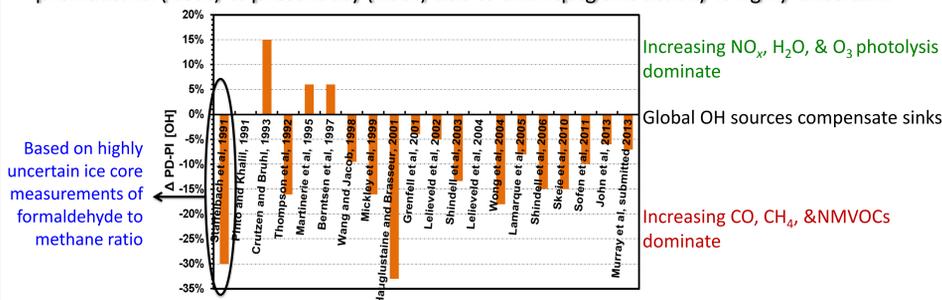


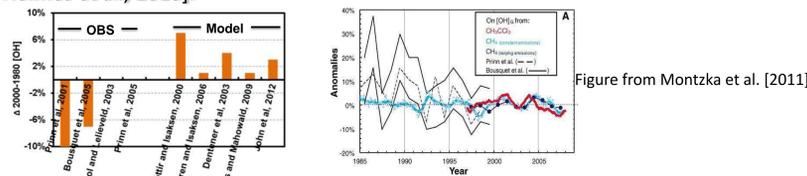
1. Motivation

Accurate measurements and modeling of tropospheric hydroxyl (OH) are key to understanding photochemical oxidation and better quantify the lifetimes of climate relevant gases, such as methane (CH₄), hydrochlorofluorocarbons (HCFCs) and hydrofluorocarbons (HFCs)

Based on previous literature, the extent to which global mean tropospheric OH has changed from preindustrial (1850) to present day (2000) due to anthropogenic activity is highly uncertain.



OH change from 1880 to 2000 inferred from methyl chloroform (CH₃CCl₃) measurements in previous studies is inconsistent with that simulated by models. However, recent analysis of CH₃CCl₃ data post 1997 indicates that global OH is generally well-buffered against perturbations [Montzka et al., 2011; Holmes et al., 2013].



Objectives

- Evaluate present day global OH in current generation chemistry-climate models (CCMs)
- Explore changes in OH and methane lifetime for present day (2000) relative to preindustrial (1850) and to 1880
- Investigate the impact of individual factors (climate and emissions) in driving preindustrial to present day OH changes

2. The Atmospheric Chemistry & Climate Model Intercomparison Project (ACCMIP)

CESM-CAM-superfast (CE)	LLNL-NCAR, USA
CICERO-OsloCTM2 (CI)*	CICERO, Norway
CMAM (CM)	CCCma, Canada
GEOSCCM (GE)	NASA GSFC, USA
GFDL-AM3 (GF)	NOAA GFDL, USA
GISS-E2-R (GI)	NASA GISS, USA
GISS-E2-R-TOMAS (GT)	NASA GISS, USA
HadGEM2 (HA)	UKMO, UK
LMDzORINCA (LM)	LSE, France
MIROC-CHEM (MI)	NIES, Japan
MOCAGE (MO)*	MeteoFrance, France
NCAR-CAM3.5 (NC)	NCAR, USA
STOC-HadAM3 (ST)	Univ. Edinburgh, UK
TM5 (TM)*	RNMI, The Netherlands
UM-CAM (UM)	NIWA, New Zealand

* CTMs

CCMs mostly driven by SST/SIC from parent coupled models

WMGHGs mostly from Meinshausen et al. [2011], CH₄ concentrations specified at the surface in most models

Anthropogenic emissions from Lamarque et al. [2010], diverse natural emissions

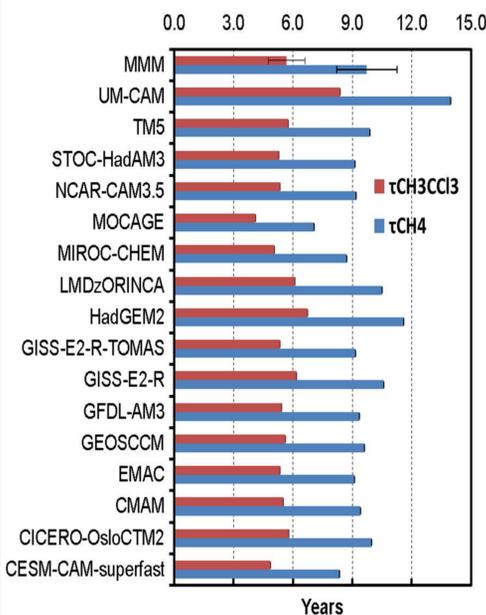
Diverse representation of stratospheric O₃ and its influence on photolysis, and tropospheric chemistry mechanisms

Time slice simulations (run for ~4 to 10 years): 1850, 1980, 2000

Fixed 2000 emission and 1850 climate

Fixed 2000 climate with CH₄ conc. and NO_x, CO, NMVOCs emissions set to 1850 individually

Present day Tropospheric CH₄ Lifetime: Large spread attributed to model-to-model differences in NMVOC emissions and simulated ozone photolysis [Voulgarakis et al., 2013]



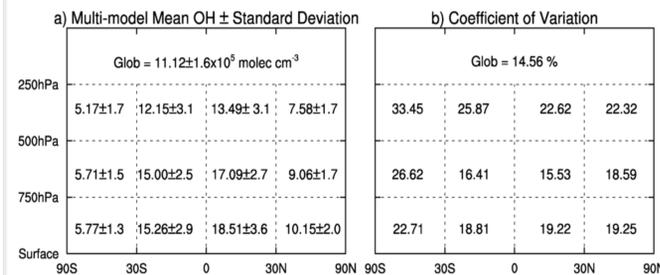
Present day multimodel mean (MMM) τ_{CH_4} and $\tau_{CH_3CCl_3}$ are 5-10% lower than observation-based estimates but within the uncertainty range

MMM $\tau_{CH_3CCl_3} = 5.7 \pm 0.9$ years
 Prinn et al. [2005] $\tau_{CH_3CCl_3} = 6.0^{+0.5}_{-0.4}$
 Prather et al. [2012] $\tau_{CH_3CCl_3} = 6.3 \pm 0.4$

MMM $\tau_{CH_4} = 9.7 \pm 1.5$ years
 Prinn et al. [2005] $\tau_{CH_4} = 10.2^{+0.9}_{-0.7}$
 Prather et al. [2012] $\tau_{CH_4} = 11.2 \pm 1.3$

MMM [OH] = $11.1 \pm 1.6 \times 10^5$ molec cm⁻³ is overestimated by 5-10% but is within the range of uncertainties

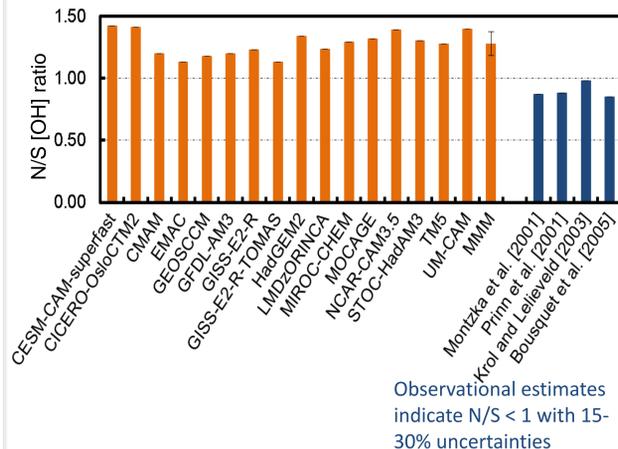
3. Regional Distribution of Present day OH: Large spread but models capture the general spatial patterns



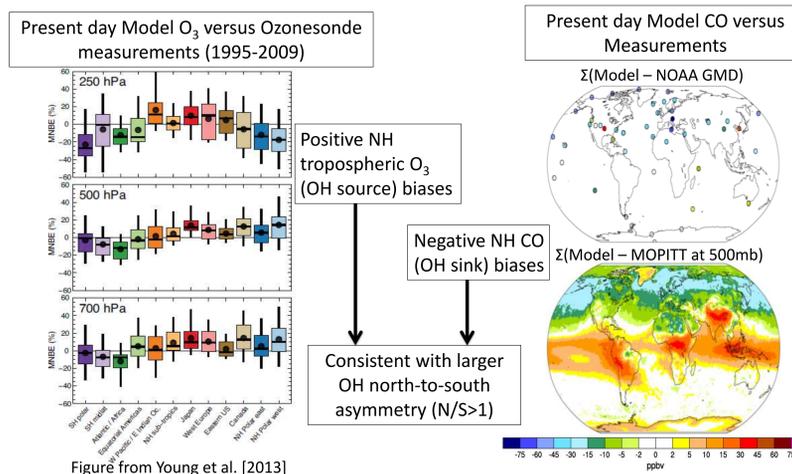
Greater diversity in upper troposphere than in mid to lower troposphere possibly reflecting biases in water vapor and clouds

Diversity is also greater in the Southern Hemisphere (SH) than the Northern Hemisphere (NH) reflecting the intermodel diversity in simulated (or prescribed) stratospheric ozone and its influence on tropospheric ozone photolysis

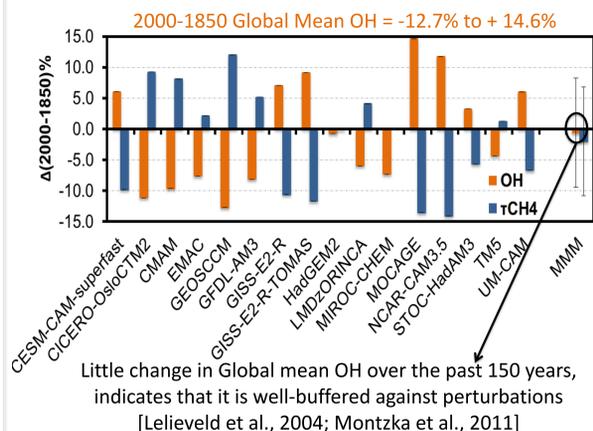
All models have greater OH in the Northern Hemisphere (NH) than the Southern Hemisphere (SH)



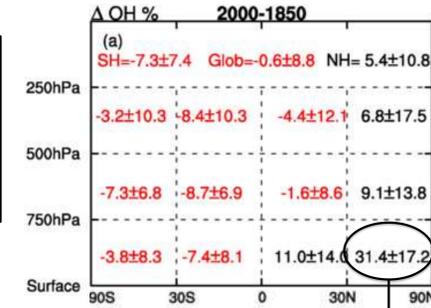
Observational estimates indicate N/S < 1 with 15-30% uncertainties



4. Present day OH relative to preindustrial

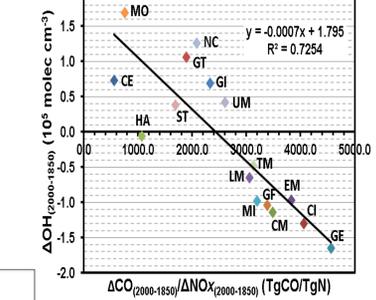


Little change in Global mean OH over the past 150 years, indicates that it is well-buffered against perturbations [Lelieveld et al., 2004; Montzka et al., 2011]



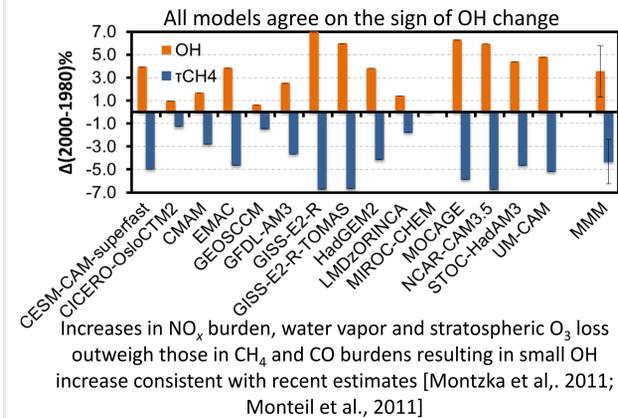
Large uncertainties exist in the magnitude and sign of OH change in different atmospheric subdomains

Different $\Delta CO / \Delta NO_x$ (~changes in OH sinks versus sources) drive diverse preindustrial to present day OH changes across models



Uncertainties in natural emissions and chemical mechanisms contribute to the spread in the interplay between OH sources and sinks

5. Present day OH relative to 1980



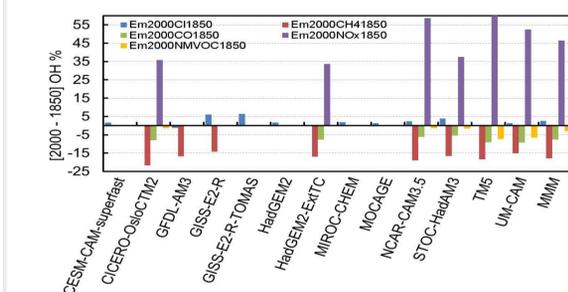
Increases in NO_x burden, water vapor and stratospheric O₃ loss outweigh those in CH₄ and CO burdens resulting in small OH increase consistent with recent estimates [Montzka et al., 2011; Monteil et al., 2011]

6. Impact of Climate Change on Tropospheric CH₄ Lifetime: τ_{CH_4} decreases by about 4 months

Warmer temperatures and higher humidity levels enhance OH-induced loss of CH₄ thereby decreasing its lifetime and contributing a negative climate feedback, consistent with previous studies [e.g. Stevenson et al., 2000]

Models	$\Delta\tau_{CH_4}$ (years)	ΔT (K)	$\Delta\tau_{CH_4}/\Delta T$ (years K ⁻¹)
CESM-CAM-superfast	-0.27	1.4	-0.20
GFDL-AM3	0.12	0.6	0.21
GISS-E2-R	-0.76	1.1	-0.69
GISS-E2-R-TOMAS	-0.70	1.1	-0.64
HadGEM2	-0.20	0.5	-0.40
MIROC-CHEM	-0.25	0.8	-0.30
MOCAGE	-0.20	0.9	-0.23
NCAR-CAM3.5	-0.37	1.1	-0.34
STOC-HadAM3	-0.46	0.6	-0.71
UM-CAM	-0.34	0.6	-0.55
MMM±STD	-0.30±0.24	0.9±0.3	-0.39±0.28

7. Influence of changes in CH₄ and anthropogenic emissions on PI to PD OH changes: NO_x > CH₄ > CO > NMVOCs



8. Future Considerations

- Accurate observational constraints on OH either through direct measurements or using proxies are needed to improve our understanding of OH as represented in models
- Questions remain about the role of aerosols (via chemistry), clouds, NMVOCs (OH recycling in low NO_x, biogenic emission changes), and lightning in driving OH changes

9. References
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