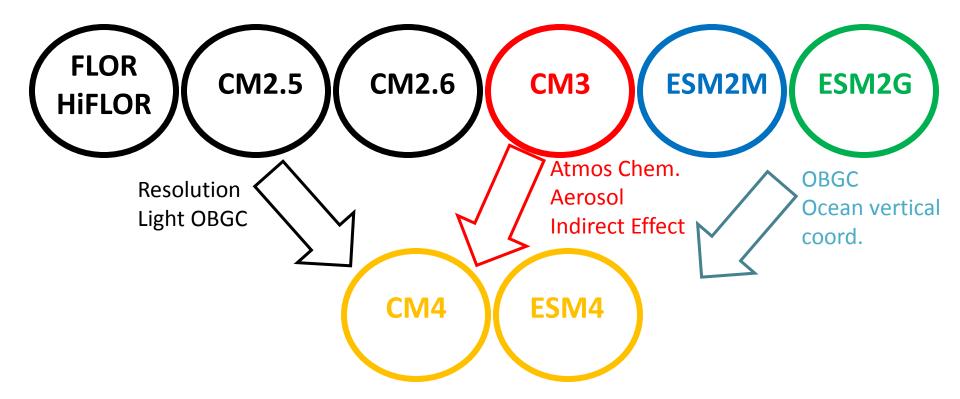
GFDL's next generation climate models: CM4 and ESM4

Presented by Michael Winton



Geophysical Fluid Dynamics Laboratory Fall Science Symposium November 2, 2017

Rationalizing GFDL's CMIP5 generation models



5-10 year Strategic Science Plan (2011) goal:

high resolution Earth System Model combining strengths of GFDL's multiple AR5 modeling streams



GFDL's CMIP6 generation models: CM4 and ESM4

	CM4 (frozen, starting DECK)	ESM4(in final development)
Atmosphere: AM4	100 km, 33 levels	100 km, 49 levels
Atmos. Chem	for aerosol (21 tracers)	aerosol+ozone (103 tracers)
Ocean: MOM6	1/4°, 75 levels	1/2°, 75 levels
Ocean BGC	BLINGv2 (6 tracers)	COBALTv2 (30 tracers)
Land	LM4.0	LM4.1 - PPA
Sea Ice	SIS2	SIS2

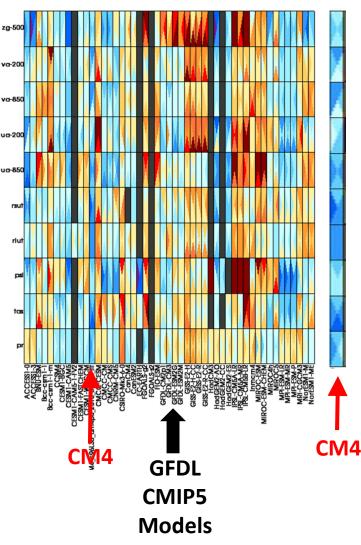
Note: All CM4 results shown are *preliminary* (based on potential vegetation historical, 1850- and 2010-forced experiments). We haven't yet run the official CMIP6 experiments with CM4.



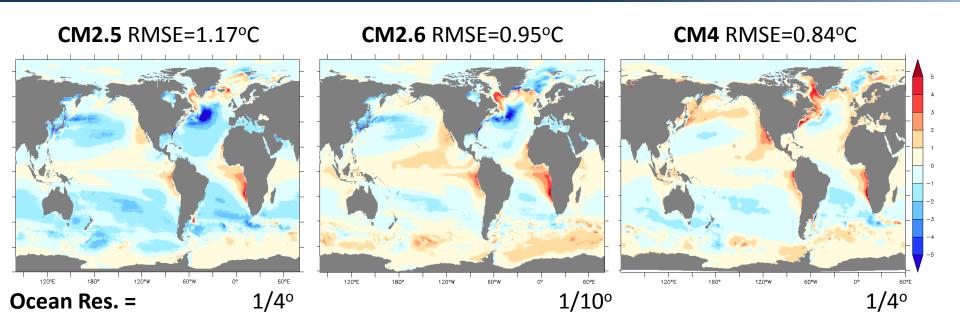
CM4 Surface Climate

- CM4's climatology is a distinct improvement over previous GFDL models
- CM4 temp., precip., OLR and reflected SW are the best in this CMIP5 ensemble
- Wind fields are good but not the best





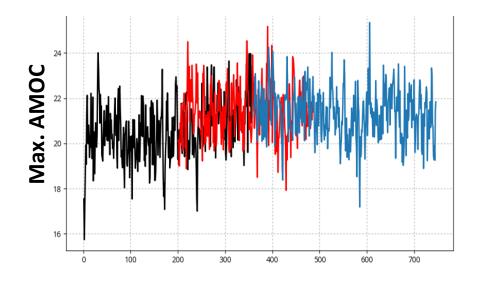
CM4 SST errors



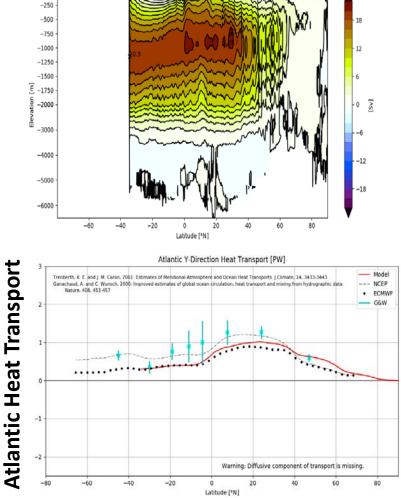
- CM4 SST error is smaller than CM2.6 (GFDL's previous best simulation)
- We expect CM4's SSTs can be improved further with higher ocean resolution (as seen refining CM2.5 to CM2.6) or with an eddy parameterization



Atlantic Meridonal Overturning



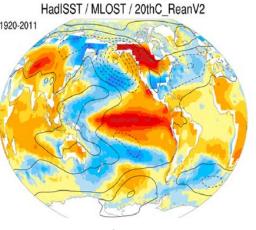
- Strong, stable AMOC
- Deep flow is too shallow and warm
- Heat transport less than observed



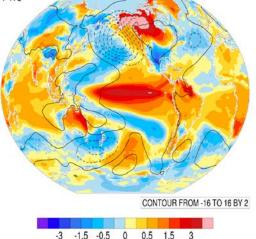
Atlantic MOC [Sv]

Variability: Improved ENSO

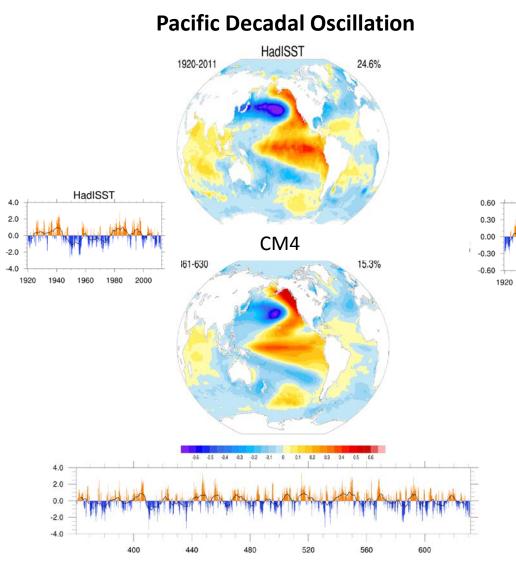
NINO3 SST spectra • ENSO magnitude is more 1920-2011 (a) NOAA ERSST.v4 obs (1957-2002) realistic than previous (b) CM4B 1850 Control_am4p0c96L33_OM4p25_30d15_tlt_bling8 **GFDL** models which din_p2_p7 0.0694 0.088 0.25 tended to be too large din_p7_1p4 0.795 0.5 0.5 din_1p4_9 0.54 CM4 iod (years) 0.554 period (YR) Delworth et al 2012 1-175 int25_1p4_9 2.32 2.38 ERSST.v3 (1880-2007 4 CM2.1 1990 (1-300 CM2.5 1990 (1-28 int50_1p4_9 3.23 3.38 0.0 0.5 1.0 1.5 2.0 2.5 °C²/octave 8 **FNSO** teleconnection int75_1p4_9 4.18 pattern is well simulated 4.39 16 0.00 0.50 1.00 1.50 2.00 2.50 (degC)²/octave RMSD(a,b) = 0.156corr(a,b) = 0.99



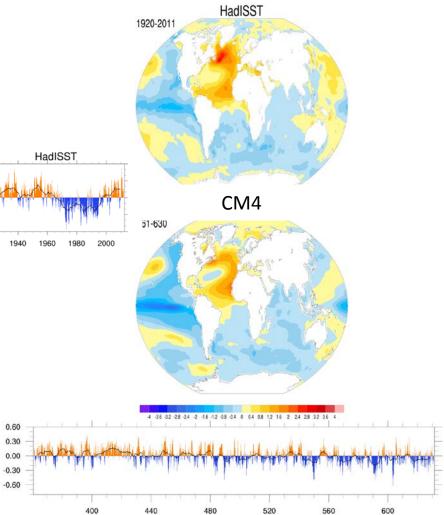
nino3.4 TS, TAS, PSL Spatial Composite (DJF⁺¹)



Variability: PDO and AMO patterns are well-simulated



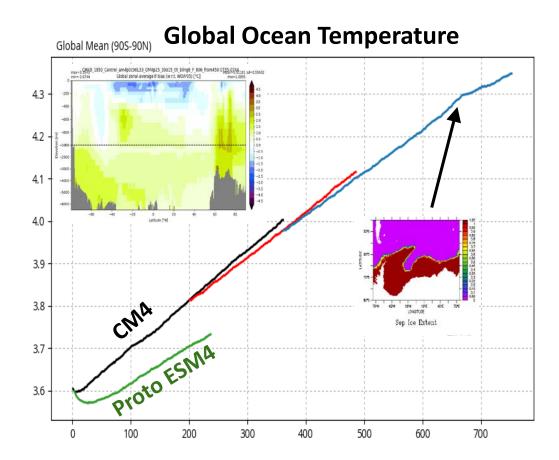
Atlantic Multi-decadal Oscillation



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Global Ocean Temperature Drift

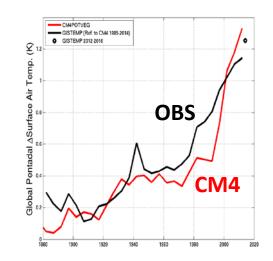
- Heat uptake is less than CM2.5 (also using 1/4° ocean)
- Heat uptake is less than the difference in heat uptake between CM2.6 and CM2.5 (eddy-permitting res. effect)
- Warming of deep water points to inadequacy of deep water formation representation (in both hemispheres)

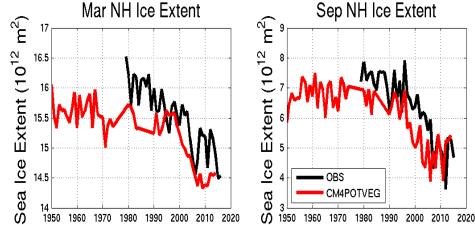




Historical Simulation: Global Temperature and NH Sea Ice Extent

- Historical warming roughly consistent with observed with possible exception of post-Pinatubo period.
- Good simulation of NH extent and its satellite era trend.
- SH sea ice low biased in summer, high biased in winter; recent observed increase is not simulated (not shown).



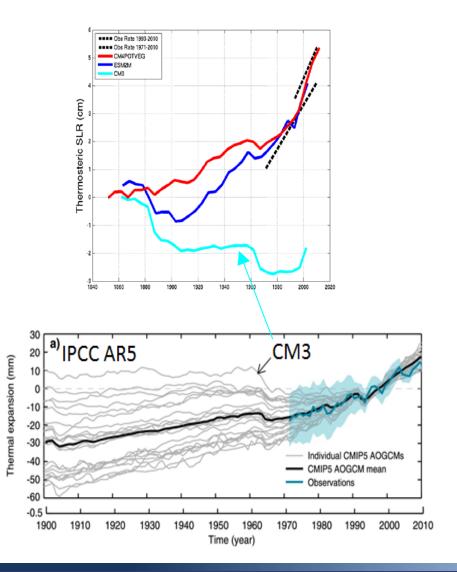


Thermosteric Sea Level Rise

CM3 thermosteric sea level rise problems:

- Excessive response to volcanoes (common to all CMIP5 models) due to lack of volcanic forcing in control experiment
- Lack of rise due to excessive aerosol forcing

CM4 has reduced aerosol forcing and improved simulation of ocean warming/thermosteric SLR





Summary

- CM4/ESM4 combine strengths of GFDL's CMIP5 generation of models into two, related models based on the same code with differing emphases on resolution and complexity.
- Expected CM4 strengths:
 - Surface climatology
 - ENSO variability; ENSO, AMO and PDO teleconnection patterns.
 - Reasonable historical climate change simulation
- Expected CM4 weaknesses:
 - \odot NADW too shallow and warm as in previous models
 - AABW formation only appears after 600 years of spin up
 - Ocean warm drift

