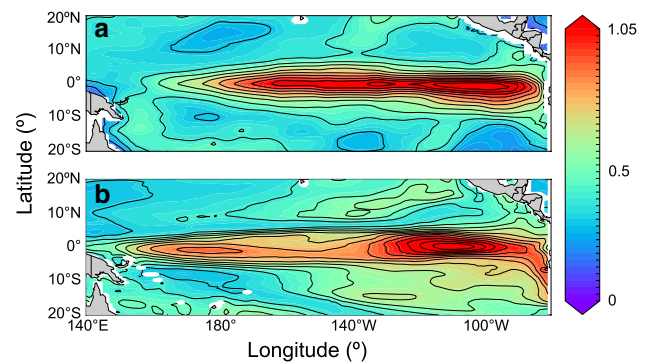
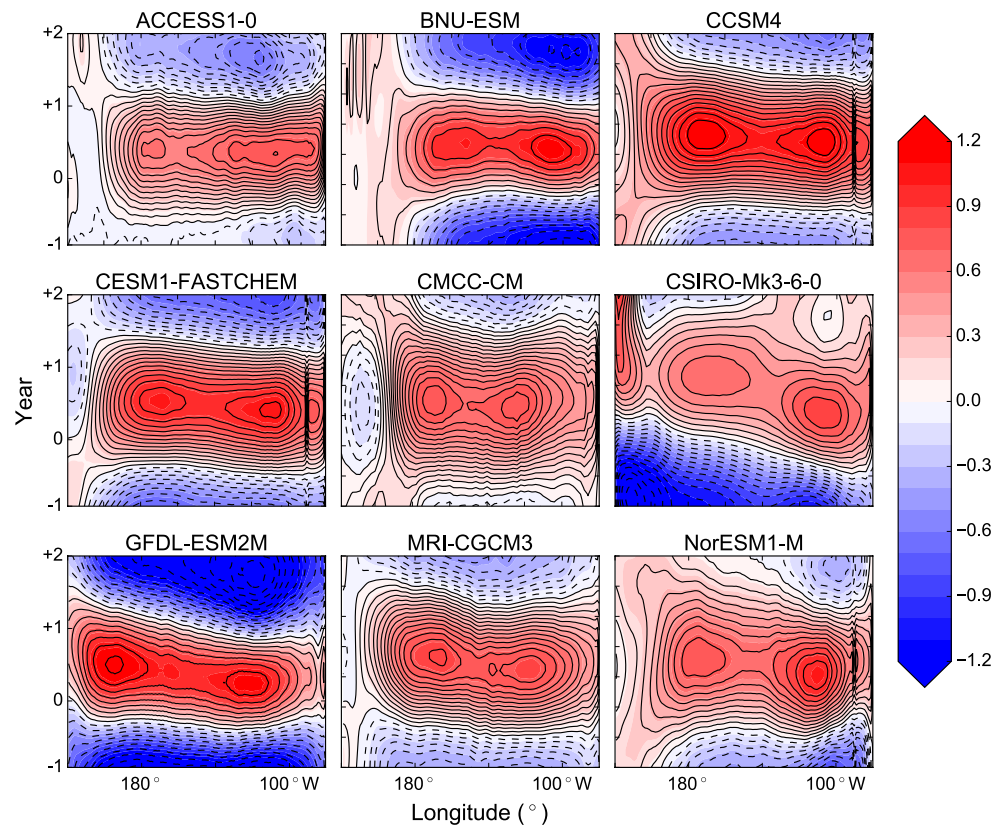


Fig. 2 Standard deviation of sea surface temperature anomalies (*shading*) in the **a** Bureau of Meteorology Research Centre SST reanalyses (Smith 1995), and **b** historical simulation of ACCESS-CM1.3. Data are in units of °C and the contour interval is 0.025 °C. Contours of the standard deviation at the 0.075 °C interval are overlaid

Fig. 3 Examples of the evolution of SST' for the 36 months surrounding composite double peaked El Niño events from *historical* simulations of nine CMIP5 models (as indicated). Data are in units of $^{\circ}\text{C}$ (with contour intervals of 0.1°C) and are averaged over 2°S – 2°N



concurrently in the eastern and central Pacific (e.g., Fig. 3). This double peaked El Niño event is common in Coupled Model Intercomparison Project version 5 (CMIP5) CGCMs (Fig. 4). A double peaked structure was also evident in the SST' variance of CMIP3 models, e.g., the CSIRO-Mk3.0 model (figure 1 of Capotondi et al. 2006).

The spatial structure of SST' is essential for determining the atmospheric response to ENSO. This is especially the case near the convectively-active region of the western Pacific warm pool, where subtle variations in SST can have large impacts on the location and intensity of atmospheric latent heating, and thereby the global atmospheric circulation. This in turn affects not only the feedbacks critical to ENSO (Choi et al. 2013, 2015), but also the structure of the atmospheric stochastic forcing (Vecchi et al. 2006b; Gebbie et al. 2007), and ENSO's remote teleconnections (Capotondi et al. 2015b; Jia et al. 2015; Yang et al. 2015; Krishnamurthy et al. 2015, 2016; Zhang et al. 2016). Thus it is important to assess and understand the biases that CGCMs have in their spatial pattern of SST' during ENSO, as well as how those biases affect ENSO behavior, remote impacts, and ENSO sensitivities to climate change.

The goal of this paper is to investigate the behavior of the CGCM double peaked El Niño event, including the mechanisms that underlie its development. We further seek to address whether the CGCM double peaked El Niño event is a realistic and likely representation of El Niño

spatial diversity, or an artifact of coupled model biases. In Sect. 2 we introduce the data and techniques used to identify and analyze the double peaked El Niño event. Section 3 presents analysis of the double peaked El Niño event in the CMIP5 suite of CGCMs. The dynamics giving rise to the double peaked event are examined in the context of the Australian Community Climate and Earth System Simulator Coupled Model version 1.3 (Bi et al. 2013a). The results are discussed and summarized in Sect. 4.

2 Data and methods

2.1 CMIP5 CGCMs

We analyze the evolution of SST' during double peaked El Niño events in *pre-industrial control* (*PiControl*) and *historical* simulations of 36 climate models submitted to the Coupled Model Intercomparison Project Phase 5 (CMIP5) database (Table 1). The nomenclature of the terms “*PiControl*” and “*historical*” follow Taylor et al. (2012). *PiControl* simulations attempt to capture the preindustrial climate equilibrium state and are simulated over several hundreds of years; *historical* simulations represent forced runs using observed atmospheric composition changes (atmospheric forcing from both natural and anthropogenic sources) from the mid-19th Century to near present day.

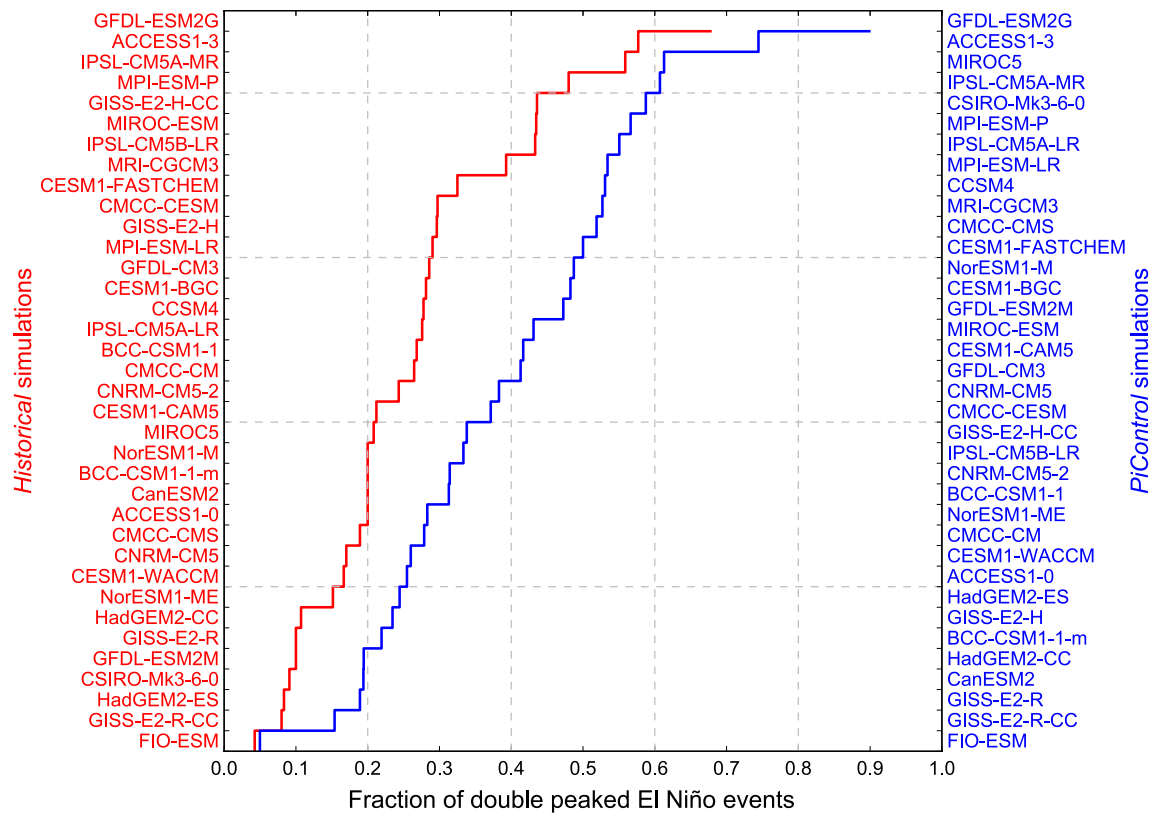


Fig. 4 Fraction of all El Niño events that are classified as double peaked events from *PiControl* and *historical* simulations of 36 CMIP5 CGCMs. Data in *blue* corresponds to *PiControl* simulations; data in *red* corresponds to *historical* simulations. The distributions

of the *PiControl* and *historical* data are statistically significantly different at the 99 % confidence level using a two-side Kolmogorov–Smirnov test (test statistic 0.44, $p = 0.001$)

To diagnose the likely mechanisms underpinning the double peaked El Niño event, monthly anomalies of SST, and all variables analyzed from the Australian Community Climate and Earth System Simulator (ACCESS) simulation, are computed by subtracting the annual cycle from the monthly mean outputs. The data are smoothed using a 13-point Parzen filter to remove frequencies of sub-annual variability.

2.2 The ACCESS model

To investigate the mechanisms underpinning the CGCM double peaked El Niño events, we analyze a *PiControl* 505-year simulation of ACCESS version 1.3 (ACCESS-CM1.3). The ocean component of the ACCESS-CM1.3 simulation is an OGCM that draws its codebase and most of its configuration from the NOAA Geophysical Fluid Dynamics Laboratory (GFDL) Modular Ocean Model version 4 (MOM4p1) (Griffies 2009). A full description of the ACCESS component models can be found in Bi et al. (2013a, b), and their implementation is described by Dix et al. (2014).

ACCESS-CM1.3 has been tested against various benchmarks important for the simulation of ENSO, finding that its performance and the magnitude of its model biases are comparable to other CMIP5 models (Brown et al. 2013; Rashid et al. 2013a, b; Kim et al. 2014; Taschetto et al. 2014; Rashid and Hirst 2015). The mean state and biases of the simulated tropical Pacific in ACCESS-CM1.3 are further discussed in Appendix 1. The ocean component of ACCESS-CM1.3 was previously analyzed in Graham et al. (2015).

2.3 Defining El Niño events

El Niño events are defined when a 5-month running mean of the unfiltered SST' in the Niño-3.4 region (5°S – 5°N , 170 – 120°W) exceeds 0.4°C for a period of at least 6 months (Trenberth 1997). If a single center of SST' warming is isolated to the eastern equatorial Pacific (east of approximately 160°W), the event is classified as an eastern Pacific El Niño. The following method is used to distinguish double peaked El Niño events. Locations of maximum warming along the equator (2°S – 2°N) are

Table 1 Summary of the El Niño spatial behaviors present in CMIP5 CGCMs based on the evolution of SST'

Model	References	Years	El Niño events	DP events	DWPE isotherm	Mean longitude (σ)	
						DWPE ($^{\circ}$ E)	Western peak ($^{\circ}$ E)
ACCESS1-0	Bi et al. (2013a)	156 500 ⁺	28 106 ⁺	6 27 ⁺	28.9 28.2 ⁺	173 (9.2) 167 ⁺ (8.4 ⁺)	189 (7.3) 184 ⁺ (8.5 ⁺)
ACCESS1-3	Bi et al. (2013a)	156 500 ⁺	26 89 ⁺	15 65 ⁺	29.2 28.7 ⁺	163 (9.1) 163 ⁺ (8.3 ⁺)	180 (6.3) 176 ⁺ (12.6 ⁺)
BCC-CSM1-1	Wu et al. (2010)	163 500 ⁺	37 115 ⁺	11 36 ⁺	28.2 28.1 ⁺	164 (11.2) 171 ⁺ (9.8 ⁺)	191 (13.2) 188 ⁺ (15.4 ⁺)
BCC-CSM1-1-M	Wu et al. (2010)	163 400 ⁺	46 114 ⁺	10 47 ⁺	29.0 28.9 ⁺	177 (13.5) 172 ⁺ (14.7 ⁺)	196 (11.6) 182 ⁺ (13.1 ⁺)
CanESM2	Arora et al. (2011); Gillett et al. (2012)	156 996 ⁺	36 268 ⁺	8 52 ⁺	28.6 27.7 ⁺	176 (13.9) 168 ⁺ (10.9 ⁺)	175 (9.2) 174 ⁺ (17.9 ⁺)
CCSM4	Gent et al. (2011)	156 1051 ⁺	36 228 ⁺	10 121 ⁺	29.1 28.6 ⁺	182 (14.7) 194 ⁺ (11.6 ⁺)	187 (18.7) 184 ⁺ (14.2 ⁺)
CESM1-BGC	Long et al. (2013)	156 500 ⁺	29 114 ⁺	9 55 ⁺	29.1 28.5 ⁺	185 (12.0) 186 ⁺ (15.9 ⁺)	191 (10.5) 189 ⁺ (8.1 ⁺)
CESM1-CAM5	Meehl et al. (2013)	156 319 ⁺	33 60 ⁺	7 25 ⁺	28.4 28.1 ⁺	175 (14.6) 175 ⁺ (13.7 ⁺)	187 (5.4) 177 ⁺ (9.0 ⁺)
CESM1-FASTCHEM	Meehl et al. (2013)	156 222 ⁺	38 48 ⁺	13 24 ⁺	29.1 29.1 ⁺	172 (12.2) 177 ⁺ (13.8 ⁺)	192 (9.8) 196 ⁺ (16.1 ⁺)
CESM1-WACCM	Meehl et al. (2013)	156 200 ⁺	34 50 ⁺	6 13 ⁺	28.8 28.6 ⁺	173 (14.7) 186 ⁺ (15.8 ⁺)	191 (9.1) 197 ⁺ (14.4 ⁺)
CMCC-CESM	Fogli et al. (2009)	156 277 ⁺	37 70 ⁺	11 26 ⁺	29.2 28.3 ⁺	180 (14.0) 181 ⁺ (13.7 ⁺)	182 (10.1) 186 ⁺ (12.6 ⁺)
CMCC-CM	Fogli et al. (2009)	156 330 ⁺	30 61 ⁺	9 17 ⁺	28.9 28.7 ⁺	172 (15.6) 179 ⁺ (17.5 ⁺)	192 (10.1) 198 ⁺ (19.1 ⁺)
CMCC-CMS	Fogli et al. (2009)	156 500 ⁺	37 106 ⁺	7 55 ⁺	28.4 28.7 ⁺	181 (14.2) 179 ⁺ (10.9 ⁺)	194 (7.5) 194 ⁺ (12.0 ⁺)
CNRM-CM5	Voltaire et al. (2013)	156 850 ⁺	41 222 ⁺	8 85 ⁺	28.4 28.5 ⁺	161 (11.6) 170 ⁺ (13.0 ⁺)	183 (23.2) 186 ⁺ (16.2 ⁺)
CNRM-CM5-2	Voltaire et al. (2013)	156 410 ⁺	36 105 ⁺	9 33 ⁺	28.2 28.0 ⁺	164 (12.8) 165 ⁺ (13.5 ⁺)	187 (7.2) 183 ⁺ (16.0 ⁺)

Table 1 continued

Model	References	Years	El Niño events	DP events	DWPE isotherm	Mean longitude (σ)	
						DWPE ($^{\circ}$ E)	Western peak ($^{\circ}$ E)
CSIRO-Mk3-6-0	Rotstayn et al. (2012)	156 500 ⁺	29 80 ⁺	3 47 ⁺	27.8 27.0 ⁺	155 (6.1) 152 ⁺ (5.3 ⁺)	189 (0.9) 193 ⁺ (10.8 ⁺)
FIO-ESM	Qiao et al. (2013)	156 800 ⁺	45 241 ⁺	2 12 ⁺	28.9 29.2 ⁺	179 (19.7) 180 ⁺ (19.7 ⁺)	204 (12.4) 195 ⁺ (16.9 ⁺)
GFDL-CM3	Donner et al. (2011); Griffies et al. (2011)	146 800 ⁺	40 184 ⁺	12 76 ⁺	28.7 28.4 ⁺	166 (13.2) 164 ⁺ (10.8 ⁺)	167 (18.9) 169 ⁺ (20.7 ⁺)
GFDL-ESM2G	Dunne et al. (2012)	145 500 ⁺	27 89 ⁺	21 80 ⁺	28.3 28.2 ⁺	156 (6.6) 156 ⁺ (7.7 ⁺)	159 (18.3) 156 ⁺ (19.1 ⁺)
GFDL-ESM2M	Dunne et al. (2012)	145 500 ⁺	30 91 ⁺	3 43 ⁺	28.6 28.8 ⁺	174 (19.3) 172 ⁺ (15.8 ⁺)	166 (19.1) 176 ⁺ (8.0 ⁺)
GISS-E2-H	Miller (2014); Schmidt et al. (2014)	156 1770 ⁺	24 162 ⁺	8 38 ⁺	28.4 28.9 ⁺	167 (14.0) 183 ⁺ (19.4 ⁺)	176 (28.9) 186 ⁺ (18.2 ⁺)
GISS-E2-H-CC	Miller (2014); Schmidt et al. (2014)	161 251 ⁺	37 71 ⁺	17 24 ⁺	28.4 28.9 ⁺	173 (17.1) 178 ⁺ (18.7 ⁺)	194 (26.6) 194 ⁺ (22.5 ⁺)
GISS-E2-R	Miller (2014); Schmidt et al. (2014)	156 1200 ⁺	28 148 ⁺	3 28 ⁺	28.9 28.7 ⁺	191 (17.9) 174 ⁺ (16.1 ⁺)	214 (16.4) 199 ⁺ (9.8 ⁺)
GISS-E2-R-CC	Miller (2014); Schmidt et al. (2014)	161 251 ⁺	24 39 ⁺	2 6 ⁺	28.9 28.8 ⁺	188 (18.7) 175 ⁺ (16.4 ⁺)	184 (8.8) 194 ⁺ (11.5 ⁺)
HadGEM2-CC	Collins et al. (2011); Martin et al. (2011)	146 240 ⁺	28 36 ⁺	3 7 ⁺	27.9 27.7 ⁺	171 (8.9) 171 ⁺ (8.5 ⁺)	206 (8.1) 179 ⁺ (0.0 ⁺)
HadGEM2-ES	Collins et al. (2011); Martin et al. (2011)	146 576 ⁺	24 90 ⁺	2 22 ⁺	28.2 28.0 ⁺	167 (9.0) 163 ⁺ (8.6 ⁺)	190 (4.5) 224 ⁺ (7.0 ⁺)
IPSL-CM5A-LR	Dufresne et al. (2013)	156 1000 ⁺	28 178 ⁺	8 98 ⁺	28.3 28.3 ⁺	157 (7.2) 151 ⁺ (6.0 ⁺)	162 (24.9) 162 ⁺ (19.1 ⁺)
IPSL-CM5A-MR	Dufresne et al. (2013)	156 300 ⁺	31 56 ⁺	19 34 ⁺	29.2 29.0 ⁺	158 (9.1) 157 ⁺ (8.6 ⁺)	174 (18.5) 162 ⁺ (20.7 ⁺)

Table 1 continued

Model	References	Years	El Niño events	DP events	DWPE isotherm	Mean longitude (σ)	
						DWPE ($^{\circ}\text{E}$)	Western peak ($^{\circ}\text{E}$)
IPSL-CM5B-LR	L'Institut Pierre-Simon Laplace Coupled Model, version 5B, coupled with NEMO low resolution	156 300 ⁺	29 63 ⁺	13 21 ⁺	28.7 28.7 ⁺	167 (12.3) 181 ⁺ (12.1 ⁺)	190 (18.7) 190 ⁺ (11.9 ⁺)
MIROC-ESM	Model for Interdisciplinary Research on Climate, Earth System Model	156 680 ⁺	17 58 ⁺	10 25 ⁺	27.7 27.9 ⁺	155 (6.3) 152 ⁺ (5.6 ⁺)	157 (13.0) 154 ⁺ (15.8 ⁺)
MIROC5	Model for Interdisciplinary Research on Climate, version 5	163 670 ⁺	23 93 ⁺	5 57 ⁺	28.5 28.2 ⁺	174 (18.8) 161 ⁺ (12.1 ⁺)	158 (15.1) 171 ⁺ (11.5 ⁺)
MPI-ESM-LR	Max Planck Institute Earth System Model, low resolution	156 1000 ⁺	27 161 ⁺	9 86 ⁺	28.4 28.6 ⁺	166 (15.2) 159 ⁺ (7.3 ⁺)	162 (12.4) 162 ⁺ (21.1 ⁺)
MPI-ESM-P	Max Planck Institute Earth System Model	156 1156 ⁺	23 196 ⁺	12 111 ⁺	28.9 28.7 ⁺	161 (8.9) 161 ⁺ (10.3 ⁺)	162 (25.0) 162 ⁺ (20.3 ⁺)
MRI-CGCM3	Meteorological Research Institute Coupled Atmosphere-Ocean General Circulation Model, version 3	156 500 ⁺	28 74 ⁺	11 39 ⁺	28.2 28.1 ⁺	170 (8.4) 164 ⁺ (7.2 ⁺)	194 (20.1) 181 ⁺ (9.9 ⁺)
NorESM1-M	Norwegian Earth System Model, version 1 (intermediate resolution)	156 501 ⁺	42 117 ⁺	9 57 ⁺	28.4 28.7 ⁺	166 (11.3) 169 ⁺ (12.8 ⁺)	180 (18.5) 187 ⁺ (8.4 ⁺)
NorESM1-ME	Norwegian Earth System Model, version 1 (intermediate resolution) with interactive carbon cycle	156 252 ⁺	30 53 ⁺	5 15 ⁺	28.4 28.2 ⁺	176 (12.0) 169 ⁺ (12.1 ⁺)	178 (13.1) 192 ⁺ (7.6 ⁺)

Unmarked data are sourced from *historical* simulations; Data marked with a ⁺ are derived from *PiControl* simulations

determined from the centers of warming that enclose SST' of a critical threshold—here, at least 75 % of the maximum SST' —for the 2 years surrounding the peak of each El Niño event. El Niño events may be “double peaked” when two separate, concurrently growing, centers of warming are identified in the evolution of the equatorial SST' that each exceed the critical threshold. The two peaks must be separated by cooler SST' . This definition allows us to distinguish between an El Niño event that evolves by propagation from east to west versus one in which the two peaks develop concurrently.

2.4 The mixed layer heat budget

The mixed layer heat budget equation used in this study is adapted from Vialard et al. (2001) and is given by

$$\partial_t T' = A'_x + A'_y + A'_z + Q' + DER', \quad (1)$$

where the symbol ∂_t represents a partial derivative with respect to time, the apostrophe ' denotes an anomalous quantity, and T' is the anomalous potential temperature integrated over the mixed layer. The term A'_x on the right-hand side represents the mixed layer averaged anomalous zonal advection defined as

$$A'_x = -\frac{1}{h} \int_{-h}^0 [\overline{u^* \partial_x T^{*'}} + u^{*'} \partial_x \overline{T^*} + \overline{u^{*'} \partial_x T^{*'}}] dz, \quad (2)$$

where u^* is the 4-dimensional (i.e., in time, latitude, longitude, and depth) zonal current, T^* is the 4-dimensional potential temperature, and the overline notation denotes a climatological quantity. The terms A'_y and A'_z in Eq. (1) represent anomalous meridional and vertical advection, respectively, and are constructed similarly to Eq. (2). The vertical velocity used to calculate A'_z is taken directly from the model output. Q' in Eq. (1) is the anomalous net surface heat flux, which can be calculated by summing the surface shortwave and longwave radiation, and latent and sensible heat fluxes, and subtracting the net shortwave radiation contribution that penetrates through the mixed layer (Q_{swout}). Q' is scaled by the mixed layer depth (MLD, h), the constant specific heat capacity of seawater ($c_p = 3989.24 \text{ J kg}^{-1} \text{ K}^{-1}$), and a constant density of seawater ($\rho_0 = 1035 \text{ kg m}^{-3}$). The term DER' in Eq. (1) represents anomalous residual processes, such as diffusion, turbulent heat fluxes, and entrainment into the mixed layer, that are not well resolved when the heat budget is calculated offline. The time-varying MLD over which the terms are averaged is denoted h , and is defined as the depth at which the density layer σ_t deviates from surface values by 0.125 kg m^{-3} (calculated offline). Derivatives are computed using centered differences. All heat budget calculations are performed on monthly mean output of u , v , w , T , and h .

An offline calculation of the heat budget equation may lead to some terms being over- or underestimated,

particularly nonlinear or eddy-related terms. For example, tropical instability waves (TIWs) that are important for the damping of SST' in the eastern equatorial Pacific on seasonal to interannual timescales require sub-monthly resolution to be adequately quantified (Vialard et al. 2001). The closure between $\partial_t T'$ calculated directly from the ACCESS models and the right-hand side of Eq. (1) will be further affected by uncertainties introduced through offline calculations. Finally, the residual term includes heat produced through mixing—a process that is not well-resolved in an offline parameterization. Nevertheless, offline calculation of heat budget terms has been used widely (Zhang et al. 2007; Huang et al. 2010, 2011; Choi et al. 2012; Graham et al. 2014) and is sufficient for our purposes of determining the dominant balance of terms giving rise to the CGCM double peaked El Niño. Equation (1) and its derivation are described in more detail in Vialard and Delecuse (1998).

In what follows, we refer to the depth-averaged (i.e., 3-dimensional, in time, latitude, and longitude, rather than the asterisked 4-dimensional) forms of the terms on the right-hand side of Eq. (2), and the corresponding terms for A'_y and A'_z . For example, in the case of A'_x , the three terms on the right-hand side of Eq. (2) simplify to $-\overline{u} \partial_x \overline{T'}$, $-u' \partial_x \overline{T'}$, and $-\overline{u'} \partial_x T'$.

3 Results

3.1 The double peaked El Niño event in CMIP5 CGCMs

The metric described in Sect. 2.3 is applied to *PiControl* and *historical* simulations of 36 CMIP5 models (Table 1). Double peaked El Niño events are common in all of the CGCMs during the period over which they are simulated (Fig. 4). Note that our selection of these 36 CMIP5 CGCMs is not dependent on them simulating a double peaked El Niño event. Several models (e.g., GFDL-ESM2G, ACCESS1-3, IPSL-CM5A-MR, and MPI-ESM-P) have a large number of double peaked El Niño events for both *PiControl* and *historical* conditions, while several others (e.g., FIO-ESM, GISS-E2-R-CC, GISS-E2-R, HadGEM2-CC) have relatively few double peaked El Niño events for both *PiControl* and *historical* conditions.

The evolution of SST' composites during double peaked El Niño events from *historical* simulations of selected CGCMs are compared in Fig. 3. Despite variations in magnitude and timing of El Niño onset between the CMIP5 CGCMs, in each model two warm peaks in SST' emerge during the first 6 months of the El Niño event. The warm peaks grow simultaneously, and separately, during the onset and development of El Niño.

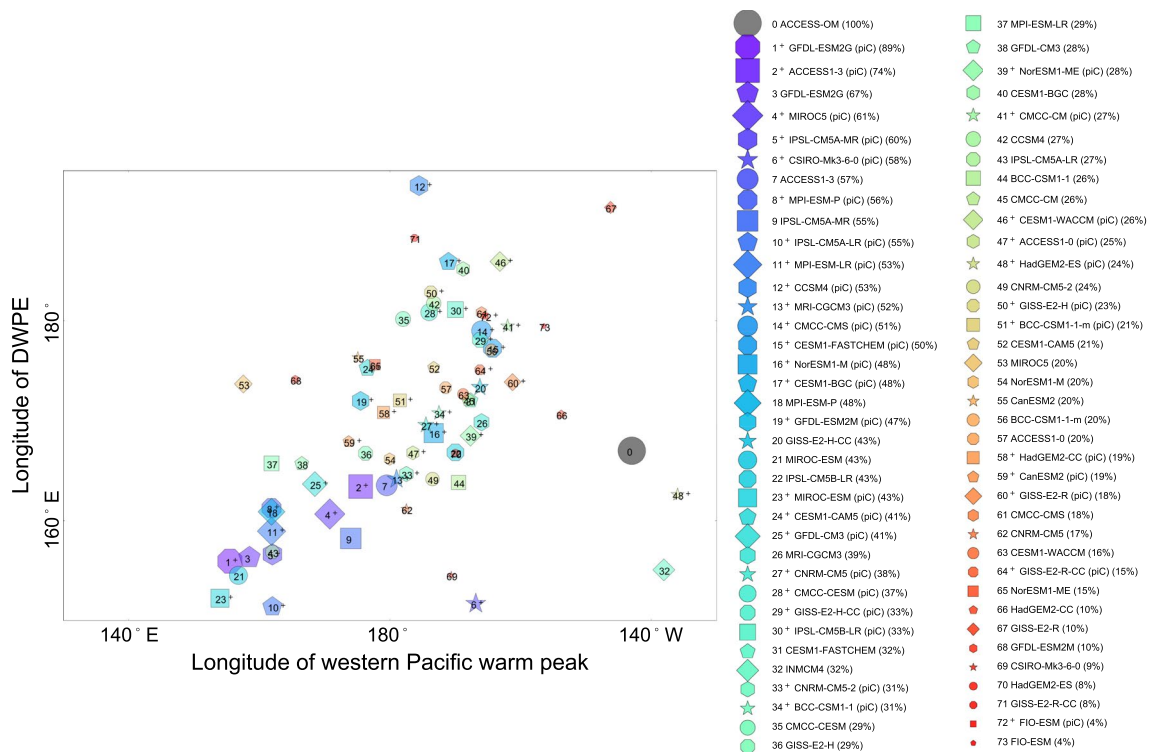


Fig. 5 Mean position of the western Pacific warm peak in a composite double peaked El Niño year versus the mean position of the dynamic warm pool edge in *PiControl* and *historical* simulations of 36 CMIP5 CGCMs. Markers representing each CGCM are sized by the fraction of double peaked events to the total number of El Niño events (see Table 1). The large grey circle represents the mean longitude of the western Pacific warm peak

Compared with *historical* simulations, *PiControl* simulations are systematically biased towards simulating more double peaked El Niño events (Fig. 4). 86 % of the 36 CGCMs simulate a greater proportion of double peaked El Niño events in pre-industrial conditions than in *historical* conditions. The mean fraction of double peaked El Niño events to all El Niño events in *PiControl* simulations is approximately 40.3 % compared with 26.2 % in *historical* simulations.

The location of the western Pacific warm peak during double peaked El Niño events varies from approximately 140°E to 140°W in the CMIP5 CGCMs. We investigate whether this is related to the magnitude of the cold tongue bias, which has been found to extend El Niño-related warming in CGCMs further westwards than observed (Taschetto et al. 2014). We use the mean location of the dynamic warm pool edge (DWPE)—the isotherm that best captures the maximum in the zonal salinity gradient—as a proxy for the magnitude of the cold tongue bias. This is because CGCMs with stronger cold tongue biases tend to simulate DWPEs further towards the western Pacific warm pool (Brown et al. 2013). The relationship between the cold tongue bias and the location of the western Pacific warm peak during

tude of a composite eastern Pacific El Niño event (x -axis) versus the mean longitude of the dynamic warm pool edge (y -axis) for a 60-year simulation (1948–2007) of the flux-forced ACCESS-OM model. This ACCESS-OM simulation does not have any El Niño events classified as double peaked using the definition in Sect. 2.3

double peaked El Niño events in the CMIP5 CGCMs is illustrated in Fig. 5. A clear pattern emerges: during double peaked El Niño events, models with stronger cold tongue biases also simulate western Pacific warm peaks located further towards the western Pacific warm pool. Furthermore, models that simulate more double peaked El Niño events tend to have DWPEs shifted further west than models with fewer double peaked El Niño events: 14 of the 20 models with the highest fraction of double peaked events simulate a DWPE west of the median ($\approx 170^\circ\text{E}$). This relationship corroborates our earlier result that the fraction of double peaked El Niño events is greater in *PiControl* simulations, where the cold tongue is strengthened relative to *historical* conditions due to a relative decrease in atmospheric CO_2 concentrations (Vecchi et al. 2006a; Collins et al. 2010; Vecchi and Wittenberg 2010; Watanabe et al. 2012). However, this relationship is not necessarily indicative of the full extent of the cold tongue bias in the model. That is, Fig. 5 only incorporates the double peaked events that meet the criterion outlined in Sect. 2.3; it does not take into account models that simulate other spatial patterns of El Niño that have not been evidenced in the observational record (e.g., both CSIRO-Mk3-6-0 and CNRM-CM5

simulate El Niño events evolving exclusively in the western Pacific warm pool) and might be a result of the cold tongue bias, or indeed other biases, in coupled models.

We found earlier that double peaked El Niño events are more prevalent in *PiControl* simulations than in *historical* simulations. Given the importance of the DWPE in generating double peaked El Niño events, we test whether the change in fraction of double peaked El Niño events from *PiControl* to *historical* simulations is related to mean state changes. Both the change in the mean longitude of the DWPE (dDWPE), and the fraction of double peaked El Niño events in *PiControl* simulations ($F(\text{piC})$; i.e., the *PiControl* mean state), are predictors for the change in the fraction of double peaked El Niño events from *PiControl* to *historical* (dF). We test the dependence of dF on each of these predictors using multiple linear regression. Considering all subsets of the two predictors, the model that yields the best fit ($R^2 = 0.40$) has the form

$$\text{dF} = a * \text{dDWPE} + b * F(\text{piC}) + c, \quad (3)$$

where a , b , and c are constant coefficients. The parameters from the regression analysis are highlighted in Table 2 and the resulting fitted data in Fig. 6. The *PiControl* mean state, $F(\text{piC})$, is found to have the greatest effect on the change in fraction of double peaked El Niño events simulated.

In what follows, we examine the evolution of heat budget dynamics during double peaked El Niño events in ACCESS-CM1.3. We note that it is possible that the CGCM double peaked El Niño event arises due to different mechanisms in different CMIP5 models; however, analysis of all CMIP5 models is beyond the scope of the current study.

3.2 The double peaked El Niño event in ACCESS-CM1.3

A total of 89 El Niño events are identified in the 505-year *PiControl* simulation of ACCESS-CM1.3. Of the CGCM El Niño events, 65 are classified as double peaked events and 10 as eastern Pacific events. In a further 12 of the remaining events two distinct peaks of warming are present, as in the double peaked El Niño event, but the SST' in

either the eastern or western peak does not meet the threshold to allow classification as a double peaked event. SST' for the ACCESS-CM1.3 double peaked and eastern Pacific El Niño events are composited. The significance of the heat budget trends from these composite events is investigated and discussed in Appendix 2 (Fig. 13).

3.2.1 Heat budget analysis

The heat budget terms from Eq. (1) are analyzed in the ACCESS-CM1.3 *PiControl* simulation to determine the mechanisms giving rise to the western Pacific warm peak of the double peaked El Niño event (Fig. 7). During the double peaked event, westerly wind anomalies generated near the DWPE (163°E) incite the growth of eastwards zonal current anomalies there. The strong zonal current anomalies occur at the maximum in the mean zonal temperature gradient (Picaud et al. 1996, 1997; Clarke et al. 2000), which is displaced further to the west than observed due to the cold tongue bias (Brown et al. 2013). This leads to the zonal advective feedback $-u'\partial_x \bar{T}$ achieving its maximum near the DWPE, and dominating the growth of the mixed layer temperature anomaly, T' , there. Warming induced by the zonal advective feedback then increases the positive anomalous mixed layer temperature gradient in the western Pacific, leading to growth of the mean zonal advection term $-\bar{u}\partial_x T'$ in the western Pacific. The climatological westward flow of the South Equatorial Current, which is up to 0.4 m s^{-1} stronger than observed in the western-central Pacific in the CGCM than in observations (Fig. 8), advects the western warm patch to the west.

We next investigate how the ACCESS-CM1.3 double peaked El Niño event differs from the eastern Pacific event (Fig. 7). The western extent of warming extends west of 160°E during both CGCM El Niño events; however, a western warm peak does not develop in the ACCESS-CM1.3 eastern Pacific event, partly due to an eastwards shift in the patterns of westerly wind stresses, which is consistent with previous studies (Rasmusson and Carpenter 1982; Kalnay et al. 1996; Wittenberg 2004). That is, during the ACCESS-CM1.3 eastern Pacific El Niño event the maximum in the westerly (i.e., anomalous) equatorial zonal wind stresses is shifted further to the east ($150^\circ\text{--}120^\circ\text{W}$) than in the ACCESS-CM1.3 double peaked El Niño event ($150^\circ\text{E--}160^\circ\text{W}$), and the westerly wind stresses in the region of the western Pacific warm peak ($150^\circ\text{E--}180^\circ$) are weaker by approximately $3.4 \times 10^{-3} \text{ N m}^{-2}$ on average during the first 24 months of the eastern Pacific event than the double peaked El Niño event. As a consequence, the zonal advective feedback is smaller by approximately $0.10^\circ\text{C month}^{-1}$ in the western Pacific during the eastern Pacific El Niño event than during the double peaked event, preventing the development of significant warming

Table 2 Fitted values and confidence intervals for the parameters in Eq. (3)

Parameter	Estimate	95 % confidence interval	
		Min	Max
a	−0.0049	−0.010	0
b	−0.44	−0.65	−0.23
c	0.043	−0.049	0.14

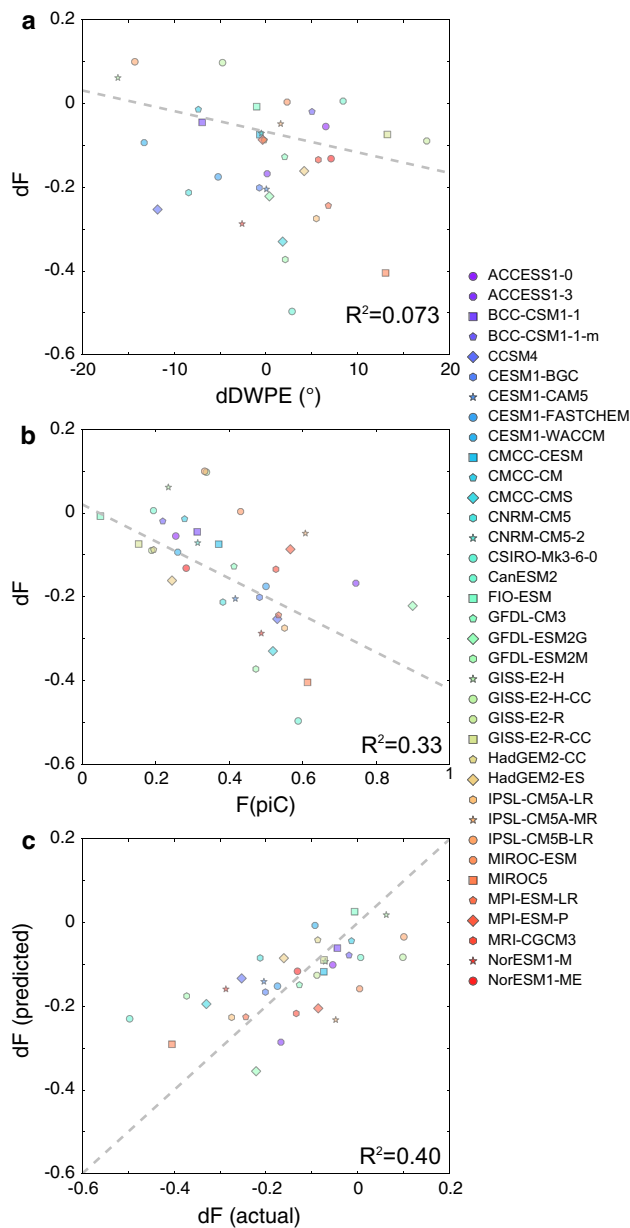


Fig. 6 The relationship between: **a** the change in the mean position of the dynamic warm pool edge (dDWPE) and the change in the fraction of double peaked El Niño events (dF) from the *historical* to the *PiControl* simulations; **b** the fraction of double peaked El Niño events in the *PiControl* simulations (F(piC)) and the change in the fraction of double peaked El Niño events from the *historical* to the *PiControl* simulations; and **c** the actual and predicted (i.e., from Eq. (3)) change in the fraction of double peaked El Niño events from the *historical* to the *PiControl* simulations. The grey dashed lines in **a** and **b** represent the line of best fit from ordinary least squares regression, and in **c** represents the 1:1 line between the actual and predicted values. The R^2 values from the multiple linear regression analysis are reported

in the western Pacific. This is consistent with observed El Niño events that develop mainly in the eastern Pacific, the growth of which is typically dominated by the thermocline

feedback (Jin 1997a, b; Yeh et al. 2014). The nonlinear meridional advection term $-v'\partial_y T'$ provides consistent warming throughout the central Pacific during eastern Pacific El Niño events. By contrast, during double peaked events this term is almost negligible in the central Pacific, but larger ($>0.15^\circ\text{C month}^{-1}$) in the western Pacific, where anomalous meridional temperature gradients are amplified.

The relative contributions of the heat budget terms to the central equatorial Pacific $\partial_t T'$ during the double peaked and eastern Pacific El Niño events are shown in Fig. 9. Here, the difference between each heat budget term during double peaked and eastern Pacific El Niño events (i.e., the bottom panels in Fig. 7) in the central equatorial Pacific (defined as the local minimum in SST' variance, 154°W) is subtracted from the difference in the western-central equatorial Pacific (the local maximum in SST' variance, 178°E). The key drivers of the western Pacific warm peak are the zonal advection terms $-u'\partial_x \bar{T}$ and $-u'\partial_x T'$, which are the result of relatively stronger anomalous zonal equatorial currents acting on the zonal temperature gradient at the edge of the western Pacific warm pool. The meridional advection terms $-v'\partial_y T'$ and $-\bar{v}\partial_x T'$ contribute to the growth of $\partial_t T'$ by meridional spreading of the equatorial SST' . The net surface anomalous heat flux Q' grows, rather than damps, the western Pacific warm peak, largely due to a positive bias in the shortwave heat flux attributed to unrealistic SST-cloud interactions in ACCESS-CM1.3 (Rashid and Hirst 2015). These unrealistic SST-cloud interactions are partly due to a climatological bias in the low cloudiness of ACCESS-CM1.3, associated with an overly strong cold tongue compared with observations, and also partly due to overly strong descending atmospheric motion. A similar result has been found in the GFDL-CM2.1 CGCM (Wittenberg et al. 2006). While residual eddy effects do contribute somewhat to generating the western Pacific warm peak, they are relatively weak compared to zonal advection, meridional advection, and thermodynamic damping contributions.

4 Discussion

Spurious double peaked El Niño events - with two warming peaks developing concurrently in the eastern and western Pacific - were found to be widespread in CMIP5 CGCMs. The location of the western Pacific warm peaks during double peaked El Niño events was correlated with the location of the dynamic warm pool edge (DWPE), a proxy for the magnitude of the cold tongue bias (Brown et al. 2013). The DWPE was as far west as 155°E in CMIP5 models. CGCMs with more westwards located DWPEs tended to simulate more double peaked El Niño events. The consistency in the response of the CMIP5 CGCMs in simulating the double peaked events serves to corroborate the cold

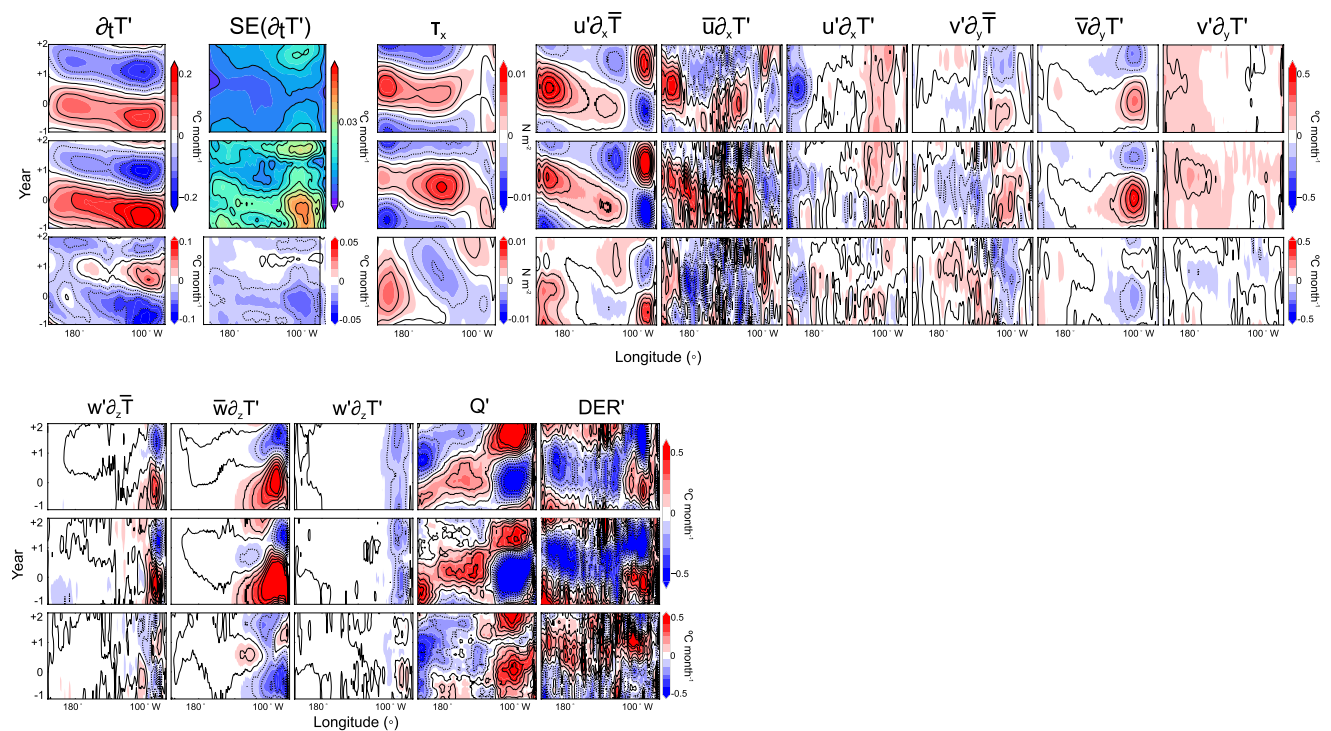


Fig. 7 Evolution of heat budget terms for the 36 months surrounding composite double peaked (*top panels*) El Niño events and eastern Pacific (*middle panels*) El Niño events from ACCESS-CM1.3. The *bottom panels* show the difference between the double peaked and eastern Pacific El Niño events. Data are averaged over 2°S – 2°N . Wind stress anomaly (τ_x) data are in units of N m^{-2} (contour interval 0.01 N m^{-2}), and the units of the remaining panels are $^{\circ}\text{C month}^{-1}$ (contour interval $0.01 ^{\circ}\text{C month}^{-1}$). The interval between 0 and +1 represents the first year of the El Niño composite event. Note the difference in the *color scale* between the tendency term and the remaining heat budget feedbacks. The terms represented in each column are,

from *left*: the mixed layer temperature tendency anomaly, the standard error (SE) of the mixed layer temperature tendency anomaly, the zonal wind stress anomaly, the zonal advective feedback ($-u'\partial_x T'$), the mean zonal advection term ($-\bar{u}\partial_x T'$), the anomalous zonal advection term, and the meridional heat budget terms. Terms represented in each column of the lower plot are, from *left*: the Ekman feedback ($-w'\partial_z T'$), the thermocline feedback ($-\bar{w}\partial_z T'$), the anomalous vertical advection term, the net surface heat flux anomaly (Q'), and the residual term DER' , namely, $\partial_t T' - (A'_x + A'_y + A'_z + Q')$.

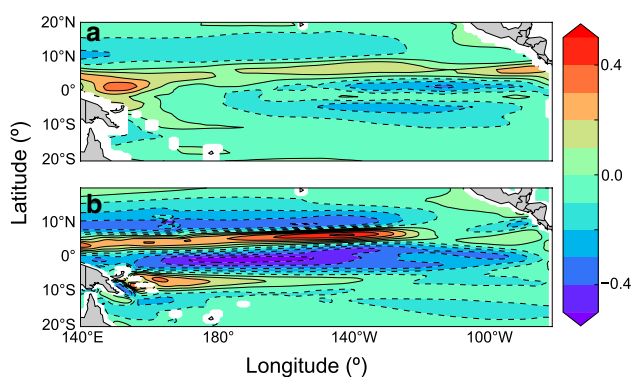


Fig. 8 Mean zonal currents over the period 1993–2005 (*shading*) derived from the **a** Ocean Surface Current Analyses Real-time (OSCAR; available at <http://www.oscar.noaa.gov>), and **b** the *historical* simulation of ACCESS-CM1.3. Data are in units of m s^{-1} and the contour interval is 0.1 m s^{-1}

tongue bias as playing an important role in generating the double peaked El Niño events, rather than this event representing a realistic “new flavor” of El Niño. Consequently, a reasonable supposition is that the ENSO behaviors present in *PiControl* and *historical* simulations of CGCMs fit within circle (iii) in Fig. 1; that is, they display some qualitatively similar features to those observed, but also simulate some unrealistic ones that are an artifact of climatological CGCM biases.

The mechanisms giving rise to the double peaked event were further investigated in ACCESS-CM1.3. During double peaked El Niño events in ACCESS-CM1.3, the westwards extension of the equatorial Pacific cold tongue region (eastern extent of the western Pacific dynamical warm pool edge) modified the location of peak warming and dynamical behavior in the western Pacific. In particular, the overly

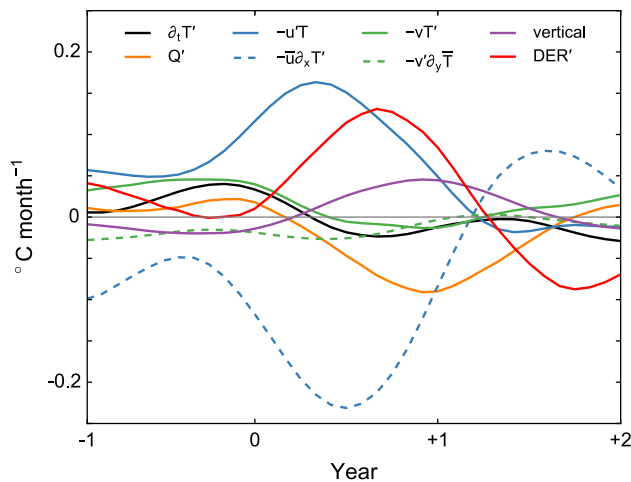


Fig. 9 Difference between ACCESS-CM1.3 heat budget terms ($^{\circ}\text{C month}^{-1}$) in the central equatorial Pacific during double peaked and eastern Pacific El Niño events. Terms are calculated as the difference in the central equatorial Pacific at 154°W subtracted from the difference in the western-central equatorial Pacific at 178°E . The heat budget terms are as in Fig. 7, but with $-u'T' = -u'\partial_x\bar{T} - u'\partial_xT'$, $-v'T' = -v'\partial_y\bar{T} - v'\partial_yT'$, and vertical $= -\bar{w}\partial_z\bar{T} - w'\partial_zT' - w'\partial_z\bar{T}$

intense and westward-extended cold tongue in CGCMs led to two biases that altered the El Niño feedbacks compared with eastern Pacific El Niño events: (1) an excessive climatological zonal temperature gradient ($\partial_x\bar{T}$) in the western equatorial Pacific that was displaced too far west of the strong climatological vertical temperature gradient in the eastern equatorial Pacific; and (2) atmospheric deep-convective cloudiness that was displaced too far west and off-equator. The westward-shifted $\partial_x\bar{T}$ led to a western displacement of the zonal advective feedback ($-u'\partial_x\bar{T}$) relative to vertical advective feedbacks ($-\bar{w}\partial_zT'$ and $-w'\partial_z\bar{T}$), generating a secondary western equatorial warm peak. In addition, the intense cold tongue displaced the atmospheric convective zones westward and poleward, leading to insufficient damping of this secondary western peak in SST' by cloud shading. These results highlight the importance of a CGCM's climatology to the dynamics and spatial structure of ENSO and motivate further attention to understanding and correcting mean state biases in CGCMs.

Here, we have focused on just one manifestation of CGCM El Niño diversity: the double peaked pattern of SST warming. Given the similarity in mechanisms giving rise to the western Pacific warm peak of the double peaked event and the central Pacific El Niño event (Yeh et al. 2014), it is possible that the double peaked El Niño event could be mistaken for a westwards-shifted central Pacific El Niño, particularly in composite El Niño events. Furthermore, differences between CGCMs can lead to behaviors that have not yet been observed (e.g., the El Niño event that evolves entirely in the western Pacific warm pool in

CSIRO-Mk3.6.0). It follows that studies of future ENSO events, such as changes in the frequency of El Niño spatial behaviors under global warming scenarios, should be cautiously interpreted in light of historical representations of El Niño diversity.

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Appendix 1: The mean state and biases in ACCESS-CM1.3

The mean SST from ACCESS-CM1.3 and SST bias, with respect to the Bureau of Meteorology Research Centre (BMRC) SST reanalyses (Smith 1995) over the period 1980–2004, are illustrated in Fig. 10. ACCESS-CM1.3 is up to 1°C cooler than the reanalysis data in the equatorial Pacific cold tongue region (180° – 100°E), and up to 2°C warmer east of 100°W along the coast of South America. ACCESS-CM1.3 displays a warm bias in the South Pacific, in the region of the South Pacific Convergence Zone, and in the tropical North Pacific (5°N , 160° – 110°W).

The standard deviation of tropical Pacific SST' is indicative of the spatial diversity in ENSO variability (Fig. 2). Variability in the eastern equatorial Pacific in ACCESS-CM1.3 is weaker than in the reanalysis data (the difference in standard deviation is up to 0.6°C at approximately 100°W), including $<0.3^{\circ}\text{C}$ from 160° – 140°W , and slightly

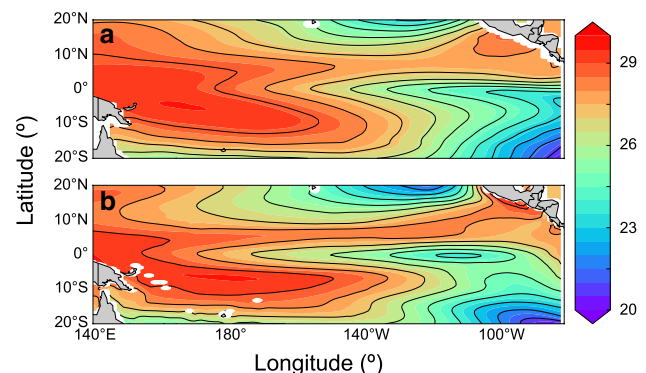


Fig. 10 Mean sea surface temperature over the period 1980–2004 (shading) in the **a** BMRC reanalyses, and **b** ACCESS-CM1.3. Data are in units of $^{\circ}\text{C}$ and the contour interval is 0.5°C

stronger ($>0.2^{\circ}\text{C}$) west of 180° longitude in a secondary western peak. Note that the standard deviation of SST' illustrated in Fig. 2 is qualitatively similar to the leading mode of an EOF analysis of ACCESS-CM1.3 SST' , which also displays the double peaked pattern of warming and represents 44 % of the SST' variability in ACCESS-CM1.3 (figure not shown).

The annual means of the equatorial surface heat fluxes for ACCESS-CM1.3 are compared with those from the Objectively Analyzed air-sea Fluxes (OAFlux; provided by the Woods Hole Oceanographic Institute (WHOI) OAFlux project, available at <http://oafux.whoi.edu>), the TropFlux reanalyses (Kumar et al. 2012), and the Coordinated Ocean-ice Reference Experiments version 2 (CORE-II, which are used to force ACCESS-OM; Griffies et al. 2012) in Fig. 11. The annual mean equatorial longwave radiation and sensible heat flux simulated by ACCESS-CM1.3 are within the range of uncertainty estimated from OAFlux, TropFlux,

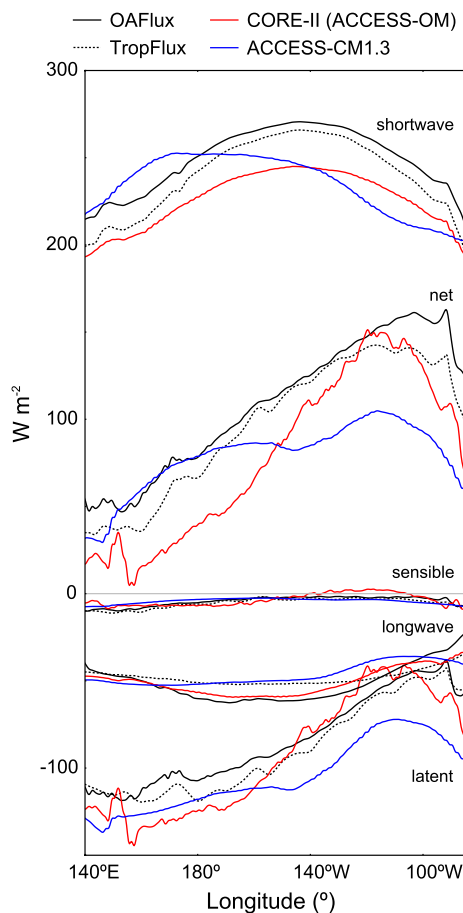


Fig. 11 Annual mean of equatorial surface heat flux variables—namely, shortwave, sensible, latent, longwave, and net heat fluxes—from ACCESS-CM1.3 (blue), the CORE-II reanalyses from ACCESS-OM (red), the OAFlux reanalyses (black solid), and the TropFlux reanalyses (black dashed). Data are averaged between 2°S and 2°N and are in units of W m^{-2}

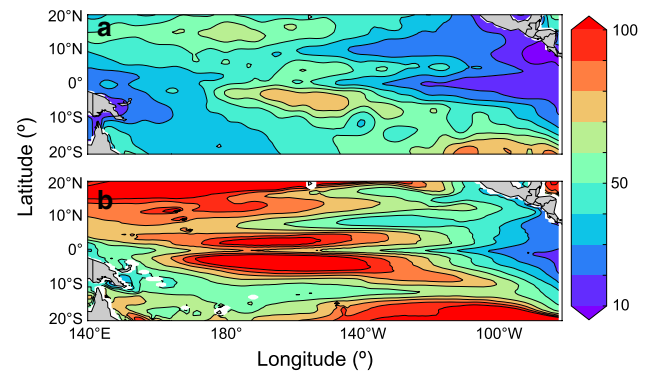


Fig. 12 Mean mixed layer depth (MLD) over the period 1980–2005 (shading) in the **a** UK Met Office (UKMO) reanalyses, and **b** ACCESS-CM1.3. The MLD is defined as the depth at which the density layer σ_t deviates from surface values by 0.125 kg m^{-3} . Contours show the bias in mean mixed layer depth with respect to the UKMO data. Data are in units of m and the contour interval is 10 m

and CORE-II. Latent heat fluxes in ACCESS-CM1.3 are up to 46 W m^{-2} less than those of the reanalyses, particularly in the eastern equatorial Pacific. Equatorial shortwave radiation values simulated by ACCESS-CM1.3 in boreal winter are up to 38 W m^{-2} different from TropFlux.

The mean state of the tropical Pacific MLD in ACCESS-CM1.3 and bias with respect to the UK Met Office (UKMO) subsurface ocean temperature and salinity data (Ingleby and Huddleston 2007) over the period 1980–2005 are compared in Fig. 12. The ACCESS-CM1.3 MLDs are up to 50 m deeper than the UKMO MLDs in bands stretching between 170°E and 150°W north and south of the equator.

Appendix 2: Significance of the double peaked El Niño event in ACCESS-CM1.3

Here, we investigate whether the composited double peaked El Niño events are significantly different from the composited eastern Pacific El Niño events. First, the double peaked and eastern Pacific El Niño events from the *PiControl* simulation of ACCESS-CM1.3 are randomly separated into two groups, groups *a* and *b*, and composited. We name these composites μ_x of sample size n_x , where $x \in \{DP1.3a, DP1.3b, EP1.3a, EP1.3b\}$. We also consider the double peaked El Niño events from the *PiControl* simulation of ACCESS-CM1.0 and separate them into two composites— $\mu_{DP1.0a}$ and $\mu_{DP1.0b}$ —with sample sizes $n_{DP1.0a}$ and $n_{DP1.0b}$, respectively.

The variable for testing the significance of the difference between composites is the Student's *t*-distribution:

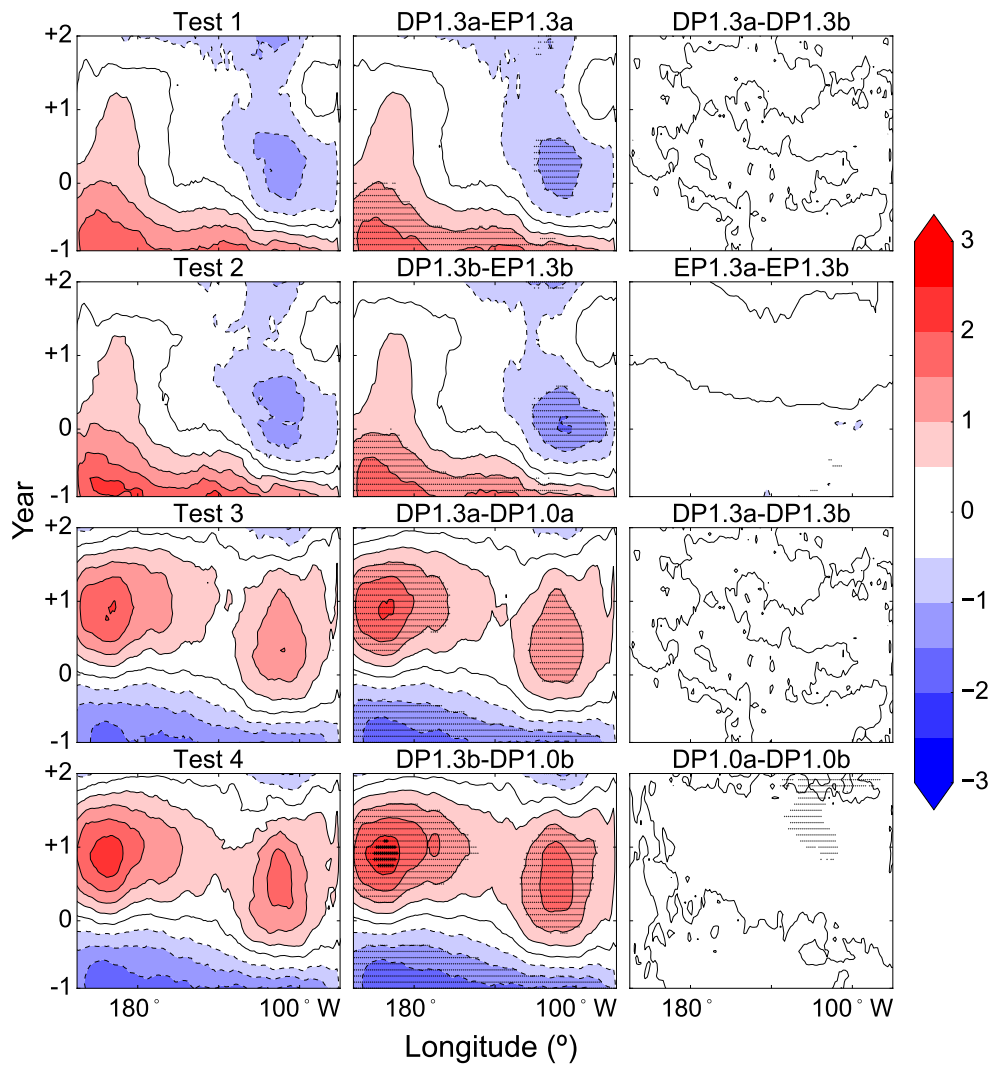


Fig. 13 Simple significance testing of SST' composites from randomly selected double peaked and eastern Pacific El Niño events in ACCESS-CM1.3 ($DP1.3a$, $DP1.3b$, $EP1.3a$, and $EP1.3b$, respectively) and double peaked El Niño events in ACCESS-CM1.0 ($DP1.0a$, $DP1.0b$, respectively) for the 3 years surrounding El Niño events. Data displayed are median t probability density function values calculated from 100 random samples of the test groups a and b .

$$t = \frac{\widehat{\mu}_x - \widehat{\mu}_y}{S \sqrt{\frac{1}{n_x} + \frac{1}{n_y}}}, \text{ and} \quad (4)$$

$$S^2 = \frac{(n_x - 1)\widehat{\sigma}_x^2 + (n_y - 1)\widehat{\sigma}_y^2}{n_x + n_y - 2}, \quad (5)$$

where $n_x + n_y - 2$ is the number of independent observations for the parameter t , and x and y represent the composited El Niño events being tested. The significance value (p value) from each test case is calculated using a two-sided Student's t -test.

The first column in each row is calculated by subtracting the third column from the second. Differences greater than one standard deviation from the mean are indicated with stippling ('.'), and differences significant at the 95 % confidence interval with crosses ('+'). In each case, significance is calculated using a two-sided Student's t test. The contour interval is 0.1

We define a simple test to establish the significance of the El Niño composite events: namely, the double peaked and eastern Pacific El Niño events are significantly different if the following conditions are satisfied during the evolution of the El Niño event (i.e., the first 24 months of the composite):

- Test 1** the differences between the $DP1.3a$ and $EP1.3a$ composites are greater than the differences between the $DP1.3a$ and $DP1.3b$ composites;
- Test 2** the differences between the $DP1.3b$ and $EP1.3b$ composites are greater than the differences between the $EP1.3a$ and $EP1.3b$ composites;

- Test 3* the differences between DP1.3a events from ACCESS-CM1.3 and DP1.0a events from ACCESS-CM1.0 are greater than the differences between the DP1.3a and DP1.3b events from ACCESS-CM1.3; and
- Test 4* the differences between DP1.3b events from ACCESS-CM1.3 and DP1.0b events from ACCESS-CM1.0 are greater than the differences between DP1.0a and DP1.0b events from ACCESS-CM1.0. The random sampling is repeated 100 times and median values for the differences between the composites, t , and p across the samples are calculated. The results for tests 1–4 are illustrated in Fig. 13.

For test 1, the median difference between DP1.3a and EP1.3a is approximately ± 2 times greater than the difference between DP1.3a and DP1.3b, which is in the range $[-0.37, 0.19]^\circ\text{C}$ for the 100 samples generated. The differences in DP1.3a and EP1.3a are greater than one standard deviation across the western-central equatorial Pacific during the 12 months prior to the peak of the El Niño event. The greatest differences in the eastern equatorial Pacific occur during the 2 months prior to and 8 months following the peak of the El Niño event. Differences between DP1.3a and DP1.3b across the 100 samples are not statistically significant. A similar result is found for test 2. Even in the *PiControl* simulations, the sample size of eastern Pacific events in ACCESS-CM1.3 is relatively small – 10 in total – such that the difference between EP1.3a and EP1.3b is likely to be biased by individual events.

The results of tests 3 and 4 illustrate that double peaked events from the ACCESS-CM1.3 model are more similar to each other than to events from ACCESS-CM1.0. Again, the median difference between double peaked events within each model simulation is small (within the range $[-0.22, 0.40]^\circ\text{C}$ for the ACCESS-CM1.0 simulation), while the median differences in double peaked events between the two models are close to $\pm 2^\circ\text{C}$ during the development of the El Niño event throughout the equatorial Pacific and in the western and eastern Pacific during the decay periods of the El Niño event (the differences are greater than one standard deviation from the mean in each case). These results provide evidence that the composite double peaked and eastern Pacific El Niño events from ACCESS-CM1.3 are sufficiently different to ensure significance in the trends analysis.

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