Cloud-radiative effects on implied oceanic energy transports as simulated by atmospheric general circulation models


Abstract. This paper summarizes the ocean surface net energy flux simulated by fifteen atmospheric general circulation models constrained by realistically-varying sea surface temperatures and sea ice as part of the Atmospheric Model Intercomparison Project. In general, the simulated energy fluxes are within the very large observational uncertainties. However, the annual mean oceanic meridional heat transport that would be required to balance the simulated surface fluxes is shown to be critically sensitive to the radiative effects of clouds, to the extent that even the sign of the Southern Hemisphere ocean heat transport can be affected by the errors in simulated cloud-radiation interactions. It is suggested that improved treatment of cloud radiative effects should help in the development of coupled atmosphere-ocean general circulation models.

Introduction

The focus of this study is on the surface energy fluxes over the oceans simulated by current atmospheric general circulation models (AGCMs), and the implied oceanic meridional energy transports. We summarize the implied partitioning of meridional energy transport between the ocean and the atmosphere and compare them with available observations. The implied oceanic meridional energy transport varies dramatically from model to model, and we show that these differences are largely due to cloud-radiative effects. This result has important implications for coupled atmosphere-ocean general circulation models.

Uncoupled AGCM simulations are performed by prescribing sea-surface temperatures (SSTs) and sea ice distributions. In a similar way, uncoupled ocean general circulation models (OGCMs) are integrated with prescribed surface wind stresses and relaxation of surface temperatures and salinity toward prescribed climatological values. Simulations produced by these models are quite realistic, but they are strongly constrained by the prescribed boundary conditions. Climate change studies must account for interactions between the ocean and the atmosphere, and so efforts are underway to couple AGCMs and OGCMs. When the boundary constraints are removed, in the coupled models, the simulated climate typically "drifts" toward an unrealistic state. To prevent this, most coupled models use "flux corrections," including ad hoc adjustments to the surface energy flux distribution (Manabe and Stouffer, 1993). We need to understand why such corrections are needed, so that ultimately they may be minimized or altogether eliminated. The results of this study shed light on this issue.

The need for a systematic and comprehensive intercomparison of AGCMs has been recognized by the World Climate Research Programme for some time. The Atmospheric Model Intercomparison Project (AMIP) (Gates, 1992) is the most recent such endeavor. AMIP simulations are 10 years long, and participating models use boundary conditions based on satellite-derived estimates of monthly mean SSTs and sea ice distributions for 1979-1988. Here we report on findings from AMIP subproject No. 5, using results from 15 AMIP AGCMs. All model results are year averages.

We must emphasize that when AGCMs are run with specified SSTs and sea ice, as in the AMIP runs, the surface radiation fluxes over the oceans are largely immaterial to the simulated atmospheric circulation. In particular, the simulated atmospheric circulation can be very realistic even if the surface radiation fluxes are not realistic. It is only when the AGCM is used as the atmospheric component of a coupled atmosphere-ocean model that the surface radiation fluxes become critical for the simulated climate.

Background

Annual-mean meridional energy transport can be defined for the atmosphere, $T_A$, the ocean, $T_O$, and the combined ocean-atmosphere system, $T_{A+O}$. Defining $\phi$ as latitude and $a$ as the radius of the earth, the annual average energy equations for the atmosphere, the ocean and the combined system are (see Fig. 1):

$$\frac{1}{\alpha \cos \phi} \frac{\partial}{\partial \phi} F_A \cos \phi = R_{up} - R_{svc} - LH - SH \quad (1.a)$$

$$= R_{up} - N_{svc},$$

$$\frac{1}{\alpha \cos \phi} \frac{\partial}{\partial \phi} F_O \cos \phi = + R_{svc} + LH + SH = N_{svc}. \quad (1.b)$$

$$\frac{1}{\alpha \cos \phi} \frac{\partial}{\partial \phi} F_{A+O} \cos \phi = R_{up}, \quad (1.c)$$

respectively, where the $F's$ represent northward energy fluxes. The vertical fluxes (downward being positive) $R_{up}, R_{svc}, LH$ and $SH$ represent the zonal mean (average along a circle of constant latitude) "top of the atmosphere" net radiative flux, and the ocean surface net radiative, latent and sensible heat fluxes respectively. The net ocean surface energy flux is $N_{ocean} = R_{ocean} + LH + SH$. A similar equation could be written for the land, but, because there is essentially no horizontal energy transport through the land, that equation is not needed here and the zonal averages of the surface fluxes in Eqs. 1.a-1.c are computed using ocean values only.

The horizontal transport of energy across a latitude circle is $T = 2\pi a \cos \phi F$, so that the equations above take the general form

$$\frac{1}{2\mu a^2} \frac{\partial}{\partial \phi} \left( T_{A+O}, T_A, T_O \right) =$$

$$\left( R_{up}, R_{up} - N_{svc}, N_{svc} \right) \cos \phi. \quad (2)$$

The northward transport may then be inferred from
\[ (T_{A+O}, T_A, T_O)(\phi) = \]
\[ 2n\alpha^2 \int_{-\pi/2}^{\pi/2} \left( R_{\text{top}} - R_{\text{top}} - N_{\text{ocean}} N_{\text{ocean}} \right) \cos \phi \, d\psi \]  
\[ \text{(3)} \]

for the total system and also for the atmosphere and the ocean separately, provided that the appropriate fluxes are known. The transports may also be obtained directly from observations of the atmosphere and the oceans. For example, Eq. 4 illustrates how atmospheric transports can be evaluated using measurements of temperature, geopotential, moisture and winds (Oort, 1983):
\[ T_A(\phi) = 2n\alpha \cos \phi \int_{\text{surface}} \left[ (c_p \rho + gz + Lq) \right] dz \]  
\[ \text{(4)} \]

Here \( c_p \) is the specific heat capacity of dry air, \( t \) the air temperature, \( g \) the acceleration due to the Earth's gravity, \( z \) the height, \( L \) the latent heat of vaporization, \( q \) the specific humidity and \( v \) the meridional wind. The brackets represent a zonal average, and the overbar an average in time. A similar equation can be written for ocean energy transport. Direct oceanic estimates are more difficult because of the lack of data on the three-dimensional distributions of oceanic temperatures and currents, but have been made at a few latitudes (Bryden et al., 1991).

The AMIP integrations are performed using atmospheric general circulation models with prescribed SSTs and sea ice distributions. The AGCMs simulate surface and top-of-the-atmosphere fluxes given these boundary conditions and the simulated working of the atmosphere, but the SSTs and sea ice do not respond to the local surface fluxes since there is no explicit oceanic adjustment. Although no explicit oceanic meridional energy transport is computed by the models, we use the oceanic energy transports implied by the simulated ocean surface fluxes (Eq. 3). Alternatively, calculating \( T_A \) from the appropriate top-of-the-atmosphere and surface fluxes allows one to calculate \( T_O \) by subtraction \( (T_O = T_{A+O} - T_A) \). Note that the implied oceanic transports obtained by this method will be in error if the long term net energy flux over land is not exactly zero, or if there is energy accumulating in the atmosphere or ocean. In our calculations, any non-zero annual mean of the globally averaged boundary fluxes (the right hand side of Eqs. 1a-1c) is removed uniformly over the globe. The imbalances in the AMIP simulations range between 10 and 10 W m\(^{-2}\). Tests have demonstrated that the assumed geographical distribution of the globally averaged energy imbalance does not alter the conclusions presented in this paper.

Results

Fig. 2a shows \( T_{A+O} \) as inferred from Eq. 3 using 4 years of net top-of-the-atmosphere radiation as observed in the Earth Radiation Budget Experiment (ERBE) (Barkstrom et al., 1990), and as simulated by the models. There is a broad range of simulated transports in Fig. 2a (especially in the Southern Hemisphere), which are generally less than observed. \( T_{A+O} \) is known with more accuracy than \( T_A \) or \( T_O \) because it is based on more reliable observations.

Fig. 2b shows observationally derived annual mean estimates of \( T_A \) by Oort (1983), Trenberth and Solomon (1994), and

Savijarvi (1988) obtained from meteorological observations, as well as the \( T_A \) results produced by the models. Trenberth used operational weather prediction analyses produced by the European Centre for Medium Range Weather Forecasts, while Savijarvi and Oort both used the same gridded radiosonde data. Savijarvi modified the radiosonde wind data to satisfy a dynamical constraint based on vorticity balance, whereas Oort did not. As shown in Fig. 2b, the results of Trenberth and Savijarvi agree very well in the Northern Hemisphere mid-latitudes where observations are best, even though they are based on independent data sources and analysis methods. They also agree reasonably well with the model simulations. The \( T_A \) estimate by Oort, on the other hand, is considerably smaller than the other observational estimates and the model results.
Figure 3a. Annual mean global ocean northward meridional energy transport: observational estimates.

Figure 3b. Annual mean global northward oceanic meridional energy transport: implied by simulations.

Figure 3c. Zonal and annual average net top of the atmosphere cloud radiative forcing: observations and simulations.

Figure 3d. Zonal and annual average net cloud radiative forcing: as simulated and observed. Annual average "hybrid" global ocean northward meridional energy transport: $T_{A+0}$ (observed) - $T_{O}$ simulated.

Estimates of $T_{O}$, based on hydrographic transects (Bryden et al., 1991, and Aagaard and Greisman, 1975) are also shown in Fig. 3a. The direct estimates are believed to be fairly accurate, especially the estimate for 65°N. Despite the general qualitative agreement seen in Fig. 3a, Southern Hemisphere data deficiencies are so severe that we cannot even be sure of sign of $T_{O}$.

Fig. 3b shows the implied $T_{O}$ obtained from $N_{con}$ fluxes of the AMIP simulations. For the Southern Hemisphere the AGCMs are in general qualitative agreement with the observations shown in Fig. 3a except for those of Carissimo et al. Although there is a large spread of results among the models, many have a maximum northward (poleward) Northern Hemisphere transport of $T_{O}$ - 2 x10^{15} W. The latitude at which this maximum occurs varies considerably from model to model. For the Southern Hemisphere, the range in the models' implied $T_{O}$ is much larger; the implied $T_{O}$ is northward (equatorward) in many cases, and southward (poleward) in only a few. The strong downward $N_{con}$ at 50°S in some of the models (see Fig. 2c), tends to force northward ocean transport between 50°S and the Equator.

It is important to identify the causes of these large discrepancies between the observed and simulated surface fluxes and implied ocean transports in the Southern Hemisphere. AGCMs are known to disagree considerably in their simulations of the effects of clouds on the Earth's radiation budget (Cess et al., 1990), and hence the effects of simulated cloud-radiation interactions on the implied meridional energy transports are immediately suspect. Unfortunately, we do not have global data on the effects on clouds on the surface energy budget, either as simulated by the models or from observations. We do, however, have observations of the effects of clouds on $R_{up}$ and thus $T_{A+0}$.

The cloud radiative forcing (CRF) is defined (Ramanathan et
The difference between the net radiation at the top of the atmosphere with the given distribution of cloudiness, and \( R_{\text{cl}} \), which we define to be the net radiation at the top of the atmosphere which would have been observed if no clouds were present but all else (e.g. temperature and water vapor) remained the same. Thus \( CRF = R_{\text{cl}} - R_{\text{d}} \). Clouds influence atmospheric heating primarily by trapping longwave energy within and beneath the cloud layer, and by reflecting sunlight back to space. Fig. 3c shows the zonally and annually averaged \( CRF \) from both the simulations and the ERBE data. There are large differences among the \( CRF \) simulated by the various models, and for the most part also between the observed and simulated cloud radiative forcing. In comparison with the ERBE data, many of the models underestimate the magnitude of the \( CRF \) at mid latitudes but overestimate it in the tropics. Poleward of 60° in both hemispheres the \( CRF \) observations are not reliable. The underestimate of the \( CRF \) in the Southern Hemisphere mid-latitudes is especially important for the implied \( T_o \) because the Southern Hemisphere oceans are much more extensive than those of Northern Hemisphere. A close comparison of these results and those in Fig. 2c suggests that there is a strong correlation between the magnitudes of the \( CRF \) and \( N_{\text{cloud}} \) and that differences in \( T_{o} \) and \( T_o \) are linked and largely due to cloud effects in the models. The discrepancies between the simulated and observed \( CRF \) suggest that the simulated implied northward \( T_o \) in the Southern Hemisphere is erroneous and is due to inadequate simulations of the \( CRF \).

We may connect the implied oceanic transport to the cloud radiative forcing using Eqs. 1-3. In particular, it follows from Eq. 3 that:

\[
T_{o} = 2\pi a^2 \int_{-\frac{3}{2}}^{\frac{3}{2}} \left[ R_{\text{cl}} + \left( R_{\text{up}} - R_{\text{cl}} \right) \right] \sin \phi d\phi \tag{5a}
\]

\[
= 2\pi a^2 \int_{-\frac{3}{2}}^{\frac{3}{2}} \left[ R_{\text{cl}} + CRF \right] \cos \phi d\phi \tag{5b}
\]

\[
= T_o + T_{CRF} \tag{5c}
\]

where \( T_o \) represents the \( T_o \) inferred from a "cloud-free" atmosphere, and \( T_{CRF} \) is that due to the radiative effects of clouds. We have computed a "hybrid" \( T_o \) defined by:

\[
\tilde{T}_o = T_{o,0} - T_o \tag{6a}
\]

\[
= (T_{o,0} - T_{o}) + (T_{o}^{ERBE} - T_{o,0}) \tag{6b}
\]

\[
= T_o + \delta T_{CRF} + \delta T_{CRF} \tag{6c}
\]

Here is the \( T_{o}^{ERBE} \) inferred from the ERBE data. The \( \delta T_{CRF} \) and \( \delta T_{CRF} \) represent the difference between the observed and simulated \( T_{o,0} \), resulting from the effects of clouds and a clear-sky atmosphere respectively. Comparison of the simulated cloudy and clear-sky fluxes has demonstrated that we can neglect the clear-sky effects on the implied \( T_o \) and thus:

\[
\tilde{T}_o \approx T_o + \delta T_{CRF} \tag{7}
\]

The resulting hybrid \( T_o \) which is based partly on observations and partly on simulations, is shown in Fig. 3d. The contrast with Fig. 3b is remarkable; in Fig. 3d, all of the hybrid results show poleward \( T_o \) in both hemispheres as a consequence of the cloud forcing "corrections."

Conclusions

The AGCMs shown here were run with fixed SSTs. The ocean energy transports inferred from these runs are not necessarily the same as the ocean energy transports that would be produced in coupled ocean-atmosphere simulations with the same AGCMs. Nevertheless, our results show that prescribed realistic SST distributions lead current AGCMs to produce surface energy budgets (Fig. 2c) that imply ocean energy transports that vary widely from model to model (Fig. 3b), especially in the Southern Hemisphere. These implied ocean energy transports cannot all be right, although they can all be wrong. Observations of the surface energy budget and/or the ocean circulation are not adequate to say which of the AGCM-implied ocean energy transports, if any, is correct. It is difficult to believe however, that the ocean energy transports could be equatorward at all latitudes in the Southern Hemisphere.

This paper presents quantitative evidence that the model-to-model variations in the implied ocean energy transports are largely due to model-to-model differences in the simulated cloud radiative forcing, which is in turn comparatively well observed quantity except at high latitudes. Our results thus indicate that as future AGCMs produce more realistic cloud radiative forcing, the simulated surface energy budget should improve. Coupled model simulations without flux corrections should also improve, although other factors are important such as the simulation of the surface fresh water flux and wind stress.

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