

# Patterns of temperature change and tropical precipitation/ cyclones

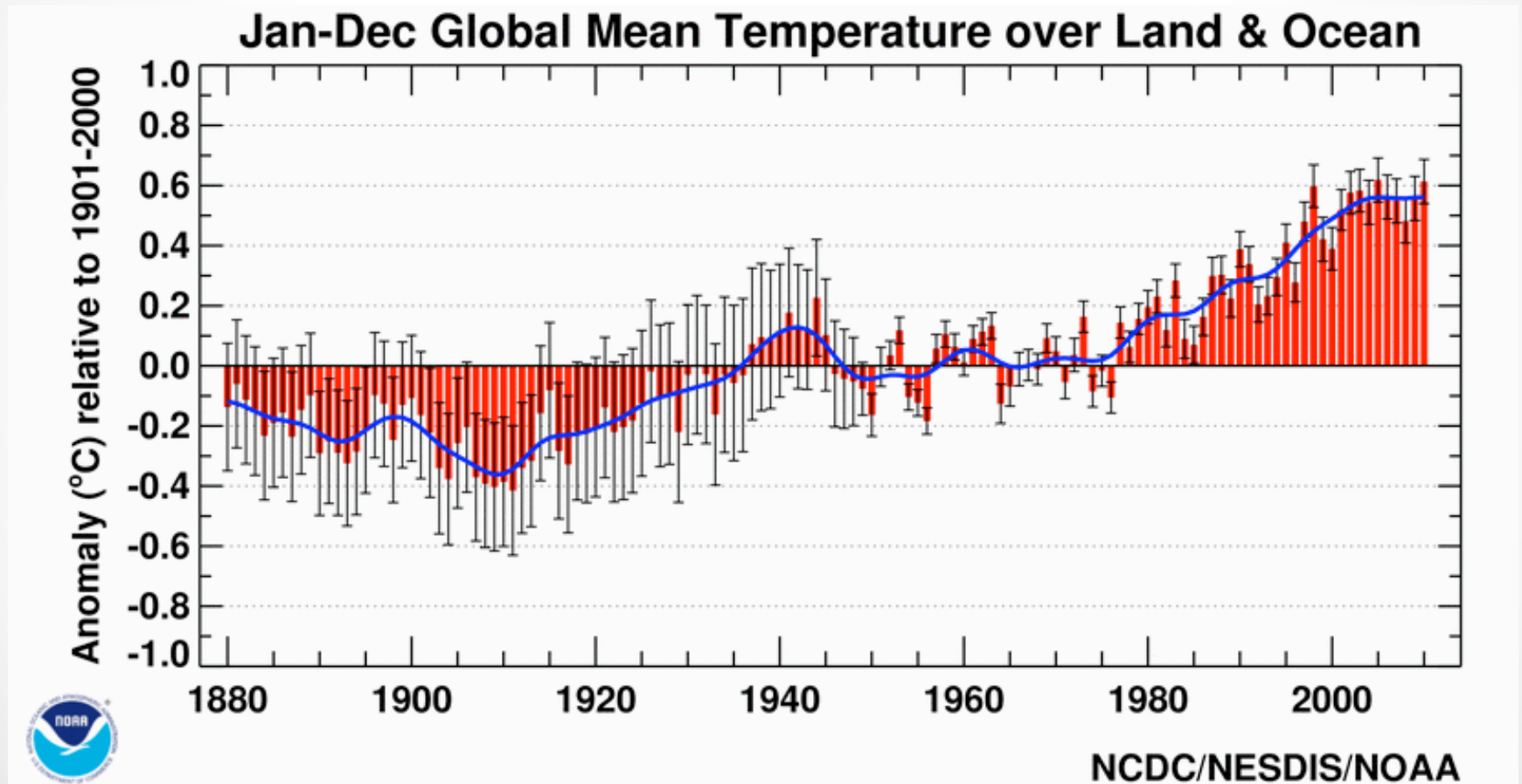
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NOAA/GFDL  
Princeton, NJ

# Summary

- Substantial progress has been made describing, understanding and attributing multi-decadal changes in global mean surface temperature.
- However, for regional tropical precipitation and tropical cyclones the “patterns” of temperature change may be more relevant:
  - Patterns in space: SST change *relative* to tropical-mean key to tropical precip and cyclone response
  - Patterns in height/depth (stability and stratification changes)
  - Patterns in time (multi-year to decadal changes)

# Global temperatures have shown clear long-term increase

(seen in other datasets, CRU, GISS, etc)

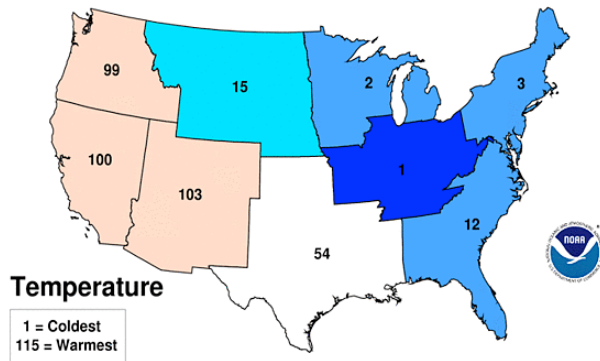


GCMs have exhibited skill in recovering these century-scale changes.

# On monthly and regional scales, changes have been more nuanced

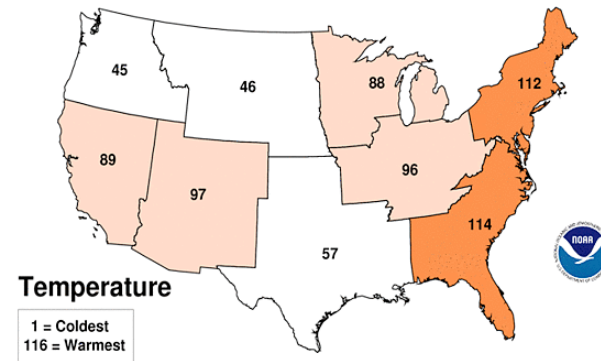
## July 2009 Regional Ranks

National Climatic Data Center/NESDIS/NOAA



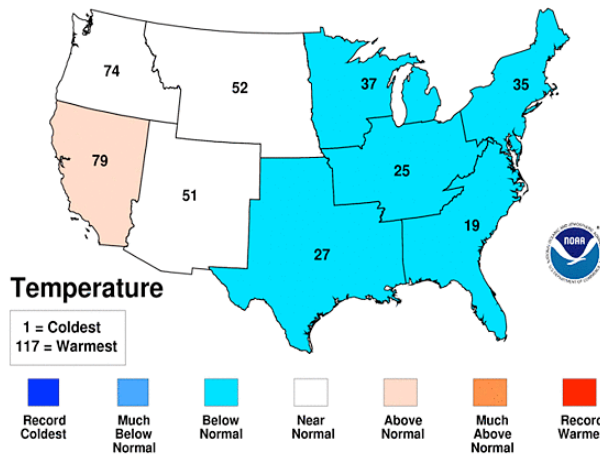
## July 2010 Regional Ranks

National Climatic Data Center/NESDIS/NOAA



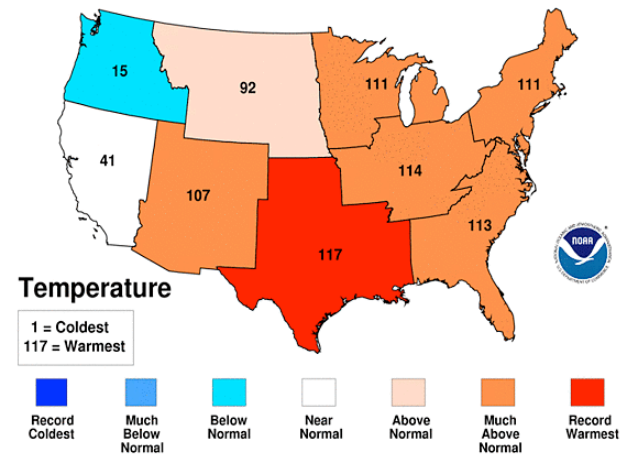
## January 2011 Regional Ranks

National Climatic Data Center/NESDIS/NOAA



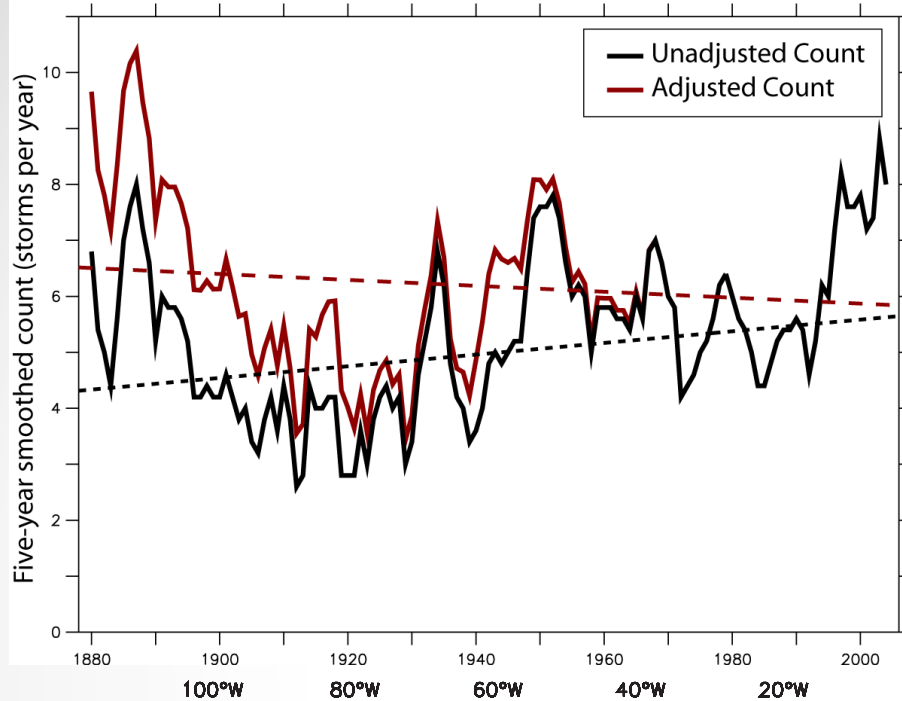
## July 2011 Regional Ranks

National Climatic Data Center/NESDIS/NOAA

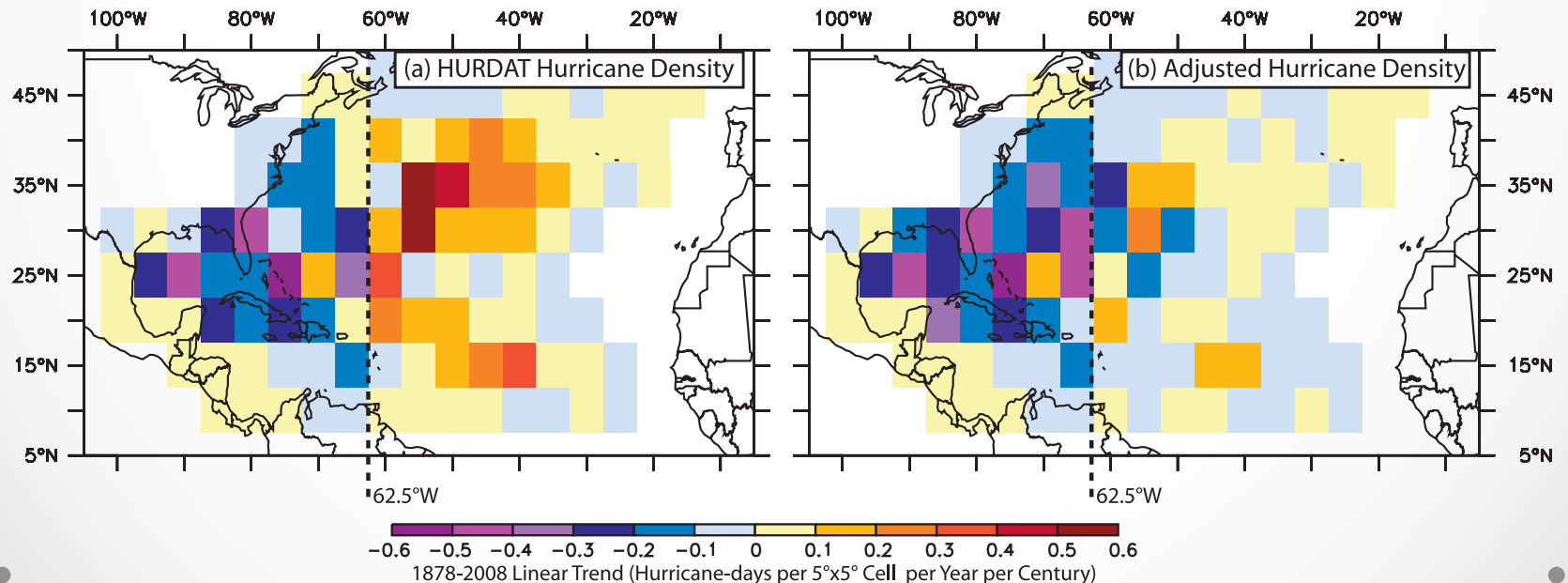




Count of Atlantic Hurricanes (Cat. 1-5)

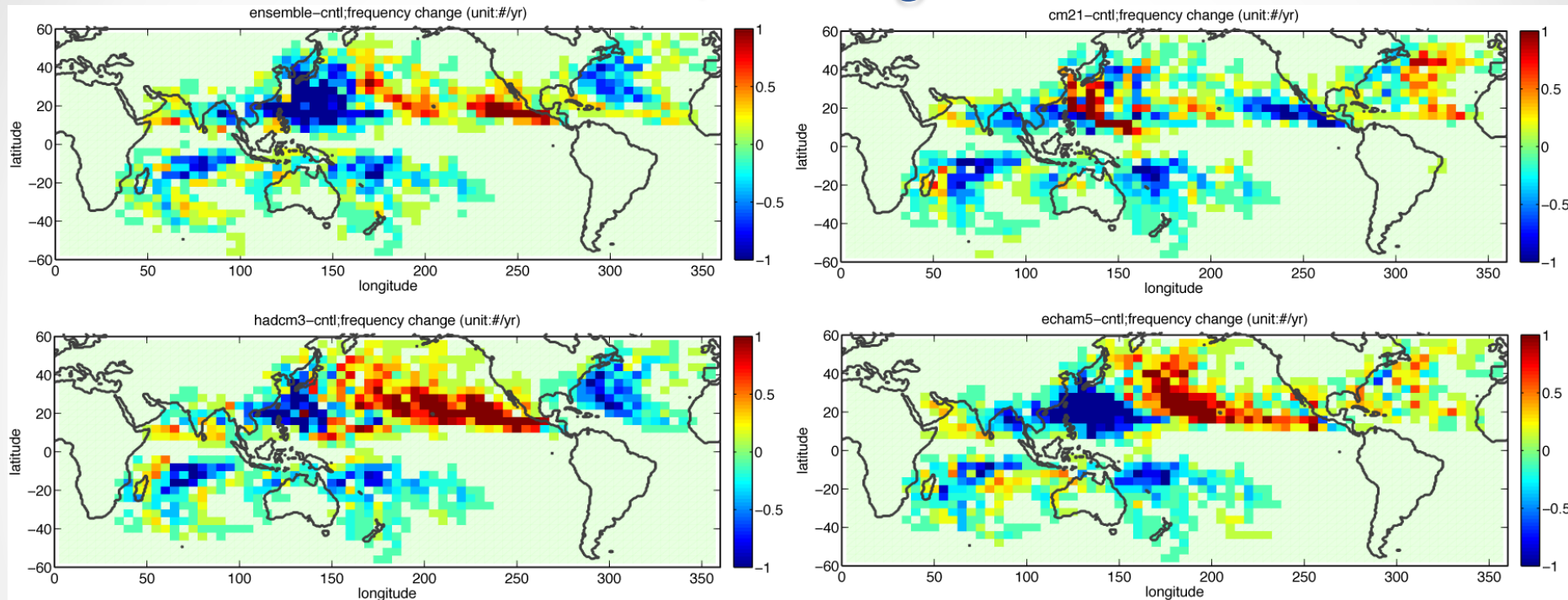


Observed hurricane changes: uncertain, not monotonic & spatially heterogeneous



*Vecchi and Knutson (2011, J. Climate)*

# 21<sup>st</sup> Cy response of TC frequency in single 50km global atmospheric model forced by four projections: uncertain, heterogeneous.



**Red/yellow = increase**  
**Blue/green = decrease**

*Adapted from Zhao et al. (2009, J. Climate)*

Regional increase/decrease much larger than global-mean.

Pattern depends on details of ocean temperature change.

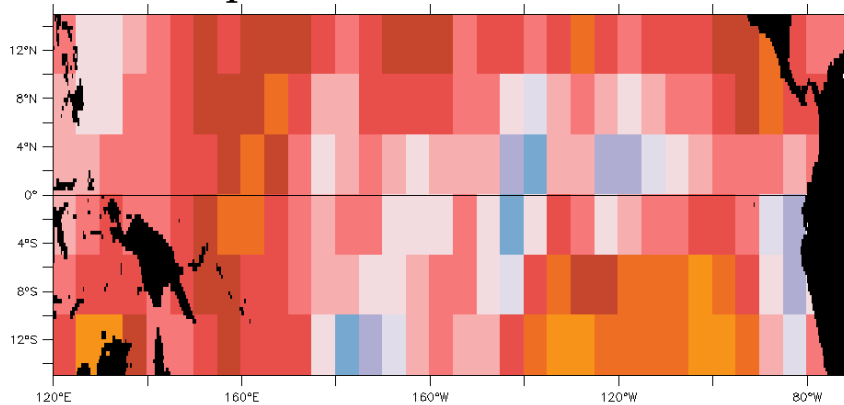
Sensitivity of response seen in many studies

*e.g., Emanuel et al 2008, Knutson et al 2008, etc*

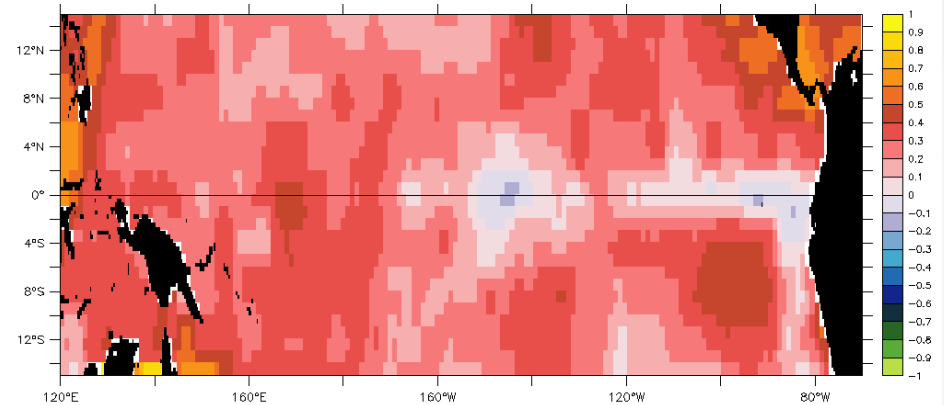
# Long-term SST trends in Tropical Pacific Uncertain

Linear trends (1880-2005) in four SST estimates.

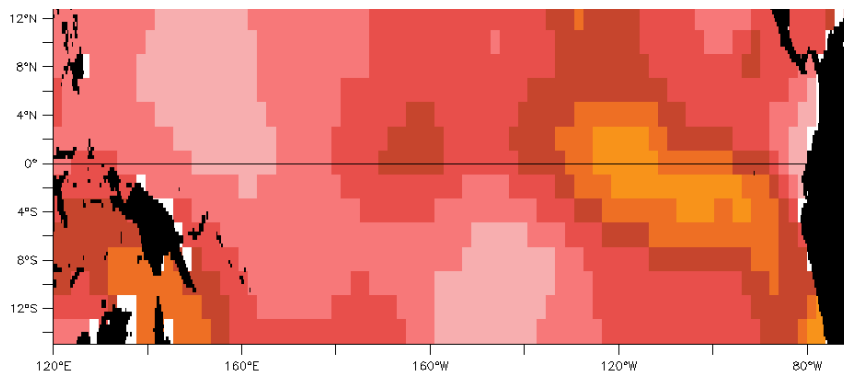
LDEO-Kaplan



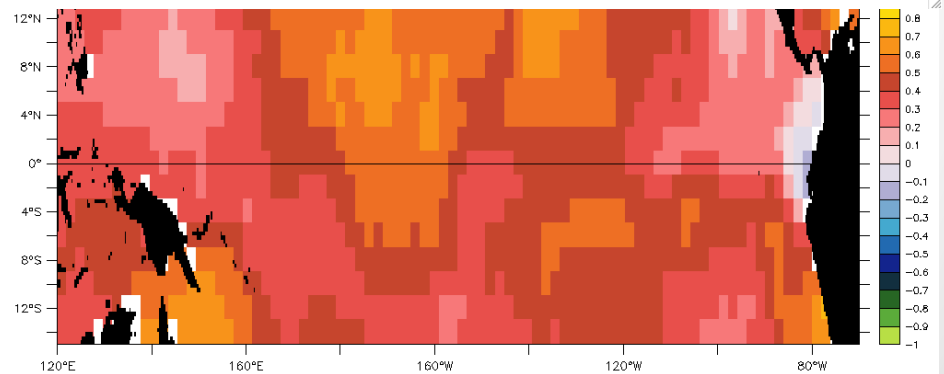
HadISST v.1



NOAA-ERSST v.2



NOAA-ERSST v.3



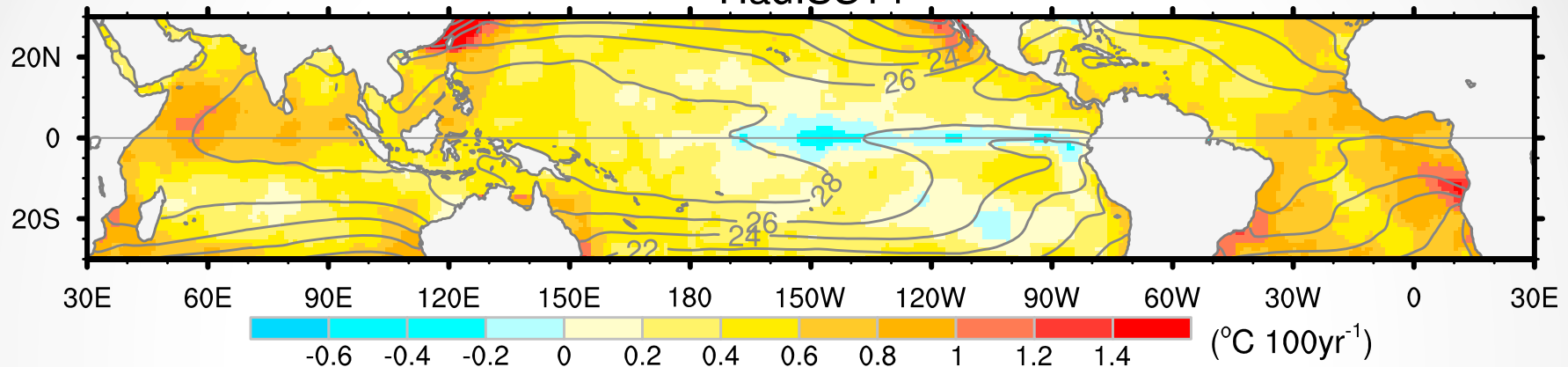
Overall warming seen in all. Structure dependent on reconstruction.

Adapted from Vecchi, Clement and Soden (2008, EOS)

# Data infilling leads to differences in east Pacific SST trends

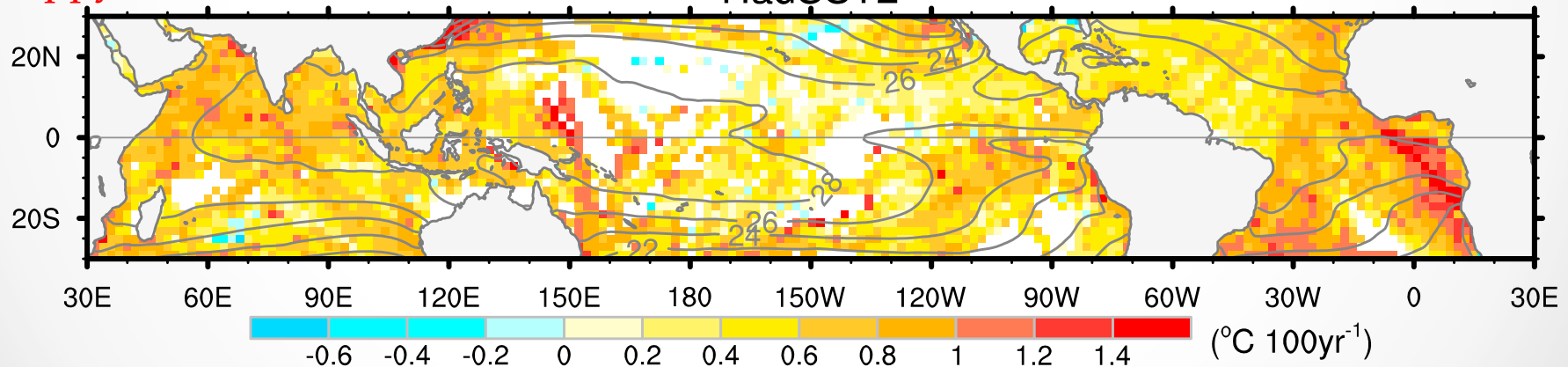
Infilled SST Product

HadISST1



"Gappy" SST Product

HadSST2



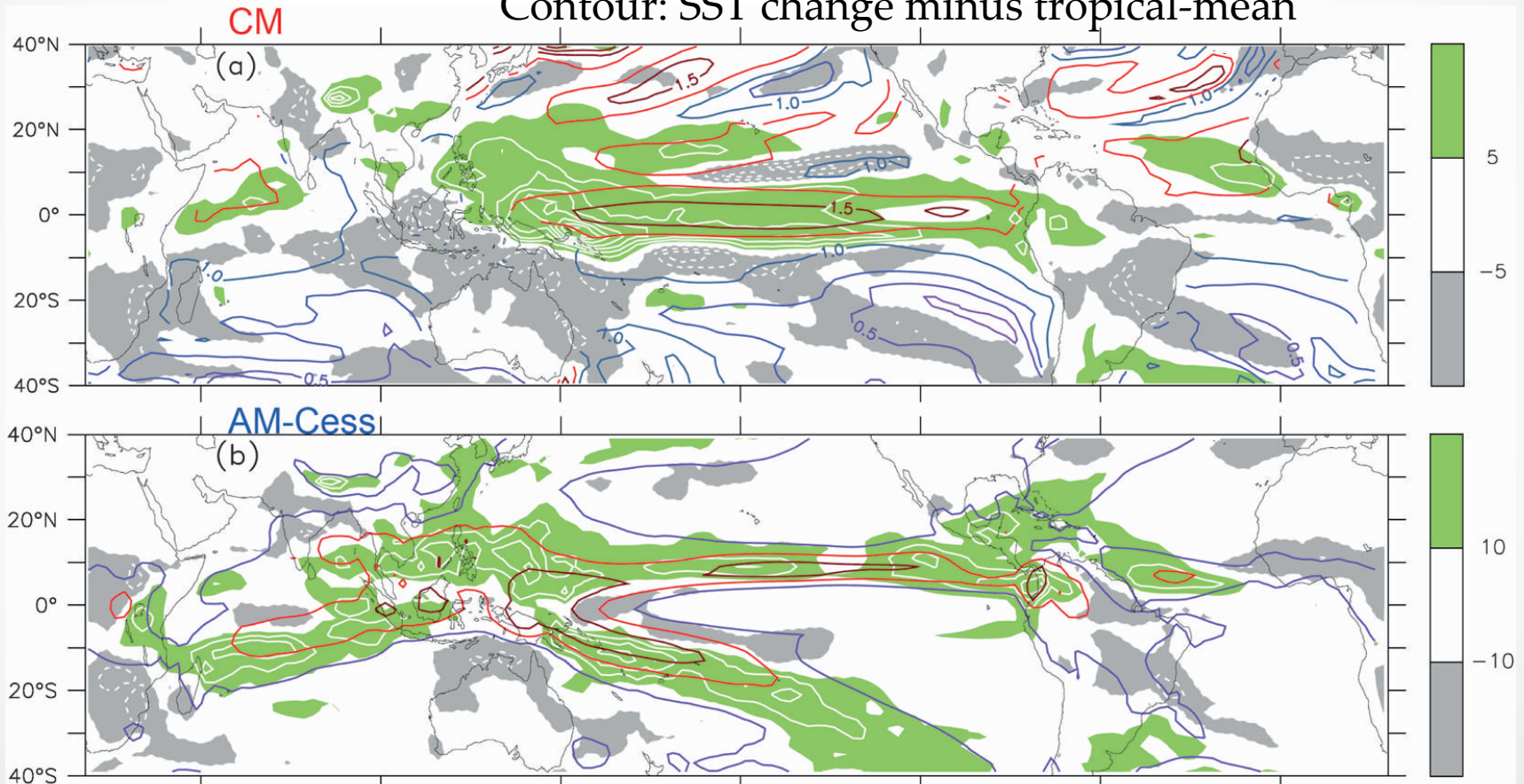
Deser et al. (2010, GRL)



# GFDL GCM precipitation response to 2xCO<sub>2</sub> and 2°C Uniform Warming

Shade: Precip response to 2xCO<sub>2</sub>

Contour: SST change minus tropical-mean



Shade: Precip response to Uniform 2°C

Contour: Control precipitation

Free tropospheric temperature changes relatively uniform (“Weak Temperature Gradient” Hypothesis, Sobel et al 2001); changes in stability follow surface temperature changes relative to tropical-mean

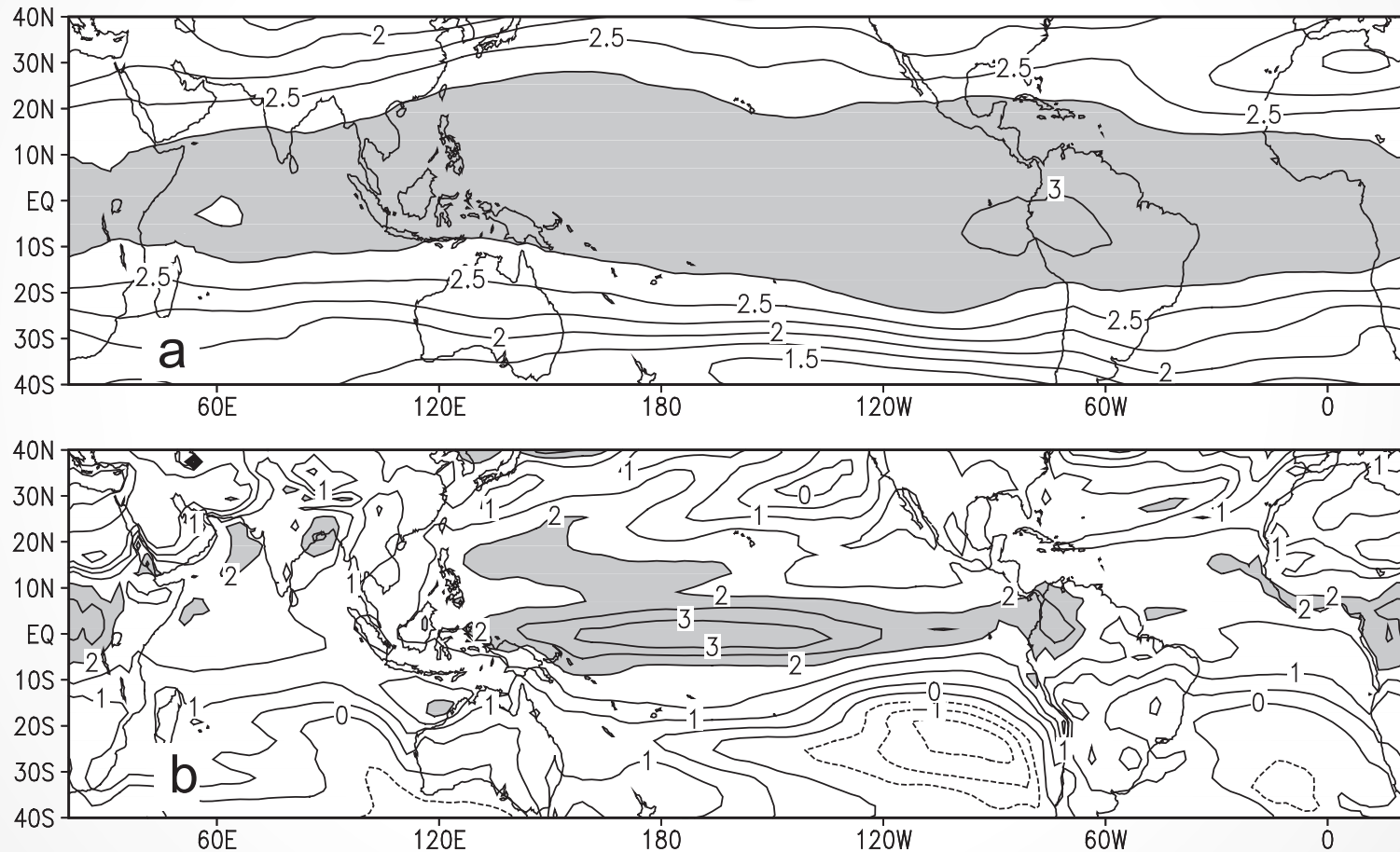
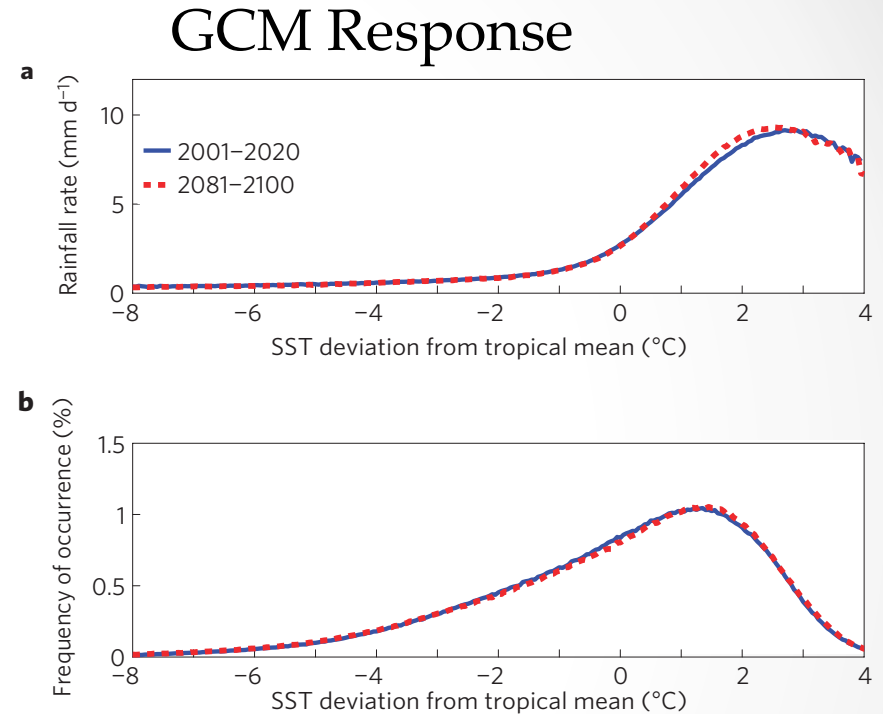
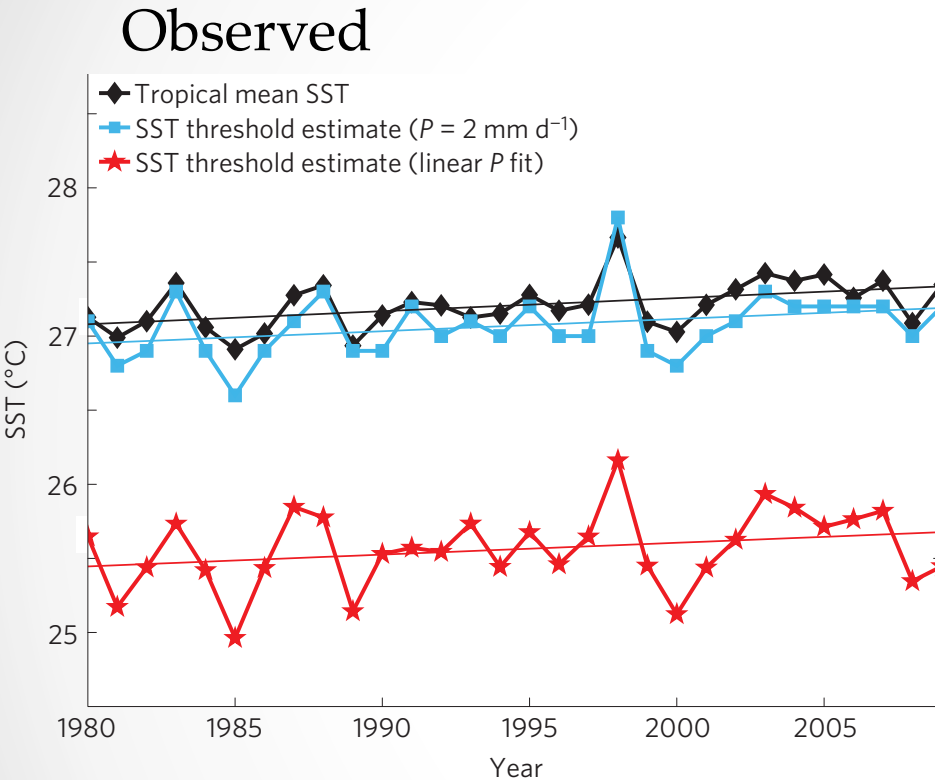


FIG. A1. Annual-mean change in CM2.1 A1B: (a) 300-hPa temperature (shading  $> 2.75$  K) and (b) gross convective instability  $I_M/c_p$  (shading  $> 2$  K).

*Xie et al. (2009, J. Clim.)*

# “Threshold” for strong precipitation seems to follow tropical-mean SST in obs. and models

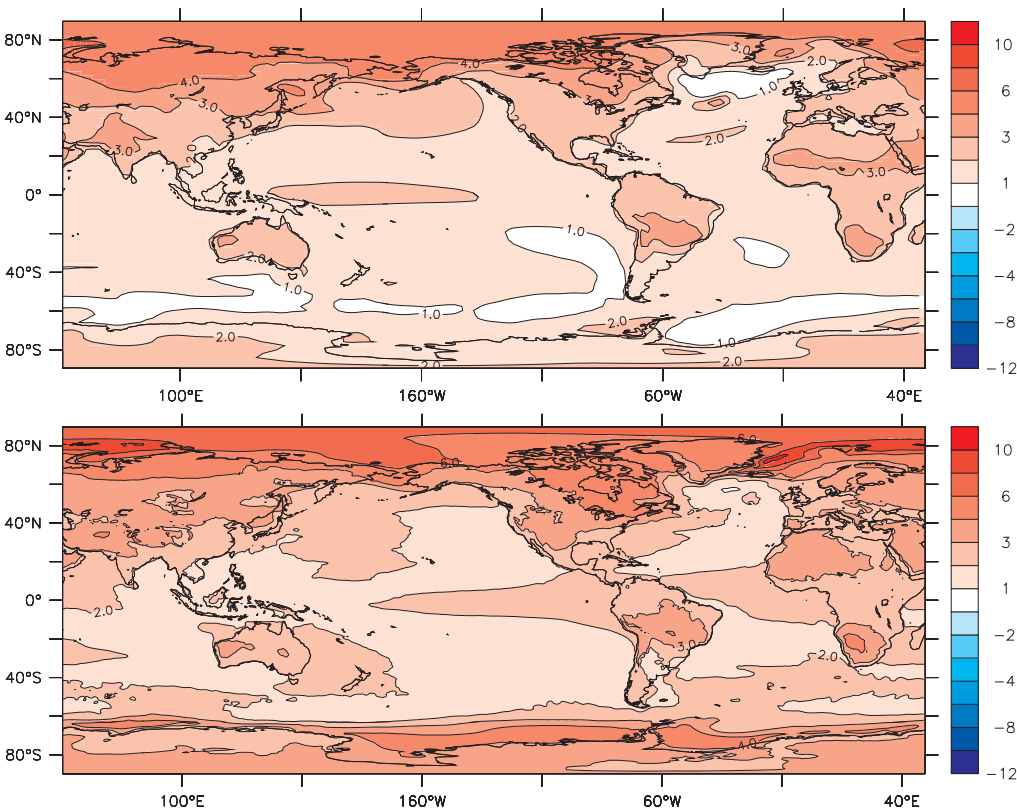


**Figure 3 | Twenty-first-century changes in rainfall rate and SST frequency distributions. a,b**, Ensemble mean rainfall rate as a function of SST (**a**) and SST frequency distribution (**b**) for 2001-2020 (blue, solid) and 2081-2100 (red, dashed) for the ten CMIP3 models of Fig. 2. SST is expressed as the deviation from the 20-year tropical mean.



# Response to 2xCO<sub>2</sub> of models of different resolution:

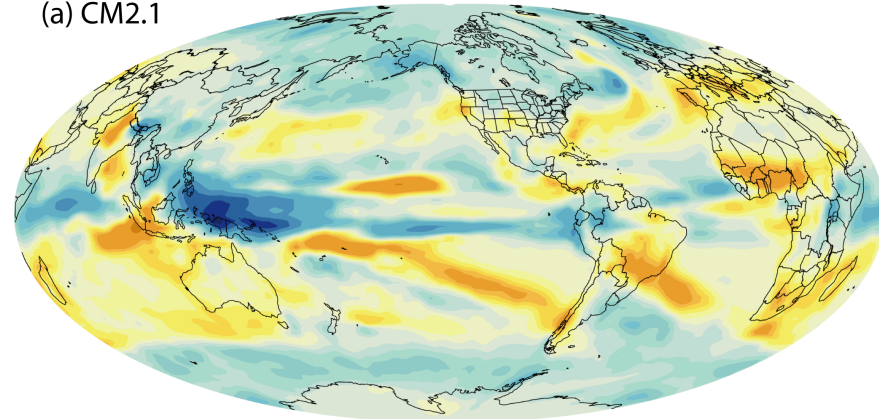
- CM2.1 (1°Ocn, 2.5°Atm)
- CM2.5 (0.25°Ocn, 0.5°Atm)



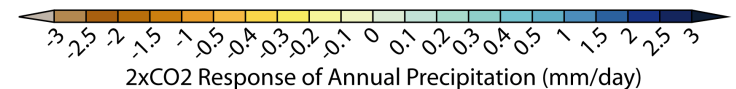
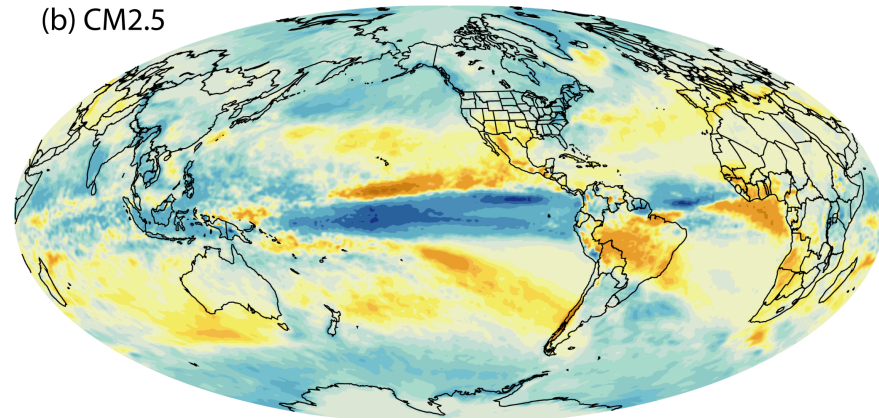
Surface Temperature

## Precipitation

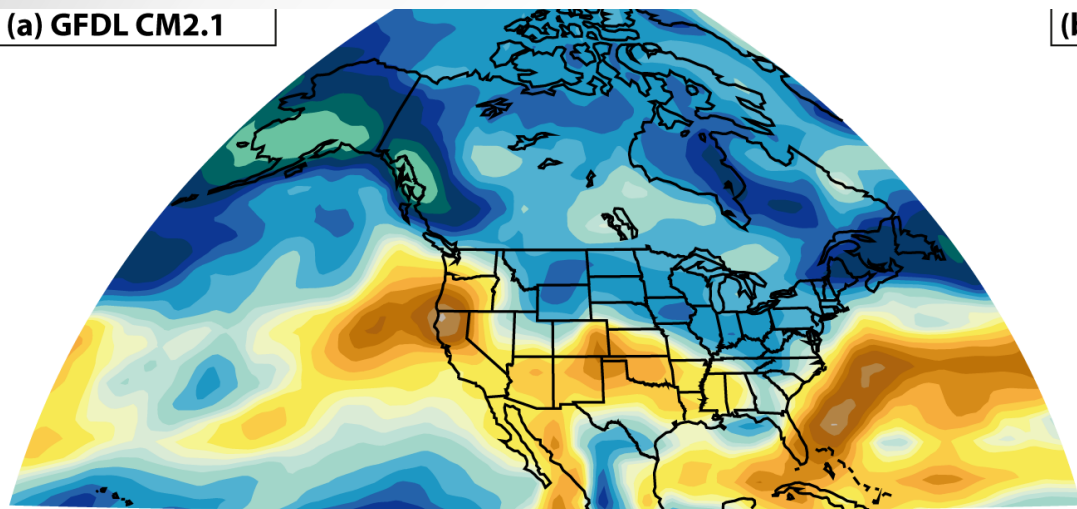
(a) CM2.1



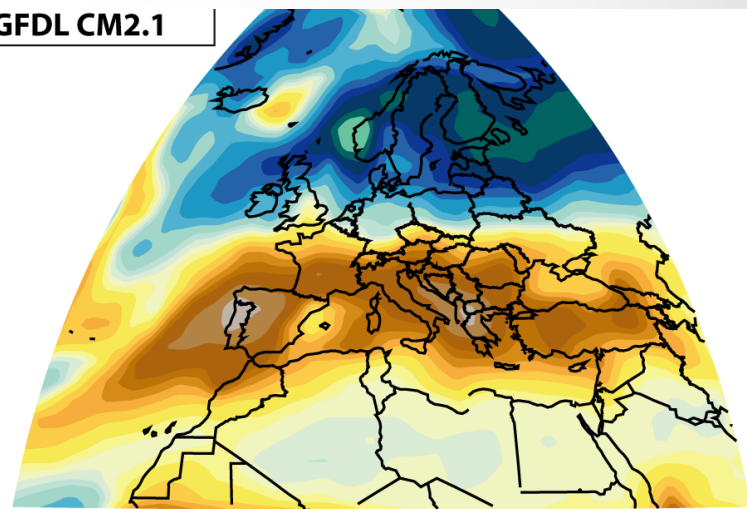
(b) CM2.5



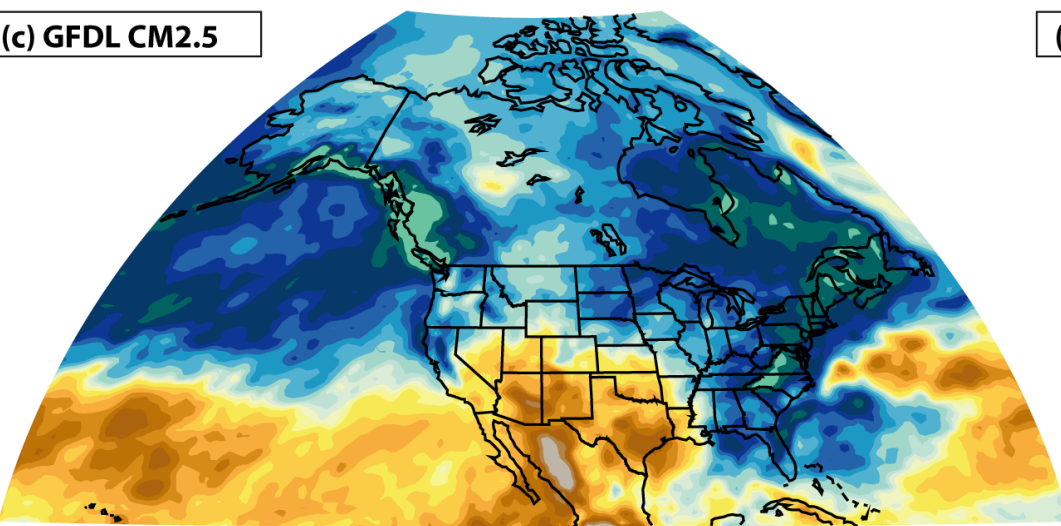
(a) GFDL CM2.1



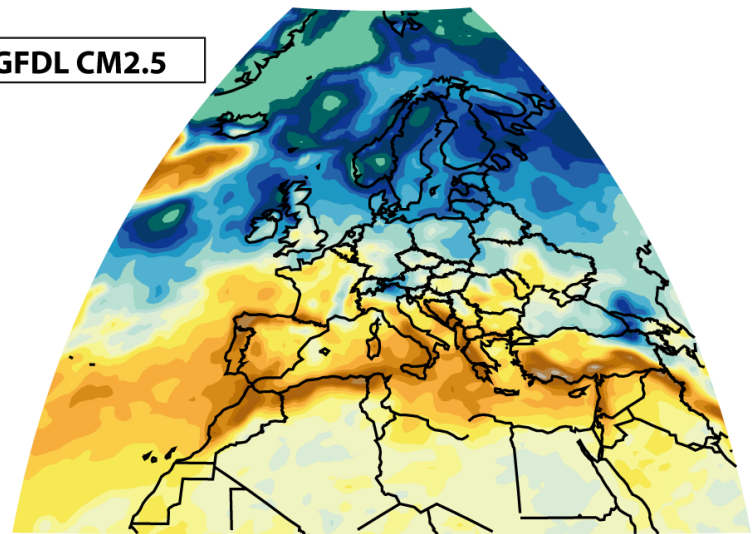
(b) GFDL CM2.1



(c) GFDL CM2.5



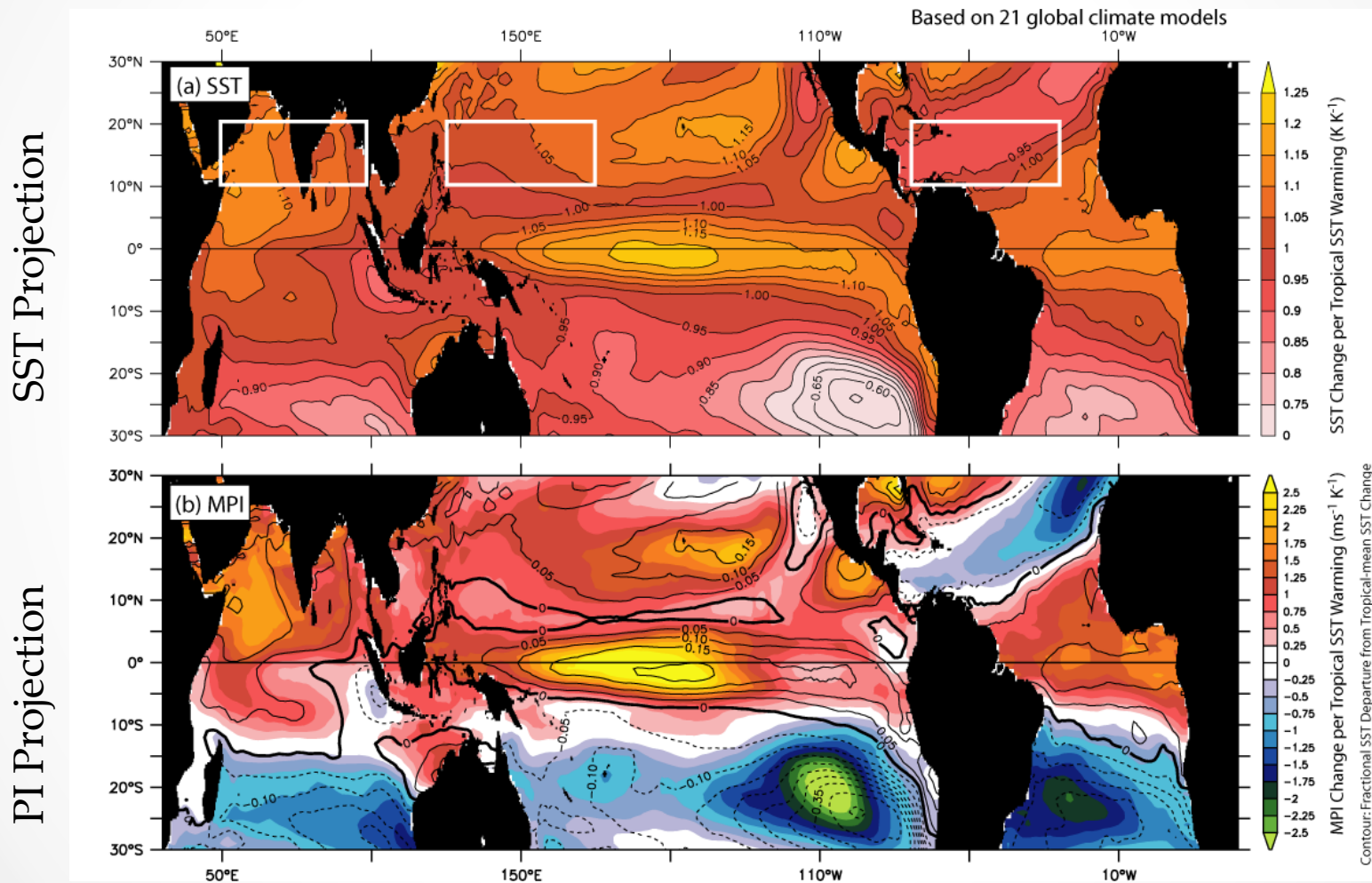
(d) GFDL CM2.5



Annual-mean Precipitation Response to 2xCO<sub>2</sub> (mm/day)

# For related reasons tropical cyclone potential intensity tracks SST relative to tropical-mean, not local SST\*

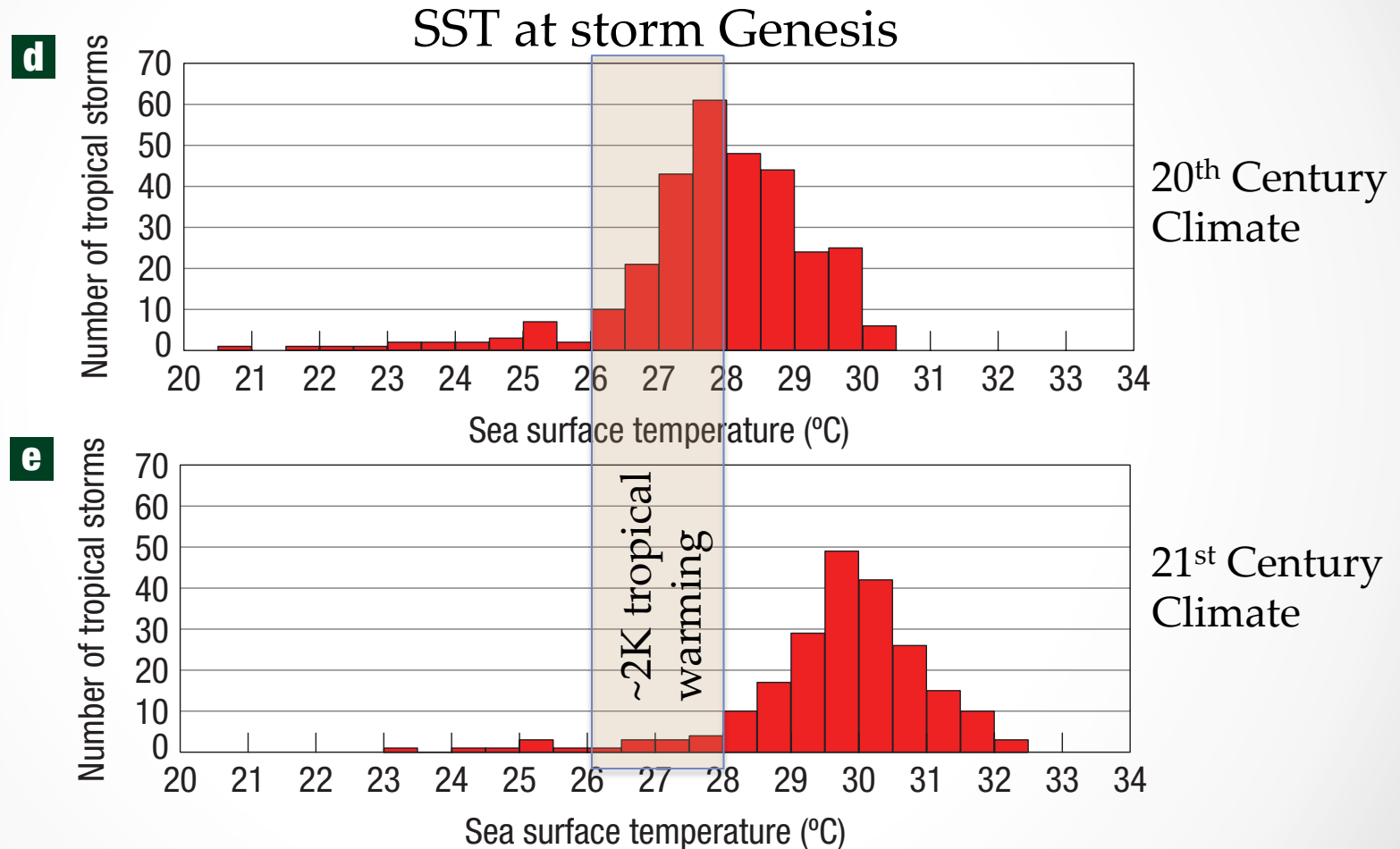
\*global-mean PI changes still need to be explained.



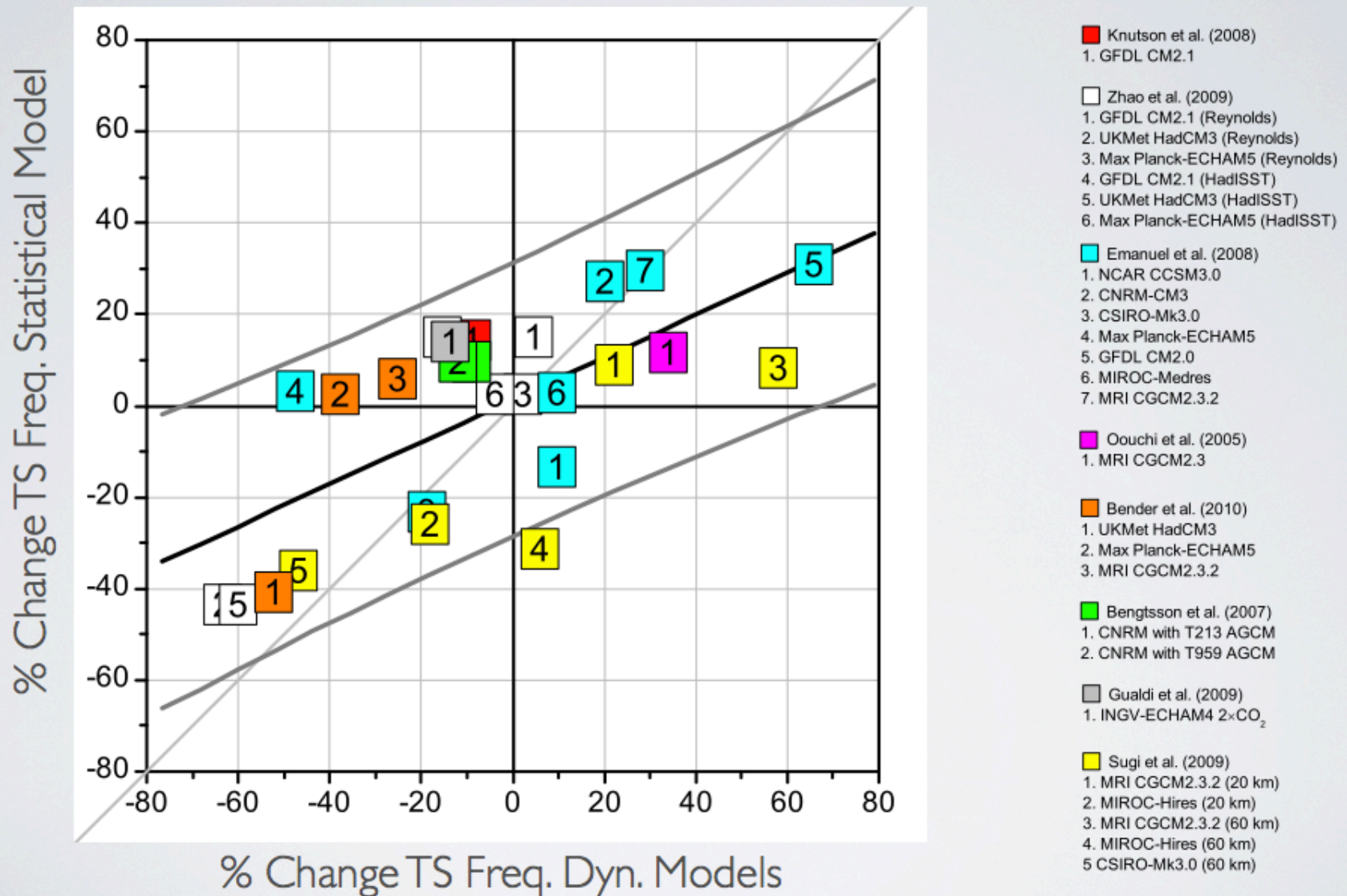
Vecchi and Soden (2007, Nature)



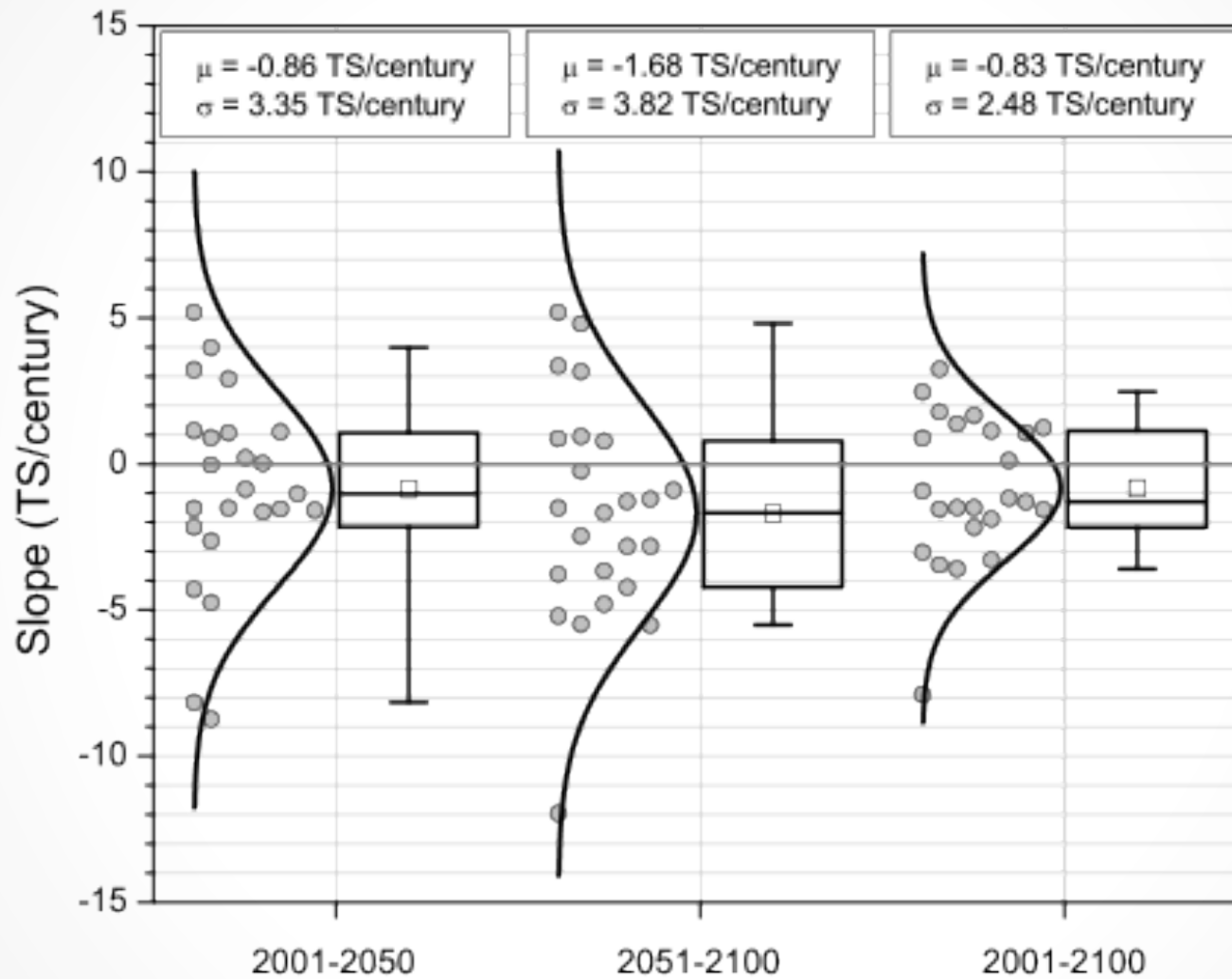
# Modeled “threshold” for hurricane genesis also depends on tropical-mean SST



# DYNAMICAL MODELS EXHIBIT CONSISTENT RELATIONSHIP TO LARGE-SCALE THROUGH STATISTICAL MODEL - ALL CONSISTENT WITH OBSERVATIONS

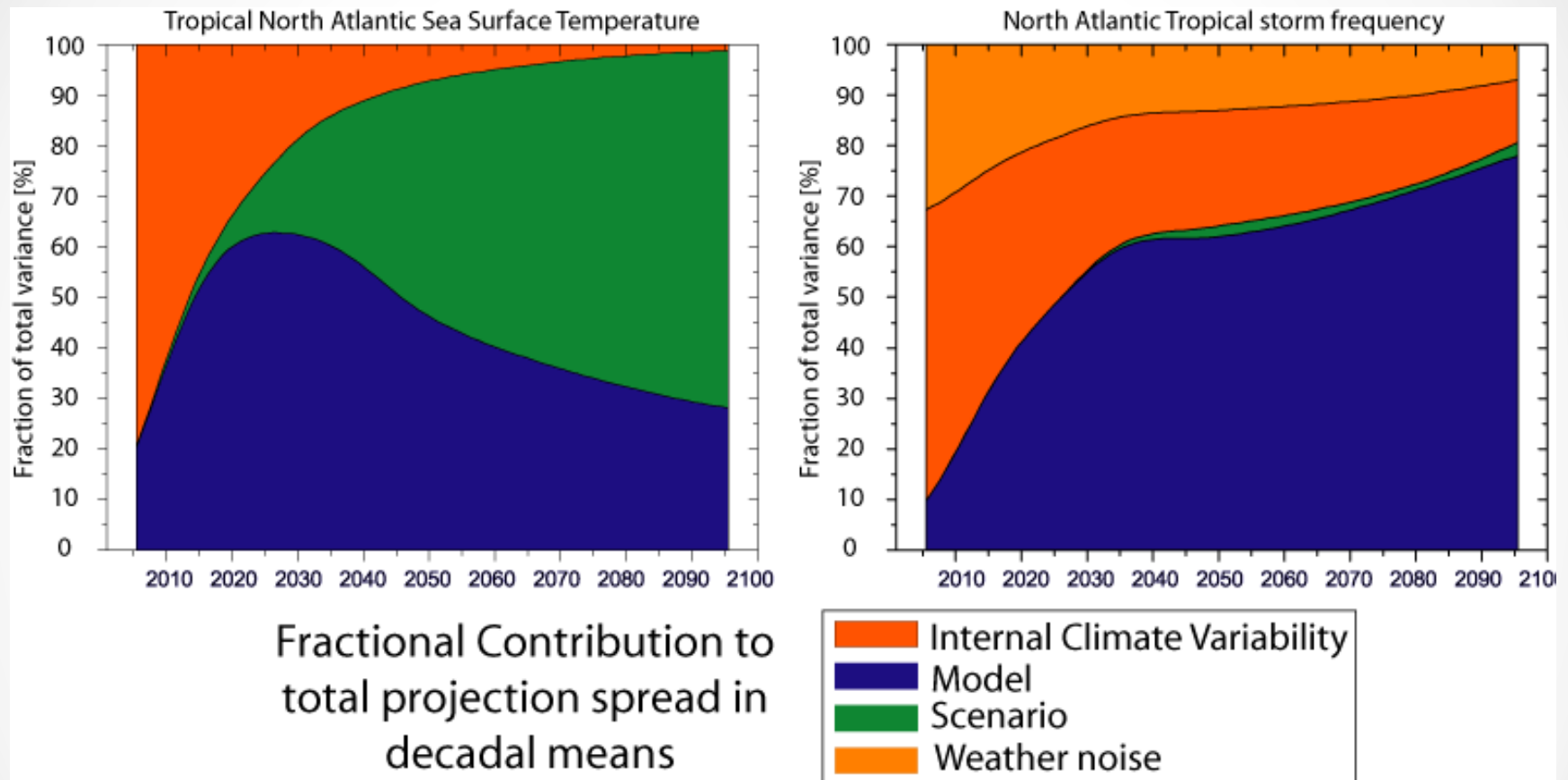


## Statistical Projections of 21st Century NA TS Trends



*Villarini et al (2011, J. Clim.)*

# Sources of uncertainty for relative SST change different from those to regional or global SST change



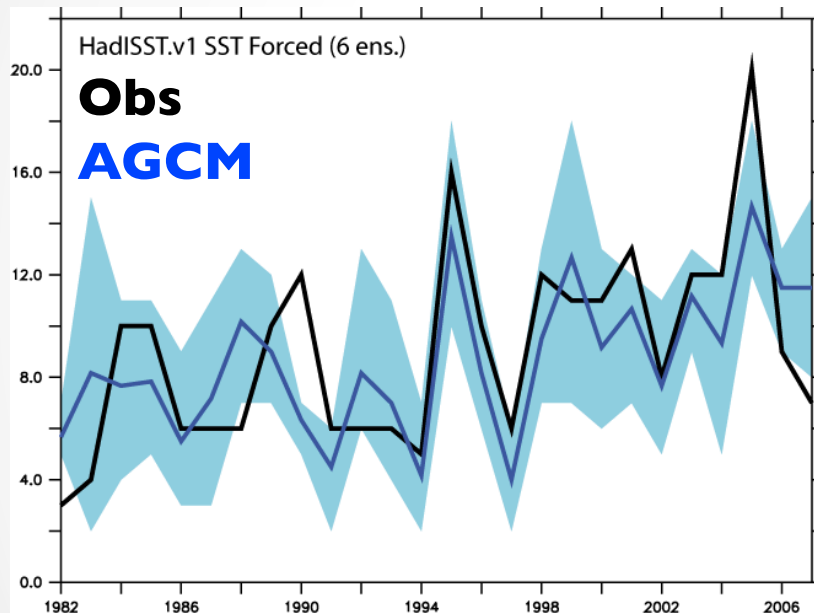
Uncertainties arising from differences in model response to forcing other key source of uncertainty for coming decades (whole century for TS's)



# Because modeled hurricane counts depend on relative SST, trends sensitive to forcing used even over satellite SST era

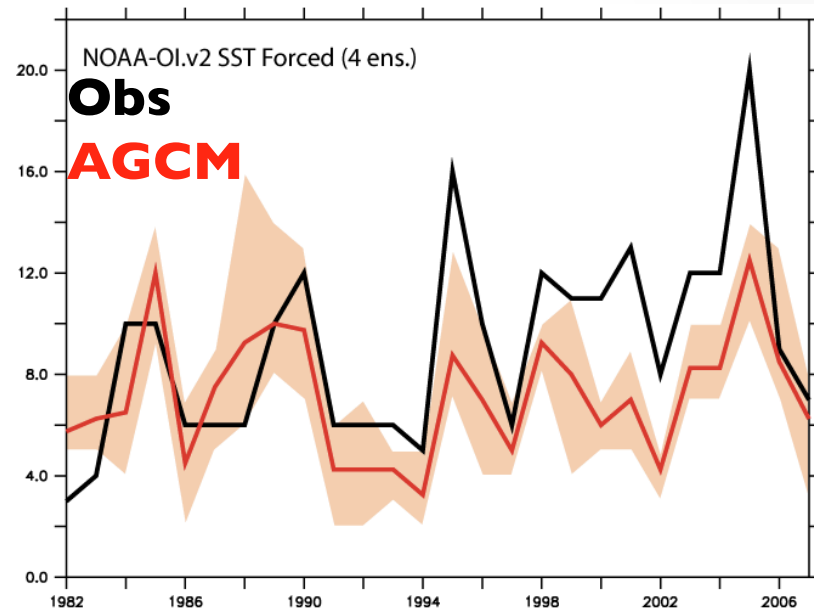
Tropical storm frequency response to same AGCM but different estimates of observed SST

HadISST forced



AGCM is 100km version of Zhao et al (2009, J. Clim.)

NOAA-OI.v2 forced



Vecchi et al (2011, in prep.)

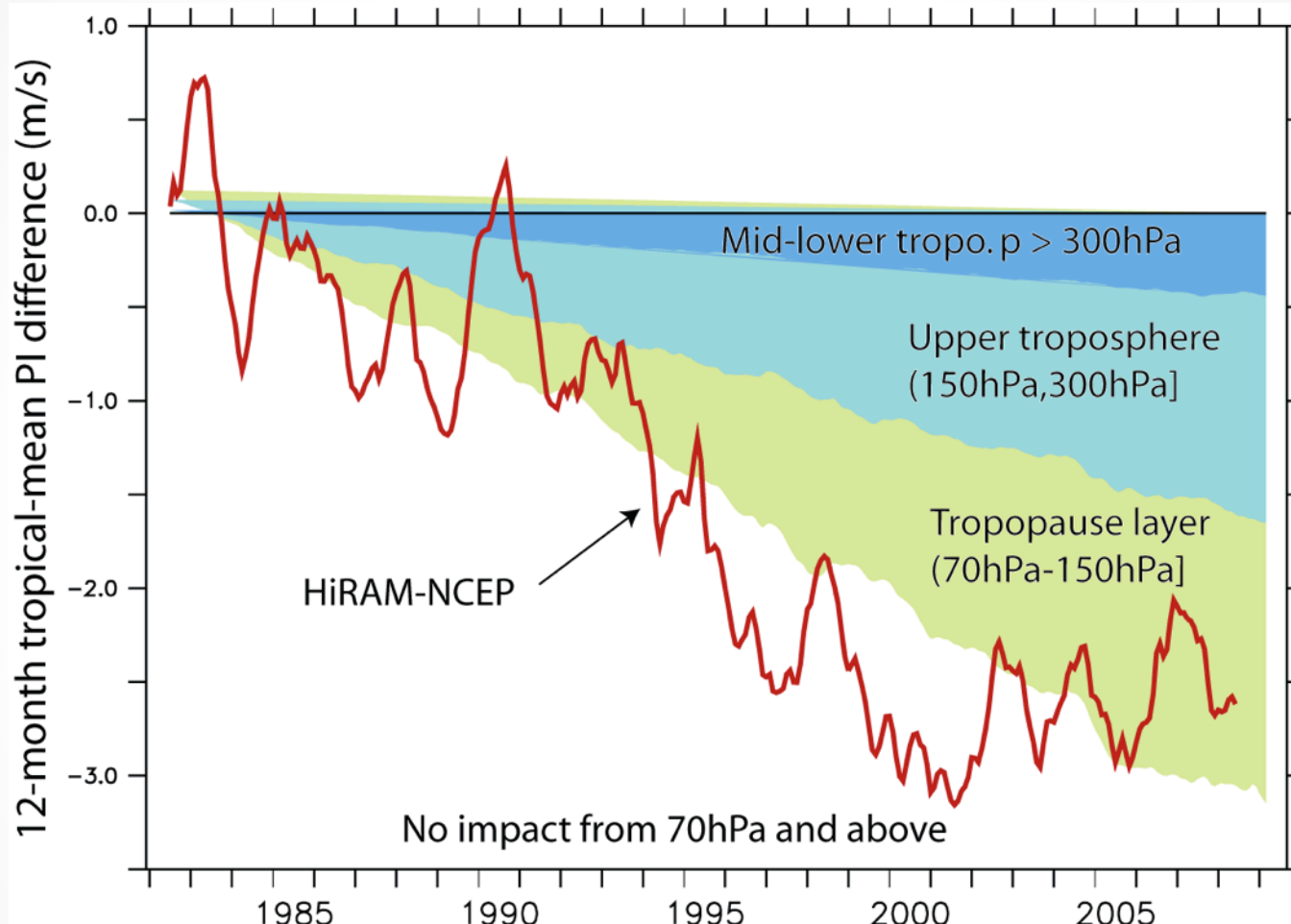
How do we confidently evaluate model skill in this context?

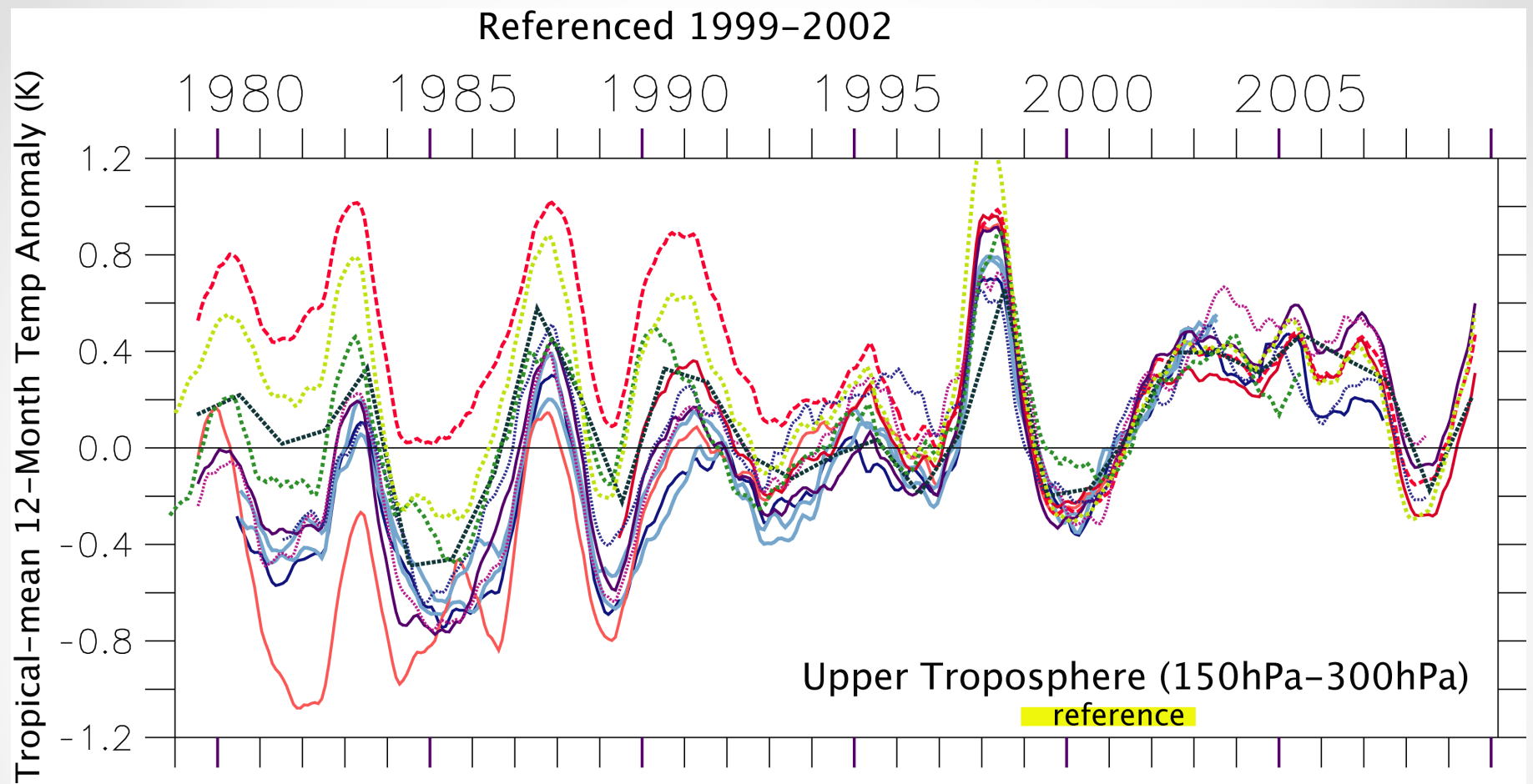
# Vertical changes

Changes in atmospheric stability impact potential intensity. Uncertainty in past changes to upper tropospheric temperature limit confidence in past (and future) PI changes.

Ocean near-surface thermal stratification key control on intensity of strongest cyclones. How does it change?

- PI Impact of swapping NCEP T-Trend with HiRAM T-Trend





SST+Rad. -Forced AGCMs

- ..... HiRAM-C180 (SST-CO<sub>2</sub>-O<sub>3</sub>)
- HiRAM-C90 (SST-CO<sub>2</sub>-O<sub>3</sub>)
- HiRAM-C90 (SST-CO<sub>2</sub>-O<sub>3</sub>-Volcano)

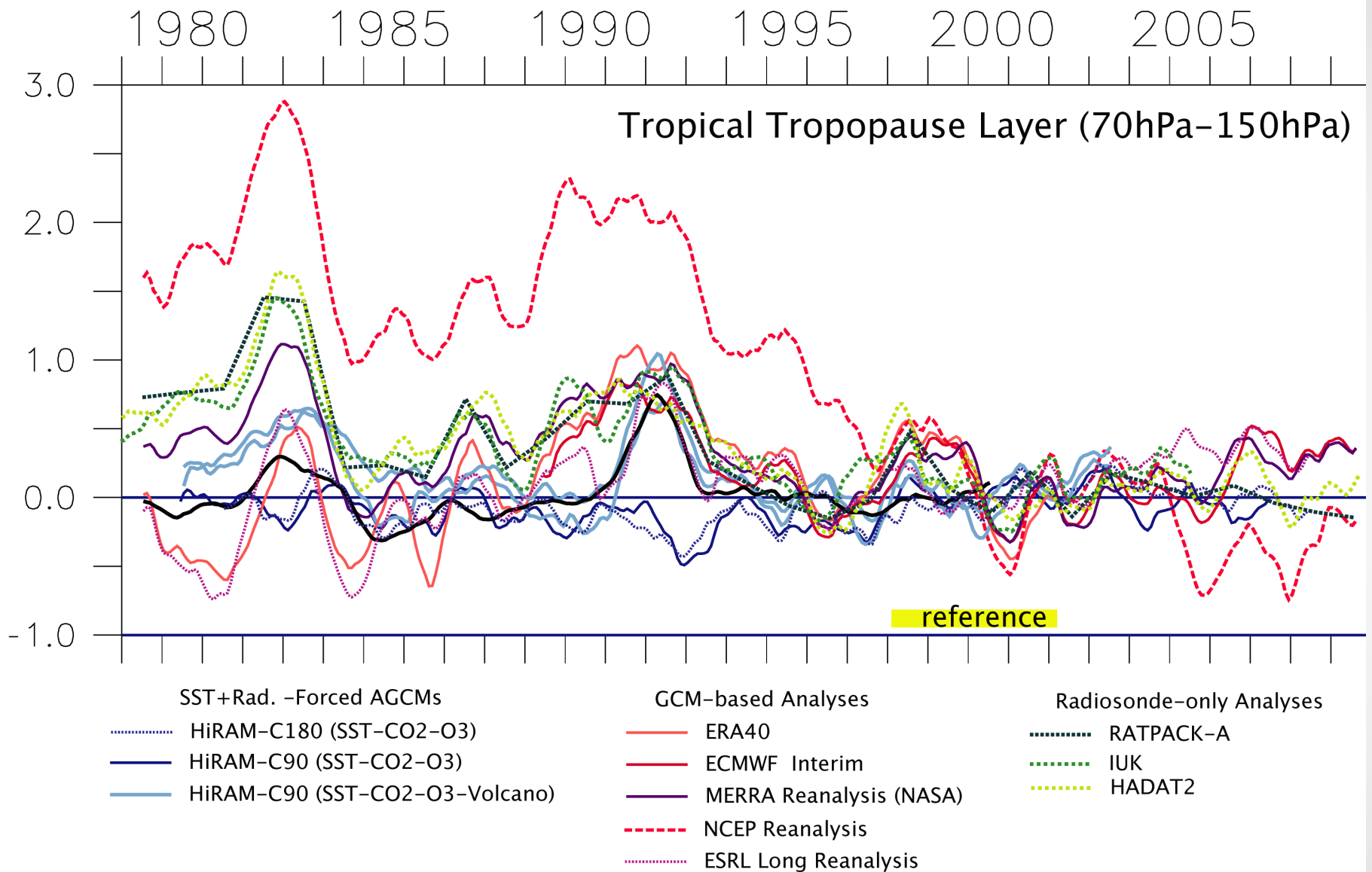
GCM-based Analyses

- ERA40
- ECMWF Interim
- MERRA Reanalysis (NASA)
- - - NCEP Reanalysis
- ..... ESRL Long Reanalysis

Radiosonde-only Analyses

- - - - - RATPACK-A
- ..... IUK
- ..... HADAT2

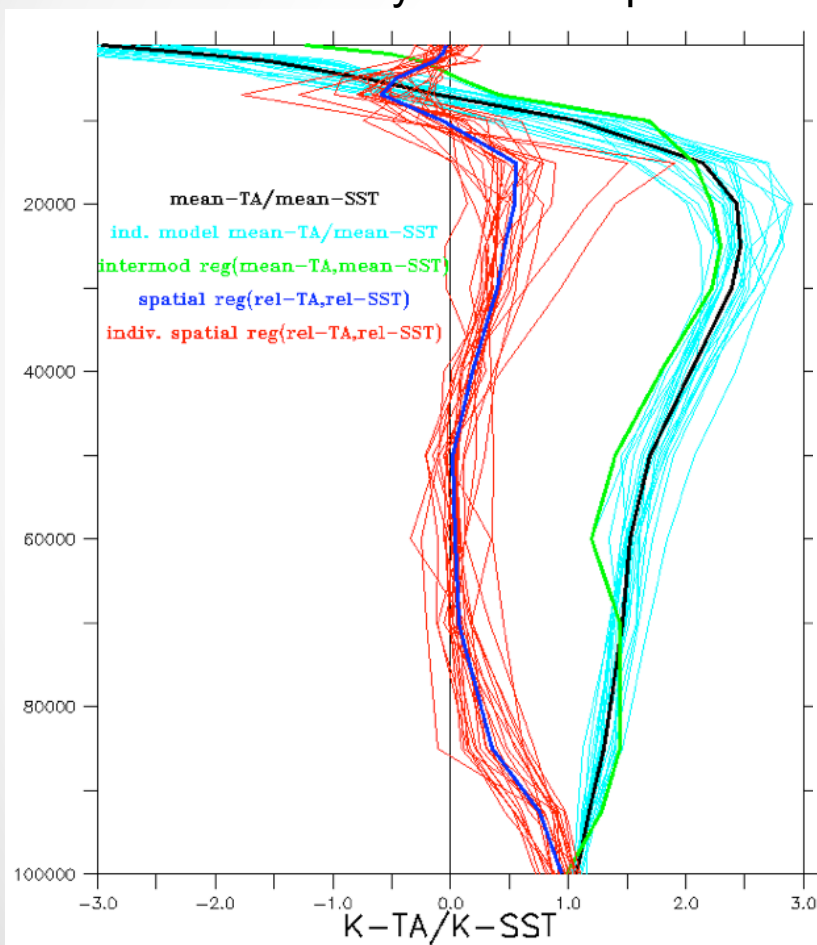
Referenced 1999–2002



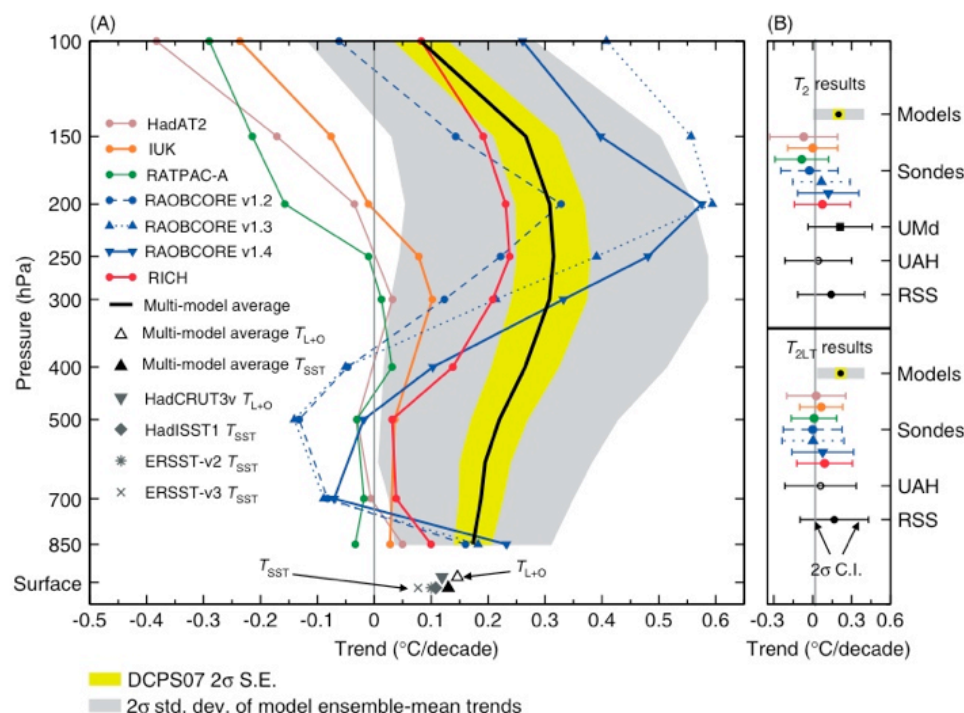
Do we know vertical structure of warming well enough?  
 Relatively small inter-model TA trend spread still gives large  $\langle \text{PI} \rangle$  spread ( $\pm 2\text{m/s}$ )

For now: assume GCM lapse rate response “reasonable” estimate?

IPCC-AR4 21st Cy model response



Observational estimates and  
 IPCC-AR4 historical runs

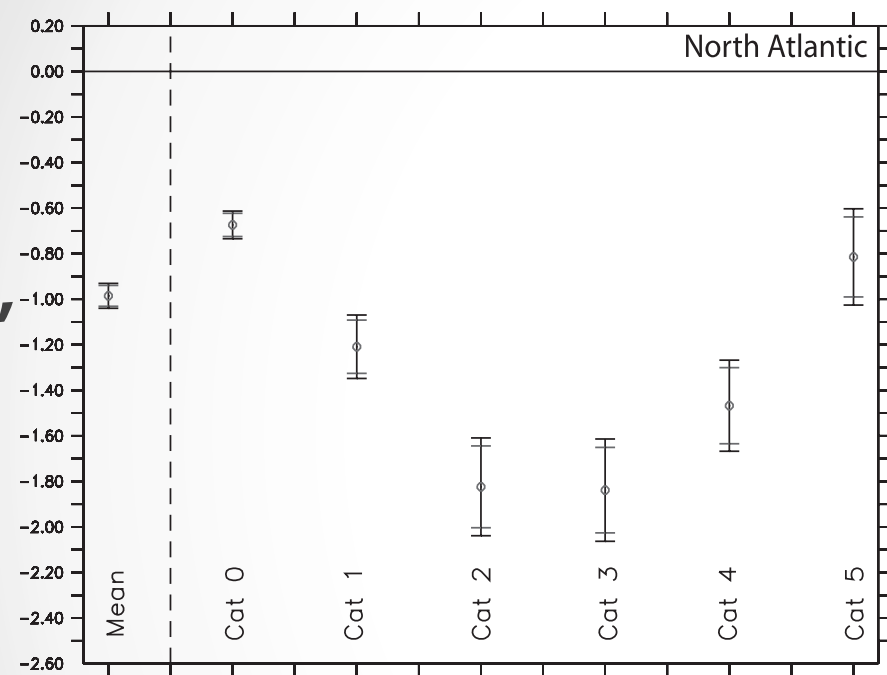


Santer et al (2008)



# Intensity of most intense hurricanes impacted by ocean stratification

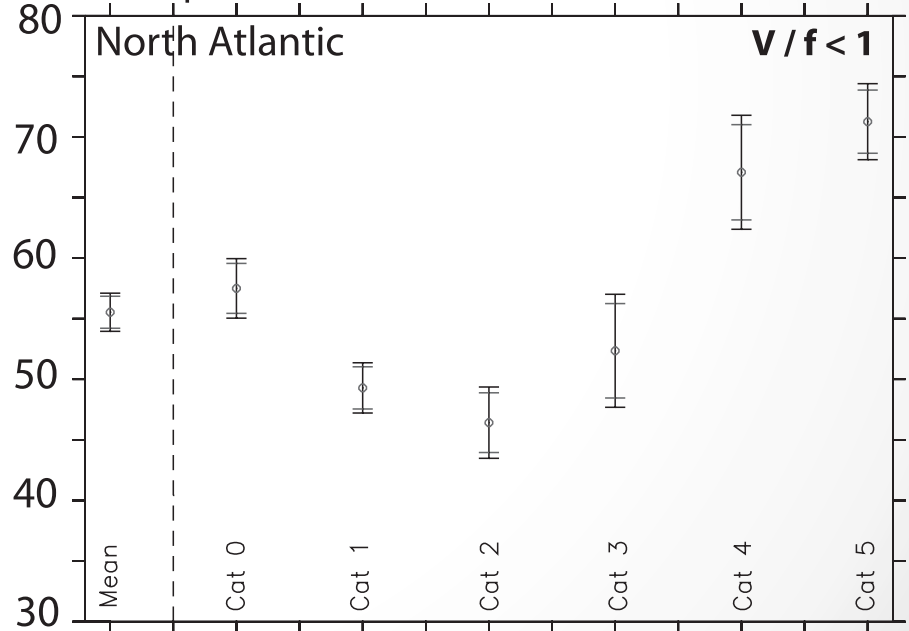
a) SST anom (K), for Day +2 minus Day (-12 to -2 average)



Strongest hurricanes tend to occur over water with light stratification (deep mixed layer).

Ocean cooling following tropical cyclones has non-monotonic relationship to intensity, indicating ocean influence.

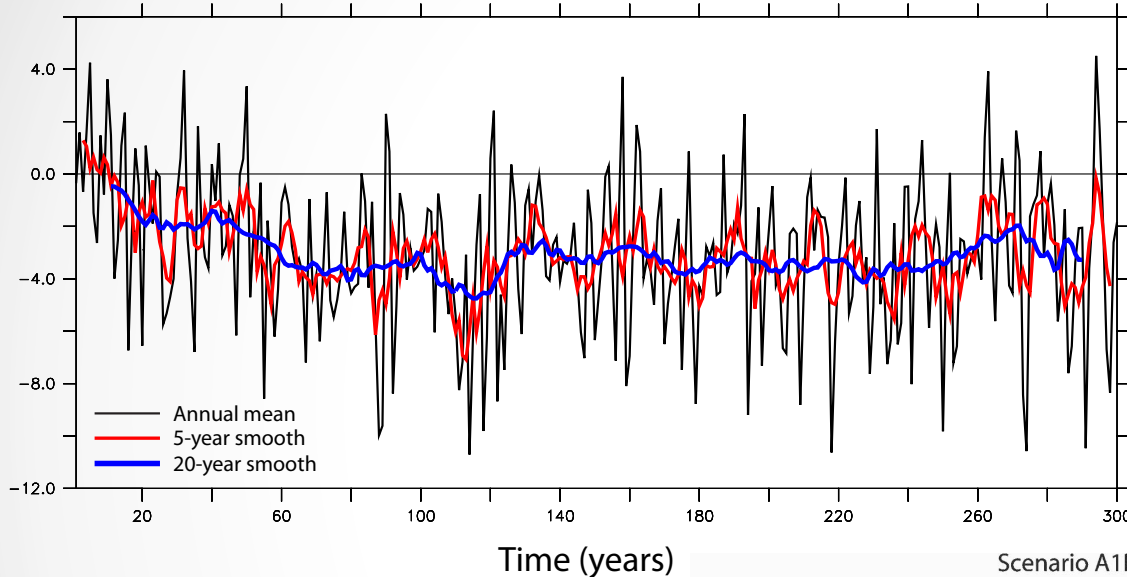
b) Depth of 2K stratification (m), CDA 2002-2007





# MLD should decrease in mean from CO<sub>2</sub>, but what about regional patterns?

CM2M: Change in 2k stratification-depth (m) for 2XCO<sub>2</sub> - Control



Thermodynamically-driven shoaling of mean mixed-layer

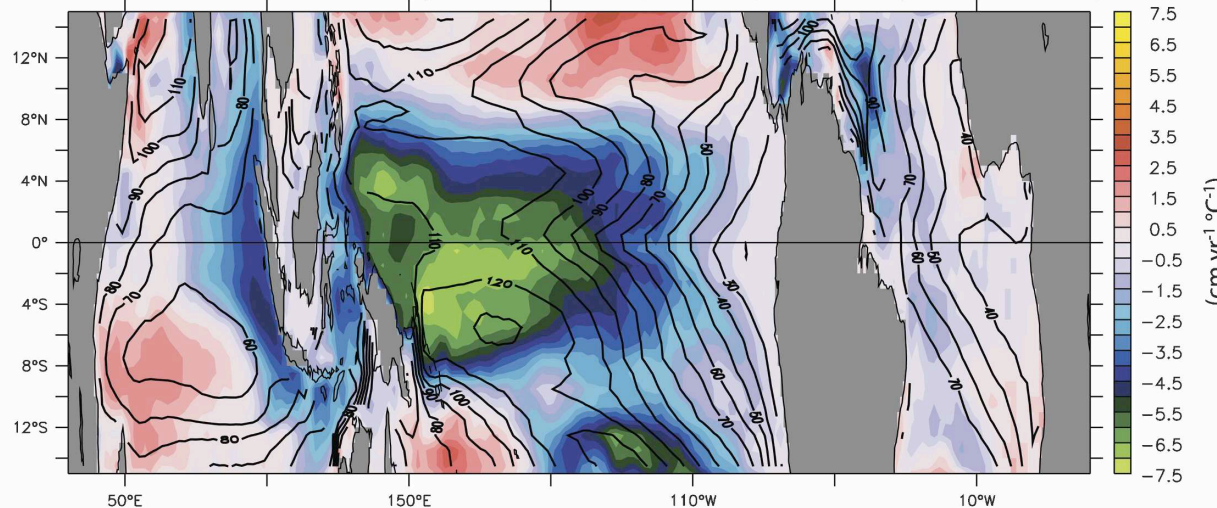
*Lloyd (2011, PhD Thesis)*

Changes in circulation can lead to local thermocline deepening

*Vecchi and Soden (2007, J. Clim.)*

Scenario A1B (720 ppm CO<sub>2</sub> Stabilization) - 2001-2100

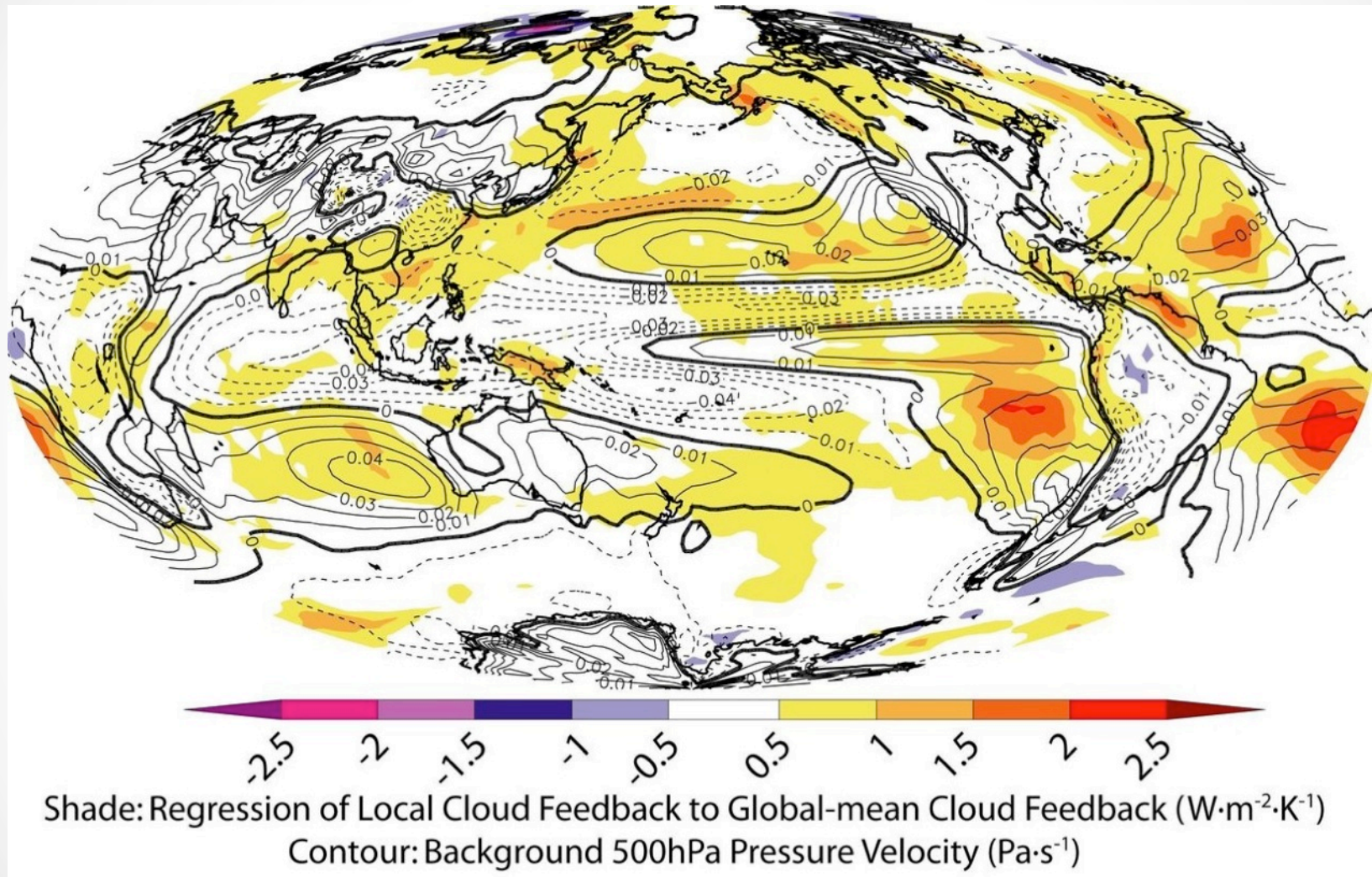
19-Model Ensemble-mean 100-year Thermocline Depth Trend (Normalized by Global SST Trend)



# Summary

- For regional tropical precipitation and tropical cyclones the “patterns” of temperature change may be more relevant:
  - Patterns in space: SST change *relative* to tropical-mean key to tropical precip and cyclone response
  - Patterns in height/depth (atmospheric stability and ocean stratification changes)
  - Patterns in time (multi-year to decadal changes)
- The past character of “patterns” of change is less well known.
- The potential mechanisms impacting “patterns” more varied (redistribution, non-uniform radiative forcing, non-uniform response to uniform forcing, etc.

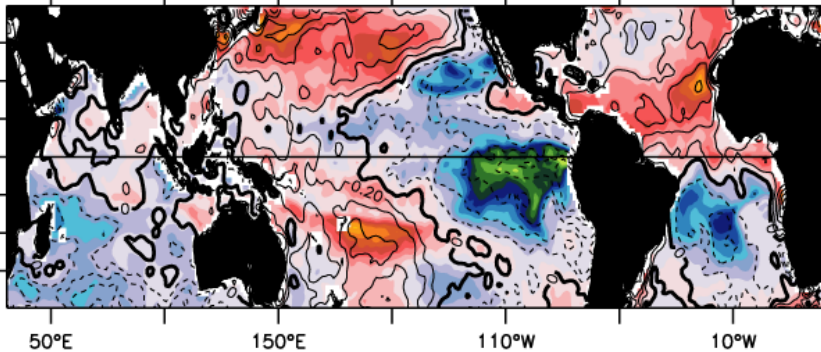
# Regional changes important to global mean: Cloud Feedbacks



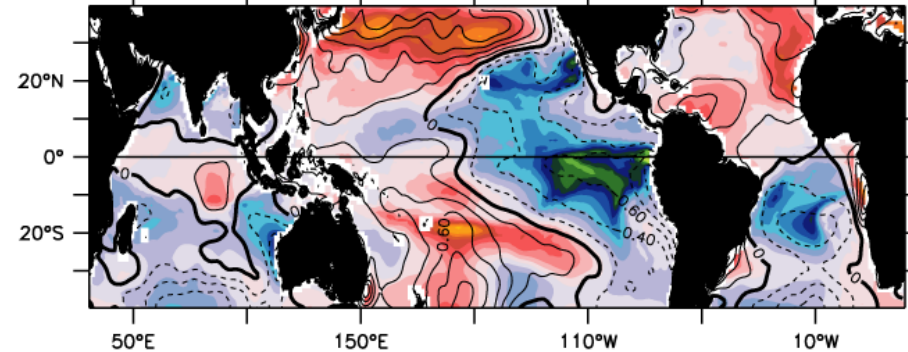


# Differences in HiRAM/NCEP Tropical-mean lapse rate trends largely explain differences in 1982-2006 potential intensity trends

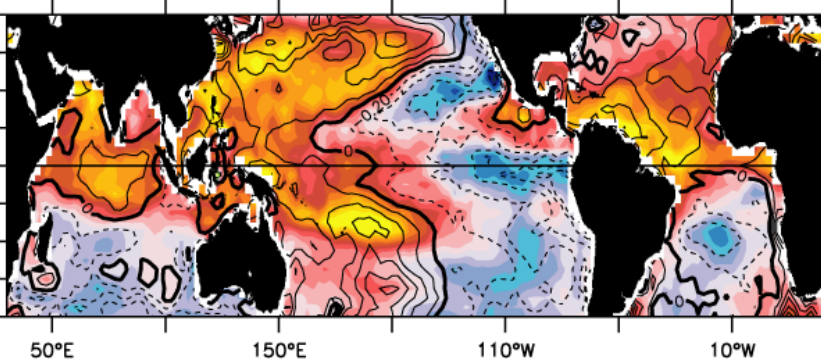
a) HadISST-forced AGCM 1: 30°S-30°N Mean = -0.29 m/s/25 years



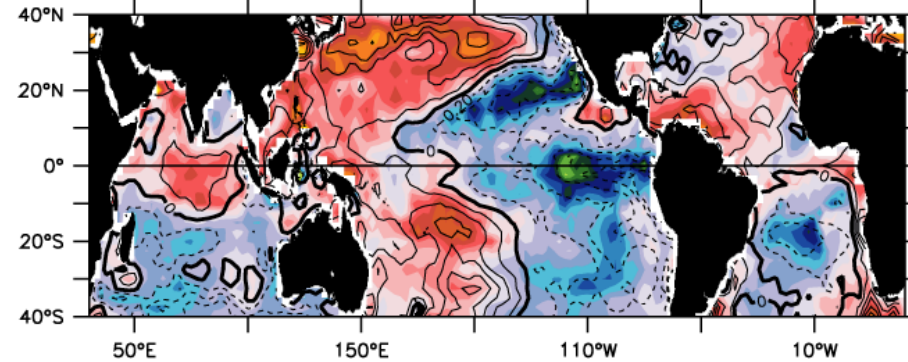
b) ERSST-forced AGCM (see text): 30°S-30°N Mean = 0.60 m/s/25 years



c) NCEP Reanalysis 1: 30°S-30°N Mean = 3.32 m/s/25 years



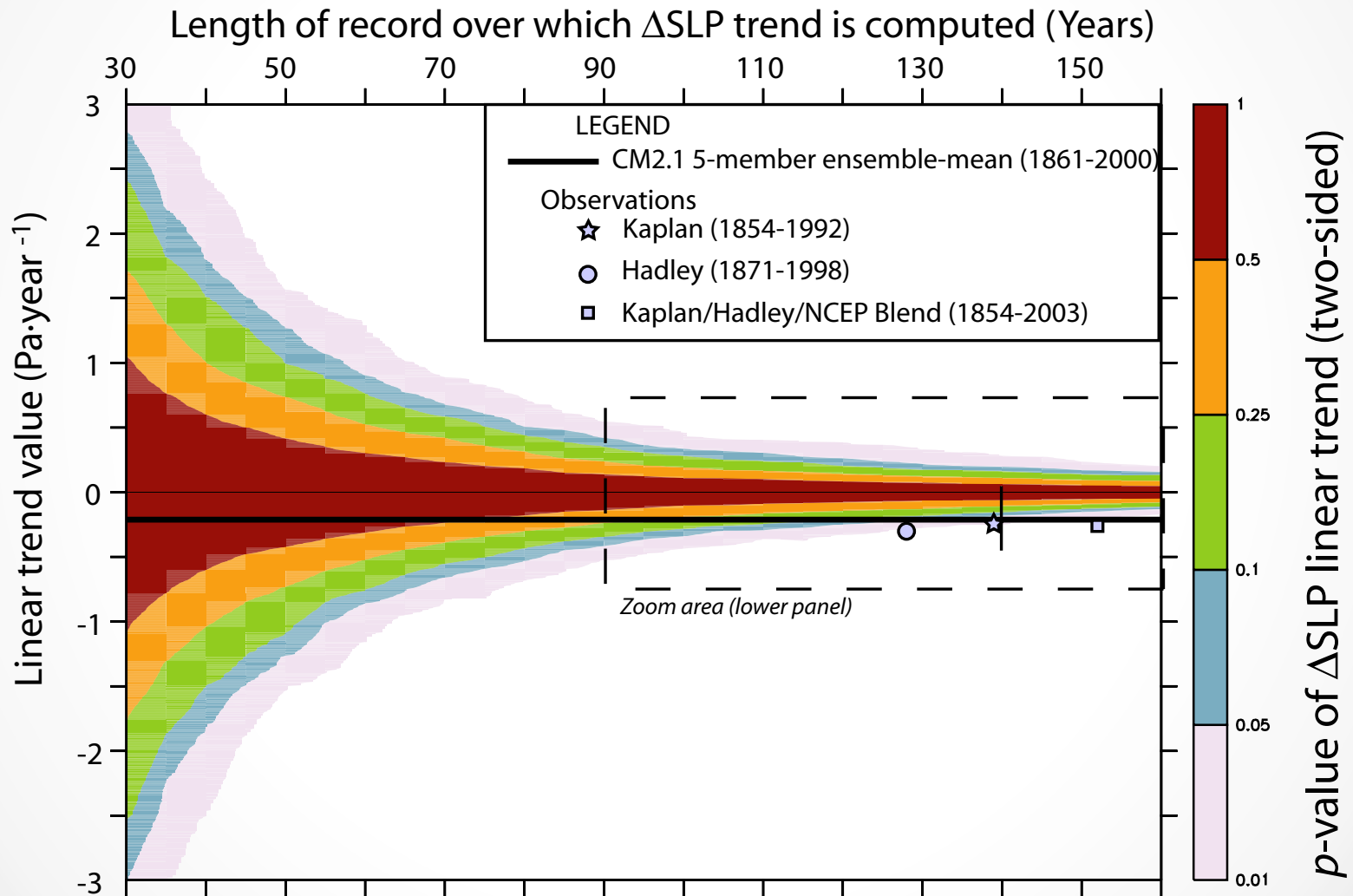
d) NCEP Reanalysis 1 Adjusted to AGCM (see text): 30°S-30°N Mean = 0.0 m/s/25 years



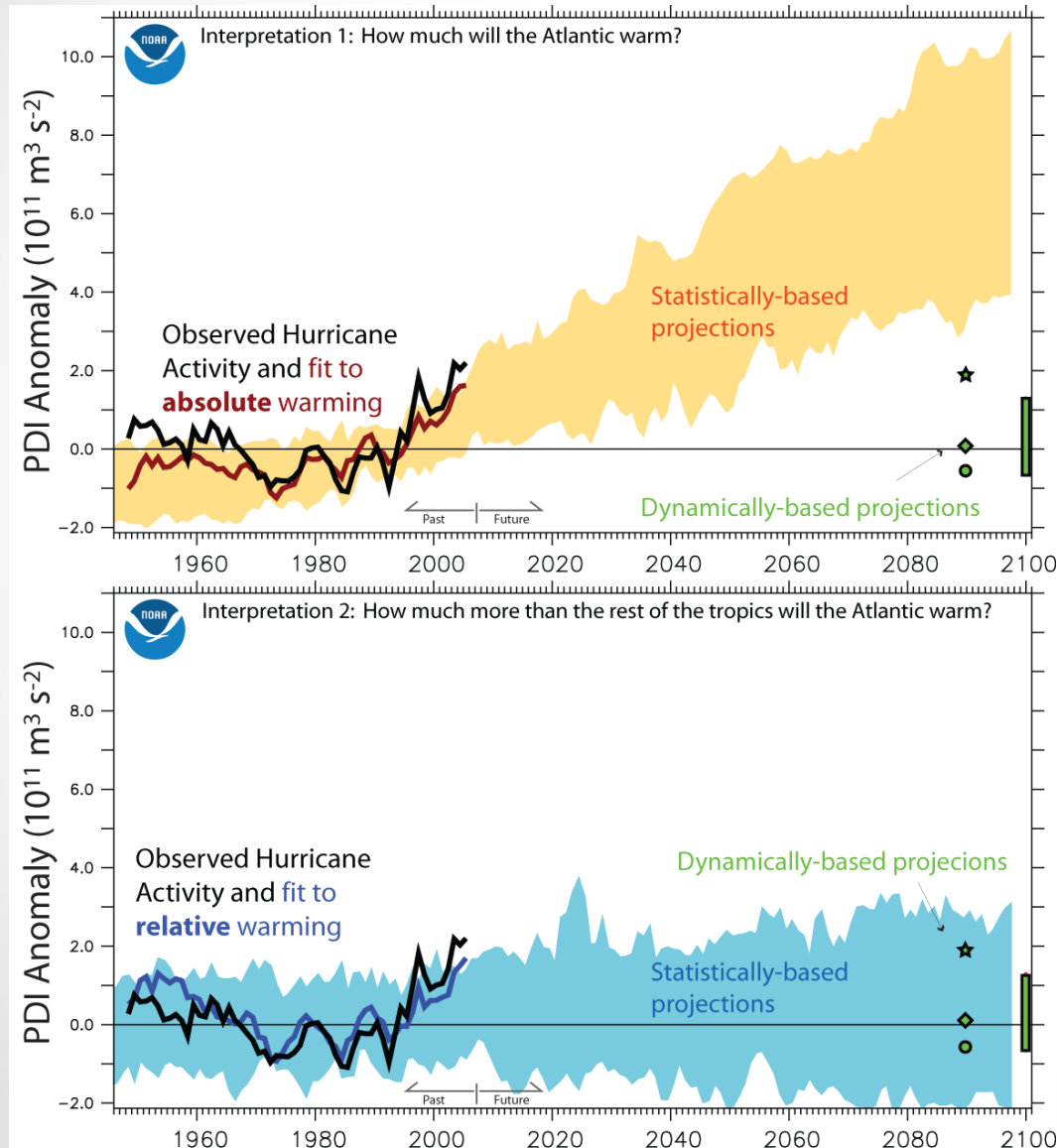
1982-2007 Linear Trend in June-November Bister and Emanuel PI (m/s/ 25 Year)



# On multi-decadal and shorter timescales internal variability can dominate Walker Circulation changes



Both absolute and relative SST consistent with recent hurricane changes, only relative SST consistent with dynamical models (*e.g.*, Vecchi et al. 2008, Villarini et al. 2011) and homogenized century scale hurricane records (Villarini et al. 2010)



Observed Activity  
Absolute Atlantic  
Temperature

Dynamical Model  
Projections

Observed Activity  
Relative Atlantic  
Temperature

Vecchi, Swanson and Soden  
(2008, Science)